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Experimental measurement of enhanced and hindered particle settling in turbulent gas-particle suspensions, and geophysical implications

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Key Points:

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11	•	We investigate the relationship between particle concentration and particle set-
12		tling velocity in turbulent air-particle mixtures.
13	•	We demonstrate an interplay between hindered settling and cluster-induced en-
14		hanced settling, depending on the Stokes number.
15	•	Clustering of 78 $\mu \mathrm{m}$ particles increases mean settling velocity, while 467 $\mu \mathrm{m}$ par-
16		ticles experience hindered settling.

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17 Abstract

The dynamics of geophysical dilute turbulent gas-particles mixtures depends to a large 18 extent on particle concentration, which in turn depends predominantly on the particle 19 settling velocity. We experimentally investigate air-particle mixtures contained in a ver-20 tical pipe in which the velocity of an ascending air flux matches the settling velocity of 21 glass particles. To obtain local particle concentrations in these mixtures, we use acous-22 tic probing and air pressure measurements and show that these independent techniques 23 yield similar results for a range of particle sizes and particle concentrations. Moreover, 24 we find that in suspensions of small particles (78 μ m) the settling velocity increases with 25 the local particle concentration due to the formation of particle clusters. These clusters 26 settle with a velocity that is four times faster than the terminal settling velocity of sin-27 gle particles, and they double settling speeds of the suspensions. In contrast, in suspen-28 sions of larger particles (467 μ m) the settling velocity decreases with increasing parti-29 cle concentration. Although particle clusters are still present in this case, the settling 30 velocity is decreased by 30%, which is captured by a hindered settling model. These re-31 sults suggest an interplay between hindered settling and cluster-induced enhanced set-32 tling, which in our experiments occur respectively at Stokes number O(100) and O(1). 33 We discuss implications for volcanic plumes and pyroclastic currents. Our study suggests 34 that clustering and related enhanced or hindered particle settling velocities should be 35 36 considered in models of volcanic phenomena and that drag law corrections are needed for reliable predictions and hazard assessment. 37

³⁸ Plain Language Summary

The propagation of turbulent dust storms, snow surge avalanches, dilute pyroclas-39 tic density currents, and volcanic plumes is controlled by the particle concentration, which 40 depends on the particle settling velocity. Most numerical models consider only theoret-41 ical single-particle settling velocities. However, in the presence of neighboring particles, 42 the settling velocity is strongly affected by gas-particle and particle-particle interactions. 43 Our main objective is to unravel the relationship between the particle concentration and 44 the particle settling velocity. We present laboratory experiments, in which the local par-45 ticle concentration is measured using two independent techniques. We observe two set-46 tling mechanisms depending on the coupling of the particles with the gas, namely, hin-47 dered settling or cluster-induced enhanced settling. These mechanisms result in settling 48 velocities significantly different from those of single particles. As an illustrative exam-49 ple, we consider the fall deposit of the 10.5 ka Plinian eruption at Llaima (Chile) and 50 argue that clusters of centimeter-sized clasts had enhanced settling velocities similar to 51 those of 10-20 cm diameter clasts. Our main conclusion is that mechanisms causing en-52 hanced or hindered particle settling velocities should be considered for accurate mod-53 els of geophysical processes, possibly through drag law corrections. 54

55 1 Introduction

Dust storms, snow surge avalanches, dilute pyroclastic density currents, and vol-56 canic plumes are examples of geophysical flows of particles in a turbulent gas. The set-57 tling velocity of particles is essential because it controls the particle concentration and 58 hence the dynamics of these flows. If the particles are suspended by the fluid, the dy-59 namics of the mixture is controlled by the coupling of the fluid and the particles as well 60 as interparticle collisions. Instead, if the particles are not suspended, they settle to form 61 a deposit or a dense basal flow controlled by particle-particle interactions. Experimen-62 tal and numerical studies revealed that the particle settling velocity may increase (Wang 63 & Maxey, 1993; Dev et al., 2019) or decrease (Nielsen, 1993; Fornari et al., 2016; Dev 64 et al., 2019) depending on the degree of turbulence, and it can be lower (Marchetti et 65 al., 2022) or higher (Del Bello et al., 2017) than predicted theoretically for a single par-66

ticle depending on the particle size, shape, and concentration. Actual particle settling
velocities different from theoretical velocities used in models may be a cause of discrepancies between observed and simulated volcanic fall deposits (Tadini et al., 2020, 2022)
and this issue is critical in the context of hazard assessment.

Dilute turbulent gas-particle mixtures are known to exhibit regions of higher par-71 ticle concentrations called clusters, which affect the mixture dynamics and the particle 72 settling velocity (Fullmer & Hrenya, 2017; Brandt & Coletti, 2022). The key mechanisms 73 that lead to particle clustering can be divided into two categories: dissipation of gran-74 75 ular temperature (i.e., fluctuations in particle velocities) due to inter-particle collisions, and relative motion between the particles and the continuous fluid. Only a few numer-76 ical and experimental studies have addressed the effects of clustering on the particle set-77 tling velocity and they revealed that clusters had a fall velocity greater than that of a 78 single particle (Wang & Maxey, 1993; Uhlmann & Doychev, 2014; Breard et al., 2016; 79 Fornari et al., 2019; Fong & Coletti, 2022). In this context, recent experiments showed 80 that the lifespan of clusters in mixtures of particles suspended by an ascending air flow 81 played a fundamental role in setting a critical concentration of particles above which ad-82 ditional particles were not maintained in the mixtures (Weit et al., 2019; van den Wilden-83 berg et al., 2020; Fong & Coletti, 2022). 84

How clustering controls the particle settling velocity and the dynamics of dilute tur-85 bulent geophysical flows remains an open question. The major challenge to addressing 86 this issue experimentally is to obtain reliable values for the solid volume fraction inside 87 3D suspensions and over a range of controlling parameters. This is due the relatively large 88 size of the particles (~100 μ m) and their concentration (~0.1-10 vol.%), which results 89 in scattering of light and makes optical observation complicated (Fong & Coletti, 2022). 90 Therefore, experimental studies are often restricted to very dilute suspensions with par-91 ticle concentrations $\ll 1$ vol.% (Fong & Coletti, 2022; Del Bello et al., 2017) in which 92 particle-fluid (two-way) interactions are expected to dominate. In the present study, we 03 use two independent techniques to investigate particle settling in suspensions with particle concentrations ~ 0.1 -3 vol.%: acoustic probing (van den Wildenberg et al., 2020) 95 and local pressure measurements. We quantify the relationship between the settling ve-96 locity (i.e., the flow velocity necessary to suspend the particles) and the concentration 97 of particles. We find that the settling velocity in suspensions of 78 μ m particles with Stokes 98 number O(1) increases with the local particle concentration and is higher than that of 99 a single particle, hence revealing enhanced settling. In contrast, suspensions of 467 μ m 100 particles with Stokes number O(100) experience hindered settling. We discuss implica-101 tions of our results for geophysical flows, focusing on volcanic mixtures generated by ex-102 plosive eruptions. 103



Figure 1. (a) The experimental setup is a cylindrical pipe of height H=1 m and inner diameter D=4 cm. An air flow is delivered into the container via a wind-box and a porous plate, and a large cylinder at the top of the pipe allows the air velocity to decrease. Interstitial air pressure is measured by two piezo-resistive sensors (PS) placed at heights $Z_1=51.5$ cm and $Z_2=56.5$ cm. Two piezo-electric transducers (S and R) placed at position $Z_3=54$ cm (in between PS_1 and PS_2) are used to measure acoustic wave propagation through the suspension. The volume probed by the local fluid pressure measurements is depicted in dashed light grey, and the volume probed by the acoustic monitoring is depicted in dashed dark grey. (b) Amplitude versus time for four typical (coherent) sound waves measured in different suspensions of glass particles with d=467 μ m and $\phi_L \sim 1$ %. Here t=0 corresponds to the time at which the ultrasonic pulse is sent by the source transducer. (c) Magnitude versus frequency obtained from fast Fourier transforms of the four the sound waves in (b).

104 2 Methods

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2.1 Experimental device and material

All experiments were performed at room temperature in a cylindrical pipe with a height (H) of 1 m and an inner diameter (D) of 4 cm (Fig. 1). At the base of the pipe

is a porous plate and the ensemble is placed on a wind-box. It was shown that the ge-108 ometry of the air inlet is important for the stability of the air flow (Bagheri et al., 2013; 109 Bonadonna & Phillips, 2003) and, therefore, the wind-box has the same inner diameter 110 as the pipe. An air supply system delivered mean air flow velocities (U) through the wind-111 box and the porous plate into the pipe. For particle sizes smaller than 300 μ m, we used 112 the compressed air available in the laboratory and set the air flow using debit controllers 113 that had a precision of less than 1 L/min. For particle sizes larger than 300 μ m, a com-114 pressor (Worthington Creyssensac RLR 35B) was used in combination with a dedicated 115 debit controller. This controller had a minimum increment of 10 L/min (with a preci-116 sion of about 2 L/min). The porous plate had a thickness of 10 mm and a pore size of 117 $20 \ \mu \text{m}$. The influence of the porous plate on the turbulent structures of the airflow is not 118 trivial. We expect that when the air exits the plate the eddy size will be roughly equal 119 to the size of the pores, and the eddy size grows -as the air flows upwards in the pipe-120 to about half the pipe's diameter. 121

At the top of the pipe, a large diameter cylinder ($\sim 2D$) allowed the air velocity 122 to decrease and prevented the loss of particles from the pipe. The pressure of the inter-123 stitial air was measured by two piezo resistive sensors installed at heights $Z_1 = 51.5$ cm 124 and $Z_2 = 56.5$ cm above the porous plate. The sensors were placed in adapters, which 125 were in direct contact with the suspension via a small hole covered by a grid (36 μ m mesh 126 size) to prevent particles entering the adapter and colliding with the sensor. Furthermore, 127 two piezo-electric transducers placed at position $Z_3 = 54$ cm (in between pressure sen-128 sors PS_1 and PS_2) were used to measure acoustic wave propagation through the sus-129 pension. One transducer was used as an acoustic source (S) while the second transducer 130 (R) detected the wave at a distance $L \sim 0.04$ m from the source. To avoid pollution of 131 the acoustic signal by particles colliding against the transducer-face, 36 μ m mesh size 132 grids were installed on the inside wall of the container in front of the transducers. A typ-133 ical experiment is performed as follows. First, a known mass of glass beads is poured into 134 the pipe, and a mean input solid volume fraction ϕ_i is calculated from the mass and the 135 assumption that the beads are homogeneously distributed over the volume of the cylin-136 der. Then, the particles are suspended in air by a vertical upward air flux at velocity U137 and the fluid pressure is measured using the two pressure sensors. Additionally, 400 short 138 pulses are sent by a piezo-electric transducer and the propagating pulses are measured. 139 For each particle size, the experiment is repeated 3 times. We experimented with nearly 140 mono disperse spherical glass particles with density $\rho_2 \sim 2500 \text{ kg/m}^3$ and mean diam-141 eters d of 49, 78, 206, 308 and 467 μ m (sieved between: 45-53, 75-80, 200-212, 300-315, 142 450-475 μ m, respectively). We filmed the suspensions with a high-speed Photron Fast-143 cam SA3 video camera, with resolution of 0.77 mm per pixel, located at 30 cm from the 144 pipe. Movies were acquired at a rate of 250 frames per second and a spatial resolution 145 of 58 x 882 pixels. Improved contrast was achieved by placing the pipe about 50 cm in 146 front of a 40 W white light panel of size 59.5 x 59.5 cm, of which 54.4 x 54.4 cm was the 147 lighting surface. 148

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2.2 Interstitial fluid pressure data

Interstitial fluid pressure arises from air-particle differential motion and the related 150 drag. In each experiment, the generated fluid pressure is measured over time using the 151 pressure sensors that take 500 samples each with a sampling rate of \sim 50 Hz. The bulk 152 solid volume fraction above each sensor may be obtained from the time-averaged pres-153 sure (P) via: $\phi_b = P/(\rho_2 g(H-Z_i))$ (Weit et al., 2019). However, this implies that the 154 suspension is homogeneous, which is not necessarily true. Therefore, in this study we cal-155 156 culate a local solid volume fraction from the pressure measurements of two pressure sensors (PS₁ and PS₂), as follows: $\phi_L = \Delta P/(\rho_2 g \Delta Z)$ with $\Delta P = P_1 - P_2$ and $\Delta Z =$ 157 $Z_2 - Z_1 = 5$ cm (the indices indicate the pressure sensors) and the volume probed is 158 depicted in Fig. 1. From $\phi_L(t)$ it can be observed that a quasi steady state is attained 159 (see Supplementary Information section 1), and only for the smaller particles do we ob-160

Table 1: Symbols and their units

Symbol	Units	Definition
$\overline{C_d}$	-	Drag coefficient for a single particle
d	$\mu \mathrm{m}$	Particle diameter
D	m	Inner diameter of pipe
Н	m	Height of pipe
U	$\rm m.s^{-1}$	Mean air flow velocity
U^*	$\rm m.s^{-1}$	Mean particle settling velocity in the suspension
U_m	$\rm m.s^{-1}$	Air velocity corresponding to ϕ_{Lm}
U_t	$\rm m.s^{-1}$	Terminal settling velocity of a single particle
$U_{cluster^*}$	$\rm m.s^{-1}$	Reduced cluster velocity (velocity in laboratory reference frame)
$U_{cluster}$	$\rm m.s^{-1}$	Cluster velocity $(U_m + U_{cluster^*})$
Re_p	-	Particle Reynolds number
St	-	Stokes number
Т	s	Cluster lifetime
Z	m	Elevation in pipe
α	$\rm m^{-1}$	Attenuation coefficient
β	-	Effective drag coefficient
η	Pa.s	Gas dynamic viscosity
κ_1	$\rm N.m^{-2}$	Air compressibility
κ_1	$\rm N.m^{-2}$	Solid compressibility
ρ_1	$kg.m^{-3}$	Air density
ρ_2	$kg.m^{-3}$	Solid density
ϕ	-	Particle volume fraction
ϕ_c	-	Maximum particle volume fraction that can be suspended
ϕ_i	-	Particle volume fraction assuming particles are homogeneously distributed in pipe
ϕ_L	-	Local particle volume fraction
ϕ_{Lm}	-	Maximum local particle volume fraction
k	$\rm m^{-1}$	complex wavenumber
σ	-	density ratio

serve a small drift due to loss of particles during the experiment. A time averaged ϕ_L may thus be calculated by averaging the 500 samples (~ 10 s).

2.3 Acoustic measurements

In each experiment, we measure the propagation of 400 short acoustic pulses with central frequency f = 0.5 MHz and repetition time of 30 ms. For each transmitted signal, the frequency dependent attenuation coefficient $\alpha(f)$ is determined from the fast Fourier transform (FFT) using: $\alpha(f) = (-1/D) \ln(A_{sus}/A_f)$, with A_f and A_{sus} the amplitudes obtained from the FFT for acoustic signals propagating through a pure fluid phase (air) and the granular suspension, respectively. The attenuation coefficient spectra are averaged over all 400 pulses to obtain a mean attenuation spectrum for each experiment.

We interpret our experimental measurements with two models. For more details we refer readers to van den Wildenberg et al. (2020) and references therein.

(i) The coupled-phase hydrodynamic model developed by Urick (1948) considers a system with a volume fraction ϕ of solid particles with density ρ_2 and compressibility κ_2 , suspended in a viscous continuous phase with density ρ_1 and compressibility κ_1 . In the long wavelength limit ($\lambda/d \gg 1$), the ultrasound attenuation coefficient in such a system is approximated as (Urick, 1948):

$$\alpha = \phi \left[\frac{1}{6} k^4 a^3 + k(\sigma - 1)^2 \frac{s}{s^2 + (\sigma + \tau)^2} \right] + \zeta \tag{1}$$

where a = d/2, the wavenumber $k = \omega/v$ with ω the angular frequency, $v \sim 349$ m/s the sound velocity in air, $\sigma = \rho_2/\rho_1$, $\tau = (1/2) + (9/4)(\delta/a)$, and $s = (9/4)[(\delta/a) + (\delta/a)^2]$. Here $\delta = \sqrt{2\eta/(\omega\rho_1)}$ is a characteristic viscous (shear wave) length, with η the dynamic viscosity of the continuous phase, and ζ is a frequency independent offset (van den Wildenberg et al., 2020).

(ii) The ECAH theory is based on scattering models at the microscopic scale $(\lambda/d \sim$ 183 1) developed by Epstein and Carhart (1953) and Allegra and Hawley (1972). The model 184 describes the interactions of scattered, visco-inertial, and thermal fields with a single par-185 ticle and its surrounding medium (Epstein & Carhart, 1953; Allegra & Hawley, 1972; Ying 186 & Truell, 1956; Challis et al., 2005). As the density contrast between the two phases (air 187 and glass) is very high, we neglect the thermal contribution (Challis et al., 2005). Fi-188 nally we combine the ECAH theory with a multiple scattering theory to correlate the 189 complex wavenumber (k) to the frequency dependent α via: 190

$$\alpha = \zeta + Im \left[\sqrt{k^2 + \frac{3\phi}{a^3} f(0) + \frac{9\phi^2}{4a^6k^2} (F^2(0) - F^2(\pi))} \right]$$
(2)

Here Im signifies the imaginary part and F(0) and $F(\pi)$ are the forward and backward 191 scattering amplitudes of the individual particles, respectively. The scattering amplitudes 192 are calculated from the scattering coefficients of the various types of waves scattered from 193 an individual particle, and we consider scattering and viscous/inertial losses and neglect 194 thermal conduction losses. ϕ is then calculated from the experimental acoustic data by 195 fitting the theoretical α from the models using Matlab's least-squares curve fitting tools. 196 In the dilute suspensions considered here, the additional free fitting parameter $\zeta \sim 0$ -197 1. 198

¹⁹⁹ 3 Results

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3.1 Acoustic probing and local fluid pressure measurements

We start by comparing the local solid volume fraction ϕ_L obtained by local fluid 201 pressure measurements and acoustic attenuation measurements. To do so, we introduce 202 a known mass of particles (d=49, 78, 206, 307 or 467μ m) into the container and choose 203 U so that ϕ_L obtained from the fluid pressure is about 1%. Then, we measure the prop-204 agation of 400 acoustic pulses. The similarity between the acoustic pulse signals demon-205 strates the reproducibility of the sound propagation through the suspension (Fig. 1b) 206 and spectral analysis of the sound signals evidences a single peak around $f_c \sim 0.5$ MHz 207 (Fig. 1c). The attenuation coefficient α , calculated from the spectral analysis, increases 208 with the frequency (Fig. 2a). In order to infer ϕ_L we compare $\alpha(f)$ with existing the-209 oretical models. For the particles smaller than 200 μ m, both the Urick and ECAH mod-210 els predict well the experimental data (Fig. 2a). In contrast, for the larger particles, only 211 the ECAH model appears to be able to describe our data (Fig. 2a). For each particle 212 size, a mean ϕ_L is calculated from the pressure measurements and from the acoustic mon-213 itoring by averaging 3 experiments. Importantly, we find overall a good agreement be-214 tween ϕ_L obtained by the local fluid pressure and acoustic methods (Fig. 2b). 215



Figure 2. Acoustic wave attenuation in suspensions of glass beads. (a) Example traces of α versus frequency for suspension of glass particles in air (blue: $d = 49 \ \mu\text{m}$, orange: $d = 78 \ \mu\text{m}$, yellow: $d = 206 \ \mu\text{m}$, green: $d = 308 \ \mu\text{m}$, purple: $d = 467 \ \mu\text{m}$). The dotted lines represent the fits of $\alpha(f)$ by the Urick model (Eq.1). The dashed lines represent the fits of $\alpha(f)$ by the ECAH44 model (Eq.2). (b) ϕ_L averaged over 3 experiments as a function of the particle diameter. Solid dots represent ϕ_l obtained from the ECAH44 model, open dots represent ϕ_L obtained from the Urick model, and triangles represent the ϕ_L obtained from the local fluid pressure. The error bars show the standard deviations over the experiments. Colors correspond to those in (a).

To test the applicability of these techniques over a wider range of particle concen-216 trations, a mass of particles of 160.08 g (corresponding to $\phi_i = 5\%$) is introduced into 217 the container at U=0. Subsequently U is slowly ramped-up and the acoustic wave prop-218 agation and the fluid pressure are measured. We use glass beads of size $d=78 \ \mu m$ and 219 467 μ m, and the experiment is repeated 3 times. Traces of $\alpha(f)$ reveal that the sound 220 attenuation increases with U as more particles enter the volume probed by the acous-221 tic wave (Fig. 3a). Since the ECAH model appears to work for both particle sizes, we 222 will use this model to infer ϕ_l from $\alpha(f)$. As before, we determine a mean ϕ_L for the pres-223 sure measurements and for the acoustic monitoring by averaging over all experiments. 224 Figure 3b shows ϕ_L as a function of U. Overall, we find that ϕ_L obtained by the two tech-225 niques are similar for all the concentrations tested. For the particles $d=78 \ \mu m, \phi_L$ is 226 first nearly 0 for the lowest flow velocities, and above $U \sim 0.4$ m/s it increases until a 227 maximum of $\phi_L \sim 0.5\%$ is reached at $U \sim 0.58$ m/s, above which ϕ_L decreases as the 228 particles are dragged above the position of the sensors. For the suspensions of larger par-229 ticles $d = 467 \ \mu m$, we observe a similar behavior, but for higher U. One important ob-230 servation is that, even though the ϕ_i is the same (5%), the maximum ϕ_L obtained dif-231 fers significantly for the two particle sizes. For the $d=467 \ \mu m$ particles we find that $\phi_L \sim$ 232 $4\% \ (\sim \phi_i)$, while for the $d=78 \ \mu m$ particles $\phi_L \sim 0.45\%$ almost ten times smaller than 233 ϕ_i . 234



Figure 3. (a) Attenuation coefficients versus frequency in a suspension of $d=78 \ \mu m$ glass beads for increasing U. (b) ϕ_L obtained from fits of $\alpha(f)$ in (a) by the ECAH44 model (black dashed lines in a). The black triangles are ϕ_L obtained from local fluid pressure. (c) Same as (a) for $d=467 \ \mu m$ glass beads. (d) ϕ_L obtained from fits in (c) by the ECAH44 model (black dashed lines in c). The black triangles are ϕ_L obtained from the local fluid pressure. Error bars represent the standard deviations over three experiments.

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3.2 Local solid volume fraction and clustering

Next, we investigate the relationship between the maximum local solid volume frac-236 tion ϕ_{Lm} , the air velocity, and the particle size. Experiments are repeated for varying 237 ϕ_i and for simplicity we use only fluid pressure measurements to calculate ϕ_L . The res-238 olution in the experiments is limited by the minimal increments of the flow meter, from 239 which the error bars in U are determined. In suspensions of 78 μ m particles, we find that, 240 for a given ϕ_i , ϕ_L increases with U to ϕ_{Lm} at U_m , after which it decreases rapidly (Fig. 241 4a). Importantly, for increasing ϕ_i , both the ϕ_{Lm} and U_m increase (Fig. 4c,d). For the 242 suspensions of 467 μ m particles, a similar behavior is found for ϕ_L and ϕ_{Lm} , with ϕ_L 243 increasing with U until its maximum value ϕ_{Lm} at U_m , and increasing ϕ_i leads to a higher 244 ϕ_{Lm} (Fig. 4b,e). However, in contrast to the smaller particles, U_m appears to decrease 245 with increasing ϕ_i (Fig. 4f). 246



Figure 4. ϕ_L calculated from local pressure measurements and presented as a function of the air velocity U. (a) Suspensions of $d=78 \ \mu m$ glass for increasing $\phi_i \ [\phi_i=0.2\%$ (yellow), 1% (light orange), 5% (dark orange), 10% (red)]. The dashed arrow illustrates the shift of the curves to the right for increasing ϕ_i . (b) Suspensions of $d=467 \ \mu m$ glass beads $[\phi_i=0.5\%$ (magenta), 1% (dark purple), 2.5% (light purple), 5% (blue)]. The dashed arrow illustrates the shift of the curves to the left for increasing ϕ_i . (c) The maximum local concentration ϕ_{Lm} versus ϕ_i from the data in (a). The line represents $\phi_{Lm} = \phi_i$. (d) The air velocity U_m corresponding to ϕ_{Lm} versus ϕ_i from the data in (b), the line represents $\phi_{Lm} = \phi_i$. (f) The air velocity U_m corresponding to ϕ_{Lm} versus ϕ_i from the data in (b). Error bars of ϕ are the standard deviations over the experiments, error bars in U_m are the increment in the air velocity (for the $d=78 \ \mu m$ particles the error bars of U_m are the size of the symbols).

We filmed each suspension at ϕ_{Lm} and the corresponding U_m , focusing on the region between PS₁ and PS₂. Frames from those movies are presented in Fig. 5 and show that regions with high particle concentrations (clusters) are present in all the suspensions, except in those with the lowest ϕ_{Lm} . This supports the idea that cluster lifespan plays an important role in the maximal concentration of particles that can be suspended (ϕ_c) as well as in the formation of dense beds at the bottom of the suspension (van den Wildenberg et al., 2020; Fong & Coletti, 2022).



Figure 5. Frames from movies of suspensions with varying ϕ_i at the relevant U_m . The movies are focused on the region between pressure sensors PS₁ and PS₂. (a)-(d) suspensions of 78 μ m particles with $\phi_{Lm} = 0.09, 0.29, 0.54 0.72\%$, respectively. (e)-(h) suspensions of 467 μ m particles with $\phi_{Lm} = 0.11, 0.64, 1.6, 3\%$, respectively. Particles are black against a white background, except for in (a) where the particles are white against a black background. Denser regions of particles are outlined in red and the arrows indicate the direction of cluster movement.

A detailed investigation of the size, shape, and particle concentration of clusters 254 is beyond the scope of this paper due to the spatial resolution of the images and is fur-255 ther complicated by the limited penetration of the light through the suspension. We thus 256 provide a phenomenological description of the suspensions. Clusters in the suspensions 257 of small particles tend to be elongated and relatively concentrated by comparison to clus-258 ters observed in the suspension of larger particles which are more rounded and less con-259 centrated (Fig.5). In circulating fluidized bed risers (with higher particle concentrations 260 compared to our experiments), cluster size shows no systematic variations with the par-261 ticle size nor the riser size (Cahyadi et al., 2017; Berruti et al., 1995).Instead, it has been 262 shown to, in some cases, vary with the radial position due to variations in radial veloc-263 ity as particles go up in the middle and down on the sides (Cahyadi et al., 2017; Berruti 264 et al., 1995). Our experiments show similar cluster trajectories (Supplementary Infor-265 mation section 4 and supplementary movies), however, variations in cluster size remain 266 elusive. These are qualitative impressions and care should be taken when interpreting 267 2D projections of 3D structures. 268

In order to accurately determine the cluster fall velocity $(U_{cluster})$, it is necessary 269 to consider long trajectories. Therefore, we filmed the suspensions of $d = 78 \ \mu \text{m}$ par-270 ticles at $\phi_L = 0.7\%$ and $U_m = 0.66$ m/s, and $d = 467 \ \mu m$ particles at $\phi_L = 3\%$ and 271 $U_m = 2.5 \text{ m/s}$, imaging the whole pipe at a frame rate of 250 Hz (see Supplementary 272 movies). The duration of those movies corresponds to the time over which a pressure mea-273 surement was performed. Small particles are lost during the time of the movie, which 274 results in progressively lighter images. However, this loss is small and does not change 275 significantly ϕ_L . To study the motion of the clusters we first produce a kymograph, which 276

is an intuitive way to represent the dynamics of the clusters during the movie in a sin-277 gle image. For each image of the movie, we plot the light intensity along a line in the 278 Z-direction near the center of the pipe (Fig. 6). Clusters that ascend and descend are 279 depicted in the kymograph as trajectories with a positive slope and a negative slope, re-280 spectively. The trajectories of these clusters show that their motion is mostly smooth 281 and continuous. We thus determine the slope and length of the trajectories correspond-282 ing to the settling velocity $U_{cluster^*}$ and lifetime T of the clusters. Note that $U_{cluster^*}$ 283 is a reduced settling velocity owing to the ascending air flow. We use ImageJ and the 284 available plugin *RidgeDetection* (Steger, 1998), which is an algorithm developed for de-285 tecting ridges and lines. We consider trajectories with a minimal length of 70 frames (cor-286 responding to 0.28 s), so that their slope is well defined. We find that the identified tra-287 jectories converge for 78 μ m particles with parameters linewidth= 3, lower threshold = 288 0.2 and upper threshold = 1.4. For 467 μ m particles, a linewidth=3, lower threshold = 289 0.2 and upper threshold=7.1 are used. For each movie, we track about 140 descending 290 clusters and show the results in histograms in Fig. 6b,c. Descending clusters appear sig-291 nificantly more numerous (more than 70% of the clusters descend, see Fig.6e) and in gen-292 eral we have the impression that they are more concentrated than ascending clusters (Sup-293 plementary movies). The average lifetime of clusters is about 0.6 s and is the same in sus-294 pensions of 78 μ m particles and in suspensions of 467 μ m particles. However, the mean 295 settling velocity of descending clusters, as determined from averaging all positive $U_{cluster^*}$ 296 for a given particle size, in suspension of 78 μ m particles is $U_{cluster^*} \sim 0.9$ m/s, while 297 in suspensions of 467 μ m particles we find that $\bar{U}_{cluster^*} \sim 1.1$ m/s. 298



Figure 6. Kymographs of clusters in suspensions of (a) $d=78 \ \mu m$ particles, and, (b) $d=467 \ \mu m$ particles. (c and d) Zoom of the kymograph in (a) and (b), trajectories of clusters are shown by the red dashed lines. (e) Histogram of the reduced velocities of tracked clusters. Negative velocities correspond to ascending clusters, while positive velocities correspond to descending clusters. Most of the tracked clusters descend. (f) Histogram of lifetime (T) of descending clusters. Red and blue (transparent) colors correspond to suspensions of 78 μm and 467 μm particles, respectively.

²⁹⁹ 4 Discussion

300

4.1 Acoustic probing and local fluid pressure measurements

The solid volume fraction in particle suspensions is difficult to measure experimentally. Our first main result is the comparison of two independent approaches to determine the local solid volume fraction (ϕ_l) in dilute but optically opaque suspensions of particles in air, namely (i) local fluid pressure measurements and (ii) acoustic probing. These techniques yield similar local solid volume fractions for a wide range of particle

sizes and particle concentrations. They thus present tools for investigating the solid vol-306 ume fraction in laboratory experiments. Local fluid pressure measurements have the ad-307 vantage that they offer ready access to the local solid volume fraction, without the need 308 of many additional parameters such as, for example, the particle size or the viscosity of 309 the fluid. Its main limitation is that it can be used only when there is a significant in-310 terstitial fluid pressure. Instead, acoustic monitoring can be used in the absence of fluid 311 pressure. It is, however, more laborious to apply and needs more physical parameters, 312 which may not be known. Moreover, in suspensions of particles larger than 200 μ m the 313 Urick model yields poor fits with unreasonably low ϕ_L indicating that this model fails. 314 Under the assumption that the viscosity/inertia is responsible for the acoustic energy 315 loss, only the second term in Eq. 1 (viscous-inertial loss) may be used (Pierce et al., 2016). 316 Indeed, the use of only this term captures our data $\alpha(f)$ for the larger particles. How-317 ever, the inferred $\phi_L \sim 3.4\%$ is larger than the $\sim 1\%$ obtained from the ECAH model 318 and from local pressure measurements (Supplementary Information section 2). Further 319 work is needed to address this issue. 320

321

4.2 Local solid volume fractions and particle settling velocities

The second main result is the dependence of U_m on the maximum local particle concentration ϕ_{Lm} , as observed in Fig. 4a&b, and for clarity we present cross plots in Fig. 7a&c. For the 78 μ m particles, U_m increases rapidly with increasing ϕ_{Lm} (Fig.7a), indicating that these particles settle faster at higher particle concentrations. Instead, for the 467 μ m particles, U_m decreases with increasing ϕ_{Lm} (Fig.7c), indicating that these larger particles settle slower at higher concentrations. In the following sub-sections we will discuss these findings in detail.



Figure 7. (a) Cross-plots of U_m versus ϕ_{lm} for suspensions of 78 μ m particles (data from Fig. 4a). The solid black curve is a power law fit. (b) Effective drag coefficient β calculated from (a) versus the Re_p . The black curve is the estimated drag on a single particle via Abraham's contribution. (c) Cross-plot of U_m versus ϕ_{lm} (data from Fig. 4b) for suspensions of 467 μ m particles. Solid black curves are fits of $U_m = p_1(1 - \phi_{Lm})^{p_2}$ and theoretical predictions using $U^* = p_1(1 - \phi)^n$ with $n = (5.1 + 0.27Re_p^{0.9})/(1 + 0.1Re_p^{0.9})$ (dashed curve). (d) β calculated from (c) versus the Re_p . The black curve is the estimated drag on a single particle via Abraham's contribution (see main text).

4.2.1 Air flow regime

329

The hydrodynamics of the macroscopic air flow in our device is described by the Reynolds number $Re = \rho_1 U_m D/\eta$. For the suspensions of 78 μ m particles at $U_m \sim$ 0.6 m/s we obtain $Re \sim 1500$, corresponding to a laminar flow, while for the suspensions of 467 μ m particles at $U_m \sim 2.7$ m/s we obtain $Re \sim 6800$, corresponding to a turbulent flow. Note that in the experiments, the presence of particles in the fluid significantly complicates the flow dynamics with respect to those in pure fluid (Tsuji et al., 1984; Jessop & Jellinek, 2014).

The motion of a spherical particle falling through a fluid is controlled by two di-337 mensionless numbers, $\sigma = \rho_2/\rho_1$ and the particle Reynolds number Re_p . The local flow 338 around the particle may be estimated via $Re_p = \rho_1 dU/\eta$, where $\rho_1 = 1.13 \text{ kg/m}^3$ and 339 $\eta = 1.81 \times 10^{-5}$ kg/(m s) are the density and dynamic viscosity of air, respectively. For 340 the suspension of 78 μ m particles, with U_m varying from 0.4 to 0.7 m/s, this gives a $Re_p \sim$ 341 1.9-3.4. For the suspensions of 467 $\mu{\rm m}$ particles, taking $U_m\sim$ 2.5-3 m/s, this gives Re_p 342 ranging from 73-88. The obtained Re_p indicates that all the suspensions studied here 343 are in the transitional flow regime $(1 < Re_p < 1000)$ (Rhodes, 2008). 344

345 4.2.2 Effective drag coefficient β

Consider a single particle in a suspension and suppose that U_m is equal and opposite in direction to its fall velocity. Ignoring fluid inertia, there are three forces acting on the particle, a buoyancy force, a gravitational force and a drag force F_d , so that: $\pi d^3(\rho_2 - \rho_1)g/6 = F_d$, where $F_d = (1/2)\rho_1 A U_m^2 \beta$, with A the surface of the sphere on which the drag acts and β the effective drag coefficient. The drag coefficient is calculated from

$$\beta = dg(4/3)(\rho_2/\rho_1 - 1)/U_m^2 \tag{3}$$

and shown as a function of Re_p in Fig. 7b,d. We find that β is a decreasing function of Re_p similar to the drag coefficient for single particle, C_d , estimated by the Abraham contribution in the transitional flow regime $C_d = [0.5407 + (24/Re_p)^{1/2}]^2$ (Abraham, 1970). Additionally, we find that for suspensions of small particles, $\beta < C_d$, while for suspensions of larger particles, $\beta > C_d$. This suggests that different mechanisms control the settling dynamics in suspensions of smaller particles and larger particles, as discussed below.

4.2.3 Hindered settling

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Particles falling in a particle-laden fluid experience a larger fluid drag than a par-360 ticle falling in a pure fluid due to the flow around neighbouring falling particles, an ef-361 fect called hindered settling (Richardson & Zaki, 1954; Yin & Koch, 2007; Dey et al., 362 2019). Consequently, the mean fall velocity in a homogeneous suspension of spherical 363 particles is less than the terminal velocity of a single isolated particle. The most widely 364 used empirical relation between the hindered settling velocity U^* and volumetric par-365 ticle concentration ϕ comes from sedimentation and liquid fluidization experiments, which 366 showed that $U^* = U_t (1 - \phi)^n$ (Richardson & Zaki, 1954), where U_t is the theoretical 367 settling velocity of a single isolated particle in a pure fluid and n is a power-law expo-368 nent that depends on the particle Reynolds number $Re_p(U_t)$. For the range of Re_p in 369 this study, the relevant hindered settling model includes a prefactor k yielding $U^* = kU_t(1-$ 370 ϕ)ⁿ (Yin & Koch, 2007; Di Felice, 1999) and here $n = 4.4 Re_p^{-0.1}$ (Richardson & Zaki, 371 1954). In order to test such a settling model, the data of U_m and ϕ_{Lm} are compared to 372 an empirical expression $U_m = p_1(1 - \phi_{Lm})^{p_2}$ which has a similar functional form to 373 the hindered settling model. 374

For suspensions of 467 μ m particles, the data are well described with $p_1 = kU_t =$ 2.9 ± 0.1 m/s and $p_2 = 4.89 \pm 2.5$ (Fig. 7c). For a single particle, $U_t \sim 3.5$ m/s, which implies that $k \sim 0.8$, in agreement with literature values (Yin & Koch, 2007; Di Felice, 1999). Moreover, the exponent obtained from the fit is very close to the improved relationship $n = (5.1 + 0.27 * Re_p^{0.9})/(1 + 0.1 * Re_p^{0.9})$, where Re_p is calculated from the theoretical settling velocity of a single particle (Garside & Al-Dibouni, 1977).

Hindered settling cannot account for the settling velocity increasing with the solid 381 volume fraction in suspensions with 78 μ m particles. Instead, this dependence may be 382 captured by power law $U_m = p_3 \phi_{Lm}^{p_4} + p_5$, with $p_3 = 2804$ m/s, $p_4 = 1.9 \pm 0.6$. Here 383 $p_5 = 0.47 \pm 0.02$ m/s represents the single particle settling velocity in the limit of very 384 low ϕ , which is higher than the predicted theoretical velocity for a single particle $U_t \sim$ 385 0.37 m/s in the intermediate flow regime (Fig. 7a). A power law relation between the 386 solid volume fraction and the settling velocity was also observed for volcanic ash par-387 ticles (Del Bello et al., 2017). On the basis of this comparison, we argue that the set-388 tling dynamics in suspensions of small particles is to large extent controlled by the for-389 mation of particle clusters as discussed next. 390

391 4.2.4 Clustering

It is difficult to quantify the particle concentration in, and shape of, clusters due 392 to the extensive light scattering by the particles, and the additional problem of inter-393 preting a 2D projection of a 3D structure. Not withstanding these limitations, the clus-394 ter velocity appears to be well defined in the suspensions (Fig. 6 and supplementary movies). 395 Analysis of the cluster trajectories suggests that the mean reduced settling velocity of 396 clusters (i.e., $U_{cluster*}$ in the laboratory reference frame) is slightly lower in suspensions 397 of 78 μ m particles than in suspensions of 467 μ m particles. However, the cluster veloc-398 ity for settling in static air is $U_{cluster} = U_{cluster*} + U_m$ yielding $U_{cluster} \sim 0.9 + 0.7 =$ 399 1.6 m/s and $U_{cluster} \sim 1.1+2.5=3.6$ m/s in suspension of 78 μ m and 467 μ m particles, 400 respectively (Fig. 8). 401

To quantify the contribution of the clusters to the overall settling velocity of the mixture, we evaluate the ratio $U_{cluster}/U_t$. For the suspensions of 78 μ m particles we find $U_{cluster}$ is about 4 times larger than U_t , resulting in a significant impact of clusters on the settling dynamics of the suspension (Fig. 8a). Instead, for the 467 μ m particles the $U_{cluster} \sim U_t$, making it possible for hindered settling to dominate (which is a settling mechanism for isolated particles) (Fig. 8b).

The picture that emerges is the following. For suspensions of 78 μ m particles, the 408 bulk $Re \sim 1500$ implies a laminar flow. However, spatial fluctuations in the particle con-409 centration may generate turbulence, an effect called cluster induced turbulence (CIT) 410 (Capecelatro et al., 2015; Fong & Coletti, 2022). CIT occurs when the mass loading $\psi =$ 411 $\sigma \phi/(1-\phi) \gg 1$, which is the case here as $\psi \sim 10$. The elongate and apparently dense 412 clusters have a higher settling velocity than individual particles, resulting in the increase 413 of U_m with ϕ_{Lm} . In contrast, for the suspensions of larger particles, the bulk $Re \sim 6800$ 414 implies turbulent flow. The decrease of U_m with increasing ϕ_{Lm} is well captured by a 415 hindered settling model (Fig. 7c). Significant clustering is also observed in these suspen-416 sions, however, as the settling velocity of the clusters is not significantly different than 417 that of a single particle, the clusters do not appear to impact the hindered settling mech-418 anism. In other words, in suspensions of small particles the clusters dominate the set-419 tling dynamics, whereas the settling in suspensions of larger particles is dominated by 420 hindered settling. The impact of clusters on particle settling may be related to the dif-421 ferent shape, particle concentration, and longevity of the clusters, which in turn may de-422 pend on how particles are coupled to fluid motion (Nitsche & Batchelor, 1997; Pigna-423 tel et al., 2011; Chen et al., 2021). 424

⁴²⁵ The Stokes number captures the ability of particles to respond to changing flow ⁴²⁶ conditions and is the ratio of the particle response time τ_p to a characteristic flow time ⁴²⁷ scale τ_f (for the largest eddies $\tau_f = D/U_m$), so that

$$St = \frac{(\rho_2 - \rho_1)d^2 U_m}{18\eta D}.$$
(4)

Clustering is sensitive to St (Bec, 2003). When $St \ll 1$ particles are well coupled to the 428 flow and follow streamlines. When $St \gg 1$ particle motion is decoupled from turbulent 429 motions. When St is of order 1, particles tend to accumulate at the edges of eddies (Maxey, 430 1987; Bec et al., 2007), leading to a local increase in particle concentration, a process that 431 can initiate cluster formation. Our 78 μ m particles at $U_m \sim 0.7$ m/s have St= 0.8, and 432 the $d = 467 \ \mu m$ particles with $U_m \sim 2.5 \ m/s$ have St= 105 with the above definition 433 of St, meaning that suspensions of smaller particles are more prone to clustering. Clus-434 ters of small particles may thus last longer and have higher particle concentrations, and 435 consequently their $U_{cluster}$ and $U_{cluster}/U_m$ are larger than for clusters of high St par-436 ticles. As a result, clustering may have a dominant role for the smaller particles, while 437 hindered settling prevails for the larger particles. Future experiments and additional prob-438 ing of particle concentration are needed to disentangle the role of the particle size and 439

the velocity profile on the persistence of clusters as well as cluster shape and concentra-tion.

4.3 Geophysical implications

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Our experiments show that the consequence of clusters on the settling velocity de-443 pends essentially on the Stokes number. Here, relative to the settling velocity of isolated 444 particles U_t , our Stokes number O(1) particles settle with a mean velocity twice as large, 445 and clusters of these particles have settling velocities that are even four times as large. 446 However, our larger Stokes number $O(10^2)$ particles experience hindered settling, with 447 settling speeds up to 30% slower than U_t , while clusters in these suspensions have set-448 tling velocity roughly equal to U_t . These types of phenomena that arise from 4-way cou-449 pling are not accounted for in some models of particle transport used to simulate vol-450 canic plumes, such as Ash3D (Schwaiger et al., 2012), TEPHRA (Bonadonna et al., 2005), 451 FALL3D (Folch et al., 2020) and PLUME-MOM/HYSPLIT (Stein et al., 2015) that con-452 sider theoretical single-particle settling velocities. Modifications to the particle drag and 453 settling velocity (Fig. 7) translate into additional uncertainty in such models. Settling 454 velocities lower or higher than that of the theoretical velocities for single particles are 455 a possible cause of discrepancy between simulated and observed fall deposit character-456 istics such as grain size and mass of deposit per unit area (Tadini et al., 2020). As high-457 lighted in the report by the National Academies of Sciences, Engineering and Medicine 458 (2017), particle-scale processes influence the large scale dynamics of volcanic eruptions. 459 Similarly, in their review of pyroclastic density currents, and relevant to geophysical tur-460 bulent gas-particles mixtures in general, Lube et al. (2020) wrote "drag law correction 461 and sub-grid models that account for mesoscale clustering must be developed". 462



Figure 8. Schematic of the settling velocities (in m/s). (a) Suspensions of $d=78 \ \mu m$ particles. Left panel, settling velocities in the presence of an upward air flow (i.e., velocities in the laboratory reference frame). The mixture is suspended, while clusters settle at a reduced velocity. Right panel, settling in the absence of an air flow. The mixture settles with a velocity $U_{mix} \sim U_m$ and clusters settle at velocities $\sim U_m + U_{cluster^*}$. For comparison the terminal settling velocity U_t of a single particle in the intermediate regime is also illustrated. Note that this velocity is significantly less than the settling velocity of the mixture and the clusters. (b) Suspensions of $d=467 \ \mu m$ particles. In this case, the settling velocity of the mixture is less than the theoretical settling velocity of a single particle in the intermediate flow regime. (c) Dacite pumice fall from the 10.5 ka Plinian eruption of Llaima, Chile (Northing 570558, Easting 0271863; photo taken 7 January 2020).

463 One fall deposit characteristic from volcanic plumes that might result from clus tering, and be diagnostic that clustering is an important process in spreading umbrella
 clouds, is the occurrence of unusually large clasts in an otherwise well-sorted deposit. As
 an illustrative example, we consider the dacite pumice fall from the 10.5 ka Plinian erup-

tion of Llaima, Chile (Schindlbeck et al., 2014). Fig. 8c shows unit IIIb of the deposit, 467 12 km from the vent. Notable is the outsized clast in a matrix of smaller pumice (most 468 of the deposit consists of cm size lapilli with a small proportion of large clasts). The in-469 termediate diameter of the large clast, measured with a ruler, is 16 cm. The median di-470 ameter of the pumice, measured by dry sieving in the lab, is 1 cm with an Inman sort-471 ing value of 0.45 (see Supplementary Information section 3 for details). 12 km is too dis-472 tal for the large clast to be a ballistic bomb. Our hypothesis is that within the spread-473 ing umbrella cloud, the lapilli settled in clusters with a similar settling velocity as the 474 large particle since they were deposited together. We assume that U_m is equal to the fall 475 velocity of the coarsest particles. For the large and median particle sizes, diameters of 476 16 and 1 cm respectively, the theoretical settling speeds assuming a density of 600 kg/m^3 477 are ~ 61 m/s and ~ 14 m/s using the Abraham (1970) drag model. This implies a roughly 478 4-fold increase in settling speed for the clusters of small particles, which is similar to the 479 enhanced settling speed in our experiments (Fig. 8) and mesoclusters in large scale den-480 sity current experiments (Breard et al., 2016). 481

More generally in volcanic plumes and spreading umbrella clouds, enhanced (or hin-482 dered) settling can cause deposits with mass per unit area larger (or smaller) and grain 483 sizes smaller (or larger) than expected for single size particles (Tadini et al., 2020, 2022; 484 Freret-Lorgeril et al., 2020; Gilchrist & Jellinek, 2021). Clustering in mixtures contain-485 ing particles with Stokes number O(1) is an alternative to ash aggregation commonly 486 assumed in simulations to explain enhanced settling (e.g., Tadini et al., 2022). The fact 487 that bimodal fallout deposits, such as the one at Llaima (Fig. 8), are relatively rare sug-488 gests effective particle size segregation occurs near the vent in most cases. On the other 489 hand, hindered settling accounts for reduced particle settling velocity, otherwise explained by non spherical particle shape or the presence of wind (Marchetti et al., 2022). In a broader 491 context, hindered settling is likely to occur in various types of geophysical flows contain-492 ing particles with Stokes number $\gg 1$, and our experiments suggest that this mechanism 493 can operate at concentrations as low as ~ 1 vol.%, which may be surprising. 494

Numerical simulations of pyroclastic currents showed that unrealistically high par-495 ticle drag coefficients, at least 30 times larger than commonly estimated for typical vol-496 canic particles ($\sim 0.5-1$), were required to match the runout distance of the ash-cloud 497 surges at Merapi, 2010 (Kelfoun et al., 2017) and Mount Pelée (Martinique), May 8th, 498 1902 (Gueugneau et al., 2020). To explain the high drag coefficient, Gueugneau et al. 499 (2020) concluded that hindered settling occurred as a consequence of the increase in par-500 ticle concentration towards the lower part of the currents, our experimental findings sup-501 port this statement. Even though it is unlikely that hindered settling alone can account 502 for the large discrepancy between their simulation and the natural case, it may be partly 503 responsible. Our experiments showing hindered settling provide insights into the pos-504 sible settling characteristics of the particles in the basal layer of these natural flows, that 505 is, with Stokes numbers $\gg 1$ and at concentrations as low as ~ 1 vol.%. 506

507 5 Conclusion

We measured the local solid volume fraction in dilute suspensions of glass parti-508 cles in air using two independent techniques, acoustic probing and local fluid pressure 509 measurements. We showed that these techniques yield similar solid volume fractions, thus 510 providing two attractive tools to investigate optically opaque mixtures in analogue lab-511 oratory experiments. We anticipate that a combination of these techniques may be use-512 ful if, for example, both the particle concentration and particle size are not known. In-513 deed, the local pressure measurements may then be used to obtain a solid volume frac-514 tion, which could be used in the sound attenuation calculation to obtain the particle size. 515

⁵¹⁶ Our experiments in a vertical pipe demonstrate that the local particle concentra-⁵¹⁷tion in dilute suspensions depends on the initial amount of particles, the particle size, and the air flow velocity. In suspensions of 78 μ m particles, higher air velocities were needed to obtain higher local particle concentrations and we argue this was due to the presence of clusters that settled faster than individual particles. In contrast, in suspensions of 467 μ m particles, higher local particle concentrations were achieved at lower air velocities as a consequence of hindered settling of the particles, despite the presence of clusters. Differences in the shape and particle concentration of the clusters may explain their different settling behaviours.

Our results have implications for modeling geophysical dilute gas-particle mixtures. The field example we discuss suggests that in volcanic plumes clusters of relatively small particles have enhanced settling velocities, similar to that of larger clasts, decreasing the sorting in fall deposits. Other geophysical flows with particles of broad grain size ranges are likely to exhibit similar flow dynamics. We conclude that the formation of clusters and their settling velocities being significantly different than that of single particles should be considered in models of geophysical processes through drag law corrections.

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539 7 Open Research

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