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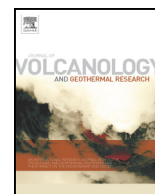
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Tephra characterization and multi-disciplinary determination of Eruptive Source Parameters of a weak paroxysm at Mount Etna (Italy)

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ABSTRACT

The determination of Eruptive Source Parameters (ESPs) is a major challenge especially for weak volcanic explosions associated with poorly exposed tephra-fallout deposits. In such a case, the combination of deposit analyses and remote sensing observations can provide fundamental insights. We use the 29 August 2011 weak paroxysm at Mount Etna (Italy) as a case study to discuss some of the challenges associated with multi-disciplinary determination of ESPs of poorly exposed tephra-fallout deposits. First, we have determined the erupted mass from a combination of field and synthetic data to fill a significant gap in data sampling; synthetic data have been derived based on extrapolation of field observations and validated based on comparisons with other tephra deposits at Etna and TEPHRA2 modelling. Second, we have combined the estimates of erupted mass and grain-size distribution as derived both from deposit observations and satellite retrievals. Analytical modelling was applied to characterize the size fractions most likely represented in satellite retrievals and tephra deposits, respectively. In addition, the Rosin-Rammler distribution fitting is shown to inform on missing parts of the grain-size distributions and reproduce a tail of very fine ash (1–20 μm) whose mass proportion is close to the satellite estimates (1.3–1.6% versus 1.9%, respectively). Finally, it was found that this very-fine-ash fraction increases as a function of satellite-derived Mass Eruption Rate for a set of eruptions for which independent estimates are available. This critical combination of field observations, analytical modelling and satellite retrievals demonstrates the potential and importance of multidisciplinary strategies for the derivation of ESPs even for small-size explosive events and poorly exposed deposits such as that of the 29 August 2011 paroxysm of Mt. Etna.

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1. Introduction

The determination of Eruptive Source Parameters (ESPs) (e.g., the vent location, Total Erupted Mass - TEM, Mass Eruption Rate - MER, Plume Height - H_p , Total Grain-Size Distribution - TGSD) is crucial to understand, assess, and forecast the hazards associated with volcanic eruptions (e.g., Costa et al., 2006; Scollo et al., 2008; Folch, 2012; Bonadonna et al., 2015; Beckett et al., 2020; Aubry et al., 2021). The radius of umbrella clouds can also represent a key ESP of large eruptions (Constantinescu et al., 2021). Typically, ESPs are derived based on the application of dedicated models either to deposit data or to real-time geophysical observations. As a result, the exposure of tephra deposits and/or the availability of accurate geophysical data are of crucial

importance to the determination of ESPs. As a matter of fact, more information is available for well-exposed deposits which are typically associated with high-intensity and large-magnitude events than for small events, which are in fact more frequent but more difficult to study (e.g., Connor et al., 2001). In addition, these small events are typically not even preserved in the stratigraphic records as they are easily eroded and mixed with soils or larger events (Gurioli et al., 2013; Kiyosugi et al., 2015). As a result, hazard assessments tend to be biased towards the largest events (e.g., Connor et al., 2001; Bonadonna, 2006). Given the low potential of preservation in the geological record, an effort is required to characterize small events during or soon after the emplacement of tephra deposits in order to have a comprehensive characterization of eruption dynamics at active volcanoes and accurate hazard assessments. Nonetheless, even in case of rapid sampling, the characterization of small explosive events is associated with a variety of challenges (e.g., small accumulations, mix with previous tephra layers, rapid erosion in case of tephra sampling, dilute plumes and clouds in

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case of remote sensing observations). In addition, a poor sampling coverage may also introduce large biases in the determination of the TEM and TGSD from tephra-fallout deposits, especially when plumes are dispersed off-shore and/or over other inaccessible areas (e.g., Andronico et al., 2014a; Bonadonna et al., 2015). In such a case, the combination of multi-disciplinary strategies might provide important insights.

Multi-disciplinary strategies have been recently developed to compute key ESPs (e.g., TEM, MER, TGSD) based on various methodologies and sensors including the combination of deposit data and geophysical observations (Scollo et al., 2009, 2015; Bonadonna et al., 2011; Corradini et al., 2016; Calvari et al., 2018; Poret et al., 2018a, 2018b; Pioli and Harris, 2019) and the combination of various sensors (Marzano et al., 2018, 2019; Mereu et al., 2020; Freret-Lorgeril et al., 2021). In particular, the combination of satellite-based observations with tephra-fallout deposit analyses has been suggested to provide additional insights into the TGSD (Bonadonna et al., 2011; Corradini et al., 2016; Poret et al., 2018a, 2018b). Indeed, satellite data have the potential to provide information on the airborne fraction of tephra, i.e., the so-called very fine ash <30 μm (Rose and Durant, 2009), that is not usually sampled on ground. In fact, even though part of the very fine ash fraction detected by satellite retrievals might fall on the ground at proximal to medial distances from the vent, part of it might remain suspended in the atmosphere for very long times (Durant, 2015).

Here we use the 29 of August 2011 weak paroxysm of Mount Etna (Italy) as a case study to illustrate various multi-disciplinary strategies for the determination of ESPs even in case of poorly exposed deposits and dilute plumes and clouds. At Mount Etna, which is one of the most active volcanoes in Europe, frequent paroxysms have occurred between 2011 and 2015 (Behncke et al., 2014; Scollo et al., 2014; De Beni et al., 2015; Calvari et al., 2018; Corradini et al., 2018; Freret-Lorgeril et al., 2018; Poret et al., 2018a; Andronico et al., 2021). Mount Etna's paroxysms are typically characterized by the emission of lava fountain-fed tephra plumes whose height can reach up to 15 km above sea level (a.s.l.) (Vulpiani et al., 2016; Calvari et al., 2018; Corradini et al., 2018). Due to the complexity of combining remote sensing and field data and to the fact that remote sensing data are not always available for the studied events, such a combination has been performed only for a few case studies (i.e., Corradini et al., 2016; Poret et al., 2018a, 2018b; Freret-Lorgeril et al., 2021).

The 29 of August 2011 paroxysm lasted 60–70 min between 03:50 and 04:53 UTC according to ground-based infrared and Doppler radar data (Calvari et al., 2018; Freret-Lorgeril et al., 2018). This eruption was associated with a transitional tephra plume that reached altitudes >9 km a.s.l. (Fig. 1A; Calvari et al., 2018; Corradini et al., 2018; Snee et al., 2021). The volcanic plume and cloud were dispersed southeastward and well detected by satellite from 04:15 to 11:00 UTC (e.g., Fig. 1B). The associated tephra-fallout deposit was sampled over 21.7 km up to the coastline (Fig. 2).

Using the 29 August 2011 paroxysm as a case study, our work mostly aims at (i) addressing the challenges associated with the determination of key ESPs (TEM and TGSD) of poorly exposed tephra deposits based on dedicated multi-disciplinary strategies and (ii) exploring the assumptions associated with the combination of observations of tephra-fallout deposits and satellite data.

2. Methodology

2.1. Deposit sampling and determination of erupted mass and plume height

A total of 13 samples were collected between 0.4 and 21.7 km from the vent (i.e., New Southeast Crater; NSEC, renamed SEC since 2020) (Fig. 2A and Table 1). It is important to note that a sampling gap of about 10 km between the proximal samples and the first medial sample exists mostly due to the presence of la Valle del Bove that does not allow access to the deposit (Fig. 2).

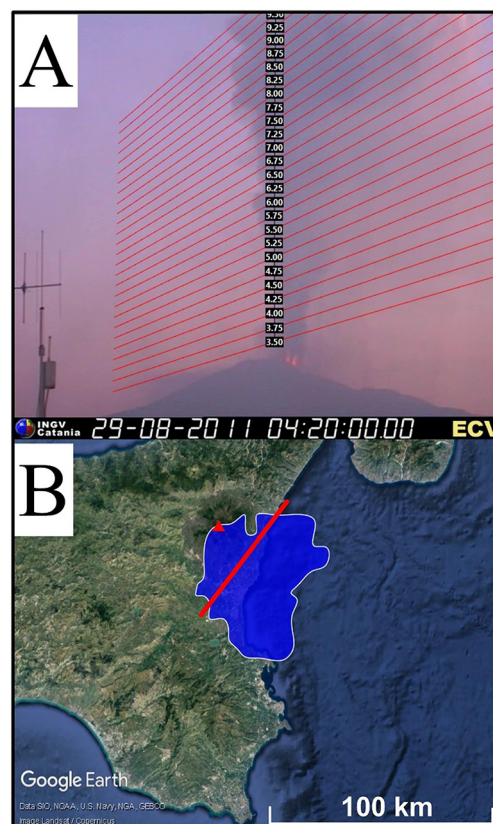


Fig. 1. A) Visual observation of the 29 August 2011 eruptive column at 04:20 UTC from Catania (INGV-OE Sezione di Catania). Isolines of plume heights above sea level are displayed in red (values are in km and reported in the middle of each isoline). B) Parallax corrected image of the tephra cloud margins (blue area) at 05:15 UTC from MSG SEVIRI. The red line indicates the transect at which the Volcanic Plume Retrieval algorithm is used to determine the mass flux of the eruptive plume/cloud (at a distance of 15 km from the vent). The red triangle indicates the location of the Southeast Crater (SEC). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

The mass of the deposit was computed based on the observed ground accumulation data as a function of the square root of the area inside the isomass contours shown in Fig. 2A using exponential fitting (Pyle, 1989), Power-law (Bonadonna and Houghton, 2005) and Weibull integration methods (Bonadonna and Costa, 2012). We also computed an isoMd_ϕ map (Fig. 2B), with Md_ϕ calculated based on Inman (1952), in order to correlate the grain-size data with the Plume height (Bonadonna and Costa, 2013a, 2013b). Finally, erupted mass and plume height were also determined based on the inversion of observed values of mass/area with the analytical model TEPHRA2 that relies on the solution of the advection-diffusion-sedimentation equation (Bonadonna et al., 2005; Connor and Connor, 2006; <https://github.com/geoscience-community-codes/tephra2>). In particular, TEPHRA2 allows to distinguish the accumulation of several particle size classes thanks to the use of a size selective diffusion law and different particle densities. The inversion employs the downhill simplex method which is a useful optimization technique that allows to find the minimum of a function of N independent variables. During inversion, we sequentially run the forward model along the downhill simplex algorithm with different sets of input parameters and end the simulation whenever the goodness-of-fit falls within the user tolerance.

2.2. Whole Deposit Grain-Size Distribution

Individual grain-size distributions, referred to as empirical GSDs hereafter (see Table 1), were obtained by mechanical sieving of all

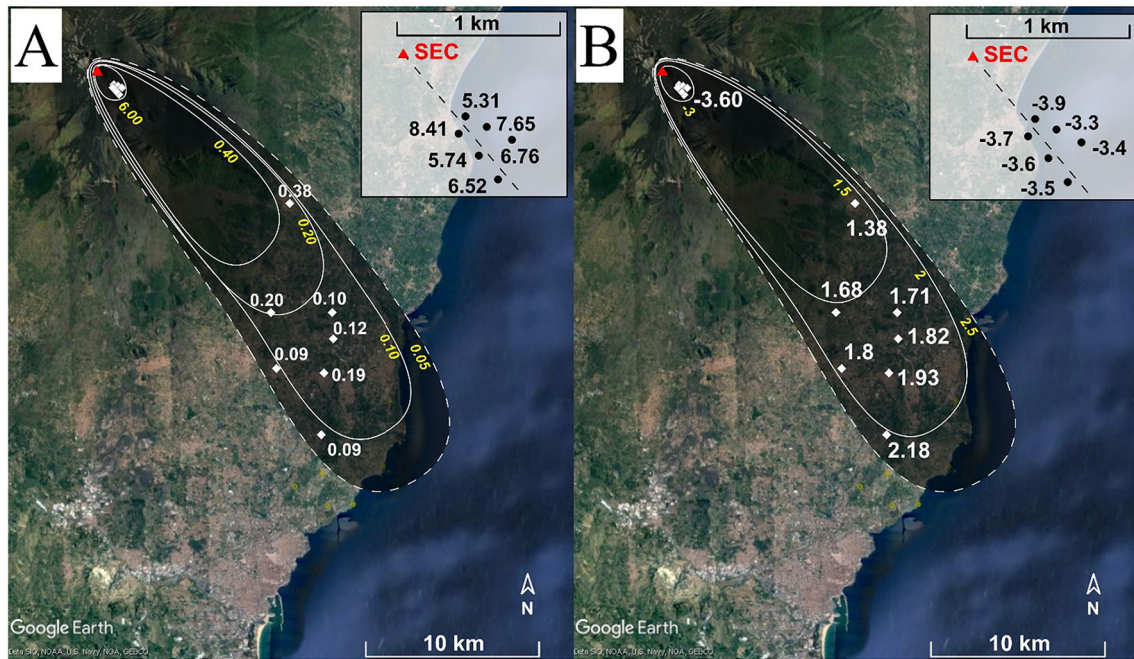


Fig. 2. Isomass (kg/m^2) (A) and isoMd_ϕ (Φ) (B) maps of the tephra-fallout deposit (see also Table 1 for more information on the samples). Isovalues are written in yellow. The red triangle indicates the SEC. Insets are a zoomed view on the proximal samples. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Table 1

Grain-Size Distributions (wt%) obtained by sieving (for individual samples) and laser diffraction (BETTERSIZER) for the cumulative residuals of the 4 Φ sieve of all samples.

		SP1	SP2	SP3	SP4	SP5	SP6	S1	S2	S3	S4	S5	S6	S7	BETTERSIZER (wt%)
Distance from the vent (km)		0.9	0.7	0.5	0.4	0.6	0.7	14.7	16.7	11.4	17.4	17.8	19	21.7	
Ground accumulation (kg/m^2)		6.52	5.74	8.41	5.31	7.65	6.76	0.09	0.20	0.38	0.10	0.12	0.19	0.09	
Φ	Size range (microns)	SIEVING (wt%)													Mean all fractions $>4 \Phi$
-6	$>64,000$	0.00	0.00	0.00	3.98	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-5.5	45,300–64,000	0.00	0.00	0.00	2.61	0.00	1.83	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-5	32,000–45,300	1.70	0.00	2.85	2.99	0.65	1.94	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-4.5	22,600–32,000	12.51	10.87	11.85	15.24	1.5	7.76	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-4	16,000–22,600	15.48	19.05	25.38	21.82	10.91	10.24	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-3.5	11,300–16,000	22.02	26.37	17.21	18.72	30.87	25.14	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-3	8000–11,300	18.63	20.43	16.11	17.58	23.20	23.20	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-2.5	5600–8000	17.76	15.95	17.12	10.20	22.23	21.51	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-2	4000–5600	7.24	5.21	5.82	3.72	6.61	5.51	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-1.5	2800–4000	3.66	1.74	2.54	1.66	3.09	2.20	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
-1	2000–2800	0.79	0.19	0.61	0.68	0.00	0.36	0.17	0.19	0.01	0.00	0.00	0.00	0.00	0.00
-0.5	1400–2000	0.19	0.00	0.35	0.47	0.29	0.00	0.00	0.18	0.10	0.16	0.00	0.00	0.06	0.00
0	1000–1400	0.02	0.00	0.17	0.3	0.00	0.32	0.04	0.18	1.53	1.35	0.13	0.02	0.02	0.00
0.5	710–1000	0.00	0.00	0.00	0.00	0.00	0.00	0.10	1.07	7.43	7.39	1.41	0.12	0.02	0.00
1	500–710	0.00	0.00	0.00	0.00	0.00	0.00	2.23	9.56	20.70	18.16	9.96	2.91	0.24	0.01
1.5	355–500	0.00	0.00	0.00	0.00	0.00	0.00	17.51	26.25	26.31	16.53	21.11	17.11	4.33	0.06
2	250–355	0.00	0.00	0.00	0.00	0.00	0.00	52.40	35.08	26.43	15.07	26.85	34.52	30.24	0.03
2.5	180–250	0.00	0.00	0.00	0.00	0.00	0.00	23.18	21.79	12.09	19.00	24.05	28.12	42.49	0.12
3	125–180	0.00	0.00	0.00	0.00	0.00	0.00	1.14	4.12	2.81	16.86	14.15	13.65	14.36	1.27
3.5	90–125	0.00	0.00	0.00	0.00	0.00	0.00	0.44	0.54	0.76	3.13	1.70	1.25	1.47	5.89
4	63–90	0.00	0.00	0.00	0.00	0.00	0.00	0.59	0.40	0.67	0.71	0.29	0.57	1.64	16.68
4.5	45–63	$>4 \Phi$ fraction													23.44
5	31.5–45														22.44
5.5	22–31.5														14.55
6	16–22														6.65
6.5	11–16														3.90
7	8–11														1.88
7.5	5.5–8														1.34
8	4–5.5	0.00	0.00	0.00	0.00	0.00	0.00	2.19	0.65	1.14	1.63	0.34	1.63	5.13	0.65
8.5	2.8–4														0.43
9	2–2.8														0.22
9.5	1.4–2														0.14
10	1–1.4														0.11
11	0.5–1														0.11

samples using $\frac{1}{2} \Phi$ intervals from -3.5Φ ($11,300 \mu\text{m}$) down to 4Φ ($63 \mu\text{m}$). The mass percentage of all sieved fractions was obtained with a 10^{-4} g resolution weighing scale. The cumulative ash fraction $>4 \Phi$ was analyzed with a BETTERSIZER laser diffraction machine at 0.5Φ intervals. In fact, the $>4 \Phi$ fraction of individual samples was too small to be analyzed individually. It is important to note that only the 7 medial samples contained particles $>4 \Phi$.

Here we distinguish between field-based Whole Deposit Grain-Size Distribution (WDGSD) and satellite-based Grain-Size Distribution (GSD_{SAT} ; see next section). It is important to note that the GSD_{SAT} combines information of both the plume and the horizontal cloud depending on plume inclination and location of plume corner, while the WDGSD only contains information of the sampled deposit. A critical combination of these two distributions is later discussed to provide the best estimate of the Total Grain-Size Distribution, i.e., the size distribution of all particles ejected during the explosive event.

One efficient way to determine the WDGSD of a non-uniform tephra-fallout deposit data set is to use the Voronoi Tessellation method (Bonadonna and Houghton, 2005; Bonadonna et al., 2015). However, the Voronoi Tessellation strategy cannot reproduce missing parts of the deposit associated with significant sampling gaps, as it is the case of the 29 August 2011 paroxysm (Fig. 2) (e.g., Alfano et al., 2016). To address this issue, we built synthetic GSDs following the strategy developed by Alfano et al. (2016). Such GSDs were obtained by fitting a Weibull function (Bonadonna and Costa, 2012, 2013a) on the variation of observed Md_{Φ} (in Φ scale; Inman, 1952) as a function of the distance from the vent. Hereafter, we refer to Md_{Φ}^* as Md_{Φ} values in metric scale. By extrapolating this fit and using the mean sorting coefficient σ_{Φ} (Inman, 1952) of all samples, we can determine synthetic GSDs at various distances from the vent that cover the sampling gap. Values of Md_{Φ} and σ_{Φ} of synthetic samples were validated based on comparison with other paroxysms at Etna and with the forward modelling of TEPHRA2 (Bonadonna et al., 2005).

Finally, we used the following Rosin-Rammler equation to fit grain-size data:

$$\text{wt}\% = e^{(-d/x_0)^l} \quad (1)$$

where wt% is the weight percentage of particle with a diameter d (μm), and x_0 and l are the length scale and the shape of the distribution, respectively; in particular, x_0 typically corresponds to the diameter at which $\sim 63\%$ of the particles are smaller (Vesilind, 1980; Pioli et al., 2019). In fact, the Rosin-Rammler fitting has been shown to best reproduce the distribution tail associated with very fine material that is often missing from the deposit (Murow et al., 1980; Pioli et al., 2019). To determine the best Rosin-Rammler parameters x_0 and l to fit our data, we used the Curve fitting application of Matlab.

2.3. Particle densities

Density of particles of different size was determined between $63 \mu\text{m}$ and 36 mm . The density of 100 scoria clasts between 16 and 36 mm selected in the proximal samples was determined using the water immersion technique of Houghton and Wilson (1989) at the Laboratory Magmas and Volcanoes (Clermont-Ferrand) (see Bonny, 2012 for more details). Densities of particles between 63 and $2400 \mu\text{m}$ selected from the most proximal of the medial samples (S3 in Table 1) were determined using water pycnometry at the University of Geneva (Eycheenne and Le Pennec, 2012; Freret-Lorgeril et al., 2019). Three pycnometry measurements were made per analyzed fraction to assess the uncertainty. Our measurements of particle density were used to model the mass decay of particles at the base of the volcanic cloud (see section 2.5).

2.4. Cloud mass and Grain-Size Distribution from satellite retrievals

We used observations made by the Spinning Enhanced Visible and Infrared Imager (SEVIRI) on board the Meteosat Second Generation (MSG) geostationary satellite to complement the determination of total erupted mass and WDGSD from tephra-fallout deposit analyses. SEVIRI is a multispectral instrument with 12 channels from visible to thermal infrared, a temporal resolution that ranges from 5 min (Rapid Scan Mode) to 15 min (Earth full disk) and a nadir spatial resolution of 3 km at sub satellite point. The ash particle retrievals have been performed using the Volcanic Plume Retrieval procedure (Pugnaghi et al., 2013, 2016; Guerrieri et al., 2015) considering the SEVIRI thermal infrared channels centered at 10.8 and $12 \mu\text{m}$ and applied to SEVIRI images from 04:15 to 11:00 UTC (resolution of 15 min). For each volcanic cloud pixel, the Volcanic Plume Retrieval procedure computes simultaneously the particles effective radius (micron), Aerosol Optical Depth (at 550 nm) and total column abundance- M_a (kg/m^2). An ash mass flux $F(t)$ (kg/s) was computed from each 15 min SEVIRI image considering only pixels at a distance of $15 \pm 1.5 \text{ km}$ from the summit craters (nearest pixels to the crater usually have large opacity that may lead to a large uncertainty in the retrievals). The flux was then computed using the following equation:

$$F(t) = l^{\text{trans}} \times v_w(t) \times \sum_{i=1}^n M_a \quad (2)$$

where l^{trans} is the transect width (in this case 3 km), v_w is the wind speed (m/s). Finally, by means of the wind speed, the time t of the flux at 15 km was then reported to the crater (0 km).

The average mass flux multiplied by the whole eruption time gives the total erupted mass (kg). Finally, the GSD_{SAT} and mass accumulation of particles belonging to the $1\text{--}20 \mu\text{m}$ fraction ($10\text{--}5.5 \Phi$) were obtained considering the whole cloud extent and combining all the images together (Wen and Rose, 1994; Prata and Grant, 2001; Stevenson et al., 2015; Corradini et al., 2016, 2018; Poret et al., 2018a, 2018b; Gouhier et al., 2019). It is important to keep in mind that particles coarser than 5.5Φ , and potentially present in the volcanic cloud, are not discriminable in the thermal infrared spectral range. The reason is that the radiative effect of these particles is approximately the same as particles of 5.5Φ . This could lead to an underestimation of the volcanic cloud total ash mass retrieved if only satellite measurements are used (see Stevenson et al., 2015, and references therein).

Following the methodology presented in Freret-Lorgeril et al. (2021), the GSD_{SAT} was computed in wt% using the ratio between the total mass of pixels containing ash particles at each time step, within a given Φ range, and the total mass of the detected volcanic plume/cloud. Finally, we averaged all individual distributions from images containing a minimum of 100 pixels with ash signal to derive the mean GSD_{SAT} of the whole event.

2.5. Mass/area decay of individual size fractions at the base of the volcanic cloud

With the aim of better understanding the relation between the tephra fraction deposited on the ground and that detected by satellite retrievals, we used the 1D integral model of Bonadonna and Phillips (2003) to estimate the mass decay of each size fraction in the umbrella cloud with distance from vent. The variation of mass per unit area (kg/m^2) for each particle size at the base of the umbrella cloud is calculated along the main dispersal axis starting from the plume corner (calculated based on the eq. 12b of Bonadonna and Phillips (2003)). Particles were considered spherical, their density was determined in the lab (section 2.3) and their terminal velocity v_t (m/s) was computed based on the drag equation of Bagheri and Bonadonna (2016a, 2016b). We used meteorological data from the hydrometeorological service of ARPA in Emilia Romagna (Scollo et al., 2009; Scollo et al., 2019) to derive wind

velocity. It is important to note that results associated with 1D modelling are associated with some uncertainties even though they provide a first approximation estimate of tephra accumulation both at the base of umbrella clouds and on the ground (Bursik et al. 1992a, 1992b; Sparks et al., 1992; Bonadonna and Phillips, 2003). In addition to the description of mass/area decay at the base of the volcanic cloud, we also inverted the tephra-fallout deposit to investigate the most characteristic position of release of individual size fractions. The height of the cloud base H_{cb} of 5.4 km (a.s.l.) was determined from the Plume Top Height H_t of 6.7 km above vent (i.e., 9.6 km a.s.l.; Corradini et al., 2018) based on the Eqs. (1) and (2) of Bonadonna and Phillips (2003).

3. Results

3.1. Total erupted mass

Fig. 3A shows the ash mass as a function of time as estimated from the different SEVIRI images. The TEM_{SAT} visible by the satellite images for the entire period (i.e., satellite-based erupted mass) has

been estimated to 2.5×10^6 kg. Interestingly, the ash mass of the imaged volcanic cloud does not vary significantly after the end of the paroxysm (04:53 UTC; vertical dashed line in Fig. 3A). The cloud is tracked up to ~400 km from the vent until 11:00 UTC.

The TEM_{DEP} (i.e., erupted mass as derived from the tephra-fallout deposit) based on the integration of the exponential, Weibull and Power-Law functions fitted on the variation of ground accumulation with the square root of the area of isomass contours (Fig. 3B; Pyle, 1989; Bonadonna and Houghton, 2005; Bonadonna and Costa, 2012) is 1.3×10^8 kg, 1.4×10^8 kg, and $1.4 \pm 0.1 \times 10^8$ kg, respectively. The average TEM_{DEP} considering all aforementioned values is $1.4 \pm 0.1 \times 10^8$ kg. The proximal integration limit for the power-law was set as established by the strategy proposed by Bonadonna and Houghton (2005). The distal limit is varied between 100 km and 300 km, i.e., the distances at which the ground accumulation becomes negligible (10^{-3} – 10^{-4} kg/m²; Fig. 3B). In any case, given that the power-law exponent is >2 , the volume estimates are not sensitive to the distal integration limit (e.g., Bonadonna et al., 2015). It is important to note that the exponential and Weibull best-fits are integrated to infinity, and, therefore,

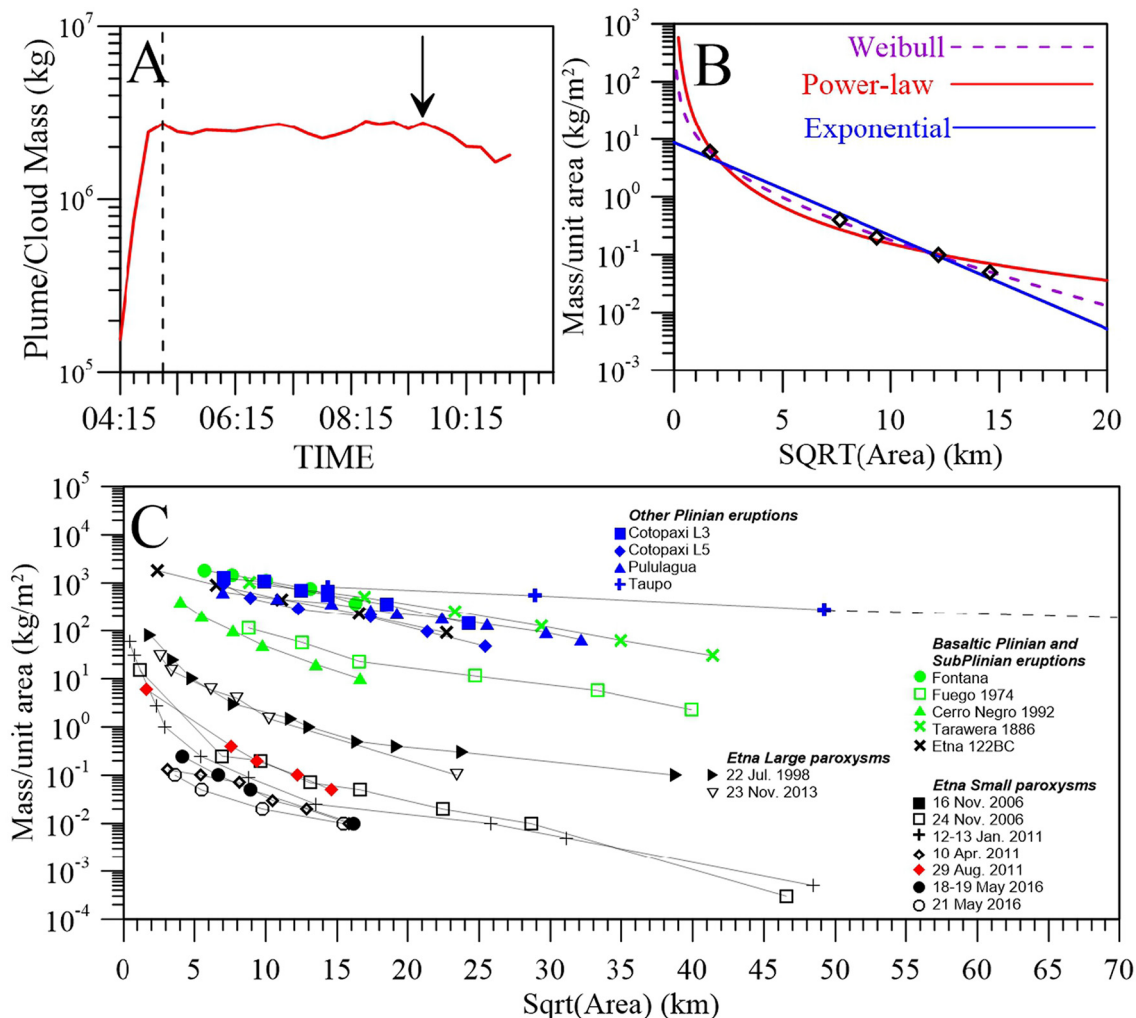


Fig. 3. A) Variation in time of plume/cloud mass (kg) as retrieved from satellite images. Vertical dashed line at 04:53 UTC represents the end of the paroxysm based on radar retrievals (Freret-Lorgeril et al., 2018). The arrow indicates the time at which the total mass starts to decrease. B) Variation of ground accumulation as a function of the square root of the isomass contour areas and associated Exponential (blue), Power-Law (red line; $y = 19.85 \times x^{-2.11}$) and Weibull (Purple dashed line) best-fits. C) Variation of ground accumulation as a function of the square root of the isomass contour areas for tephra-fallout deposits produced by small and large Etna paroxysms, Basaltic Plinian and Subplinian eruptions and other Plinian eruptions. References: Cotopaxi L3 and L5 (Biass and Bonadonna, 2011); Pululagua (Volentik et al., 2010); Taupo (Walker, 1980); Fontana Lapilli (Costantini et al., 2009); Fuego 1974 (Rose et al., 2008); Cerro Negro 1992 (Connor and Connor, 2006); Tarawera 1886 (Walker et al., 1984); Etna 122 BC (Coltelli et al., 1998); Etna paroxysm: 22 July 1998 (Bonadonna and Costa, 2013b); 16 November 2006 (Andronico et al., 2009); 24 November 2006 (Andronico et al., 2014b); 12–13 January 2011 (Andronico et al., 2014a); 10 April 2011 (Freret-Lorgeril et al., 2021); 23 November 2013 (Andronico et al., 2015; Poret et al., 2018a); 18–19 and 21 May 2016 (Edwards et al., 2018). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

they include all deposited material but not what remained suspended in the atmosphere. Moreover, the proportion of material that sedimented on land was about 7.1×10^7 kg based on the integration of the power-law fitting with a distal integration limit of 12 km, i.e., the approximate distance between the vent and the coastline along the main plume dispersal axis converted in square root of area (Figs. 2A and 3B). Using an event duration of 60–70 min (Calvari et al., 2018; Freret-Lorgeril et al., 2018), the average MER based on the TEM_{DEP} is equal to $3.5 \pm 0.3 \times 10^4$ kg/s.

Regardless of the poor deposit exposure, the thinning trend of the 29 August 2011 paroxysm (red squares in Fig. 3C) is in good agreement with the trend of other small events at Etna; associated values of MER are also in agreement (between 10^2 and 10^4 kg/s) (see also Appendix A). Note that values of MER of strong paroxysms have been estimated between 10^5 and 10^6 kg/s and such events clearly display higher ground accumulation than those obtained for small paroxysms (Fig. 3C). The inversion of tephra ground accumulation using the model TEPHRA2 (Bonadonna et al., 2005; Connor and Connor, 2006) gives an erupted mass of $1.8 \pm 0.1 \times 10^8$ kg (see Table B1 in Appendix B). Based on the inversion of the tephra deposit, we used TEPHRA2 to compute an isomass map (i.e., forward modelling; Fig. 4A) and the ground accumulation at each sampling site (Fig. 4B) which provide very similar results than the observed data (e.g., Table 1). In particular, on 13 computed datapoints, 12 of them lie within $\pm 50\%$ from the perfect match (1:1 line) with the observed samples.

3.2. Determination of plume height

Based on the ground accumulation data, we can find an inversion solution with TEPHRA2 that provides a good agreement with satellite observations of plume height, i.e., 11.1 ± 1.4 km a.s.l. against 9.0–9.6 km a.s.l., respectively. Moreover, plume height can also be derived based on Md_ϕ^* values converted in cm (Bonadonna and Costa, 2013a, 2013b). As an example, Fig. 5A shows the Weibull fit of Md_ϕ^* values with square root of area of associated contours for eruptions with style ranging between Plinian (Cotopaxi Layer 3 and 5, Pululagua 2450 BP) to sub-Plinian (Fuego 1974), with sub-Plinian grain-size data decreasing faster. Interestingly, Md_ϕ data of the 23 November 2013 strong paroxysm of Etna (derived from Poret et al., 2018a) plots in between the sub-Plinian Fuego 1974 and the Plinian style eruptions, while the 29 August

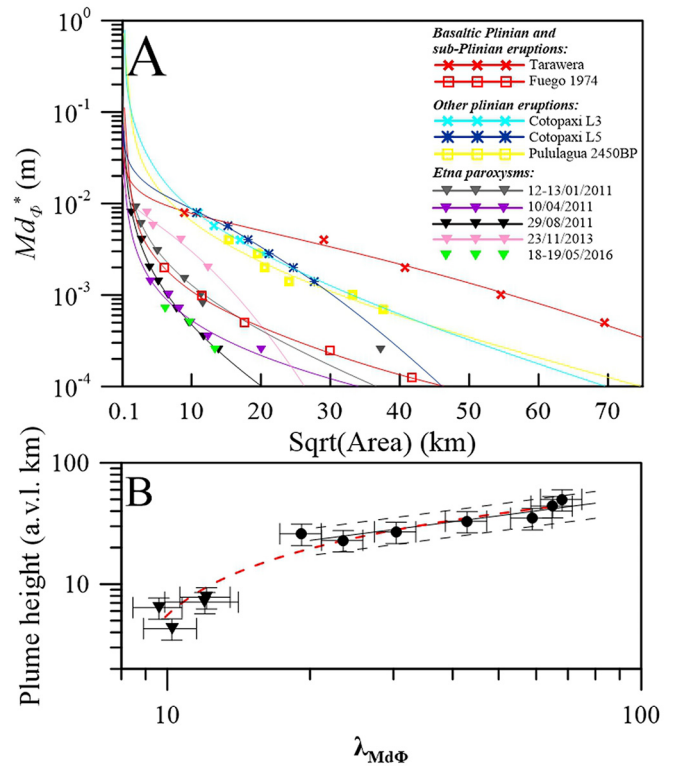


Fig. 5. A) Variation of $isoMd_\phi^*$ values (in m) as a function of the square root of the associated contour areas. In “Other Plinian eruptions” Cotopaxi L3 is andesitic, Cotopaxi L5 is basaltic-andesite and Pululagua 2450BP is dacitic; figure modified from Bonadonna and Costa (2013a, 2013b). B) Variation of plume heights above vent level (a.v.l. km) as a function of the Weibull parameter λ_{Md_ϕ} (Bonadonna and Costa, 2013a, 2013b). References are in Bonadonna and Costa (2013a, 2013b) with the exception of the 12–13 January 2011 (Andronico et al., 2014a), 10 April 2011 (Freret-Lorgeril et al., 2021), 23 November 2013 (Poret et al., 2018b) and 18–19 May 2016 (Edwards et al., 2018). See Appendix A for more details.

2011 and 10 April 2011 weak paroxysms plot below the Fuego 1974 trends. Bonadonna and Costa (2013a, 2013b) had used Md_ϕ data available in literature to determine H_t based on an empirical trend for

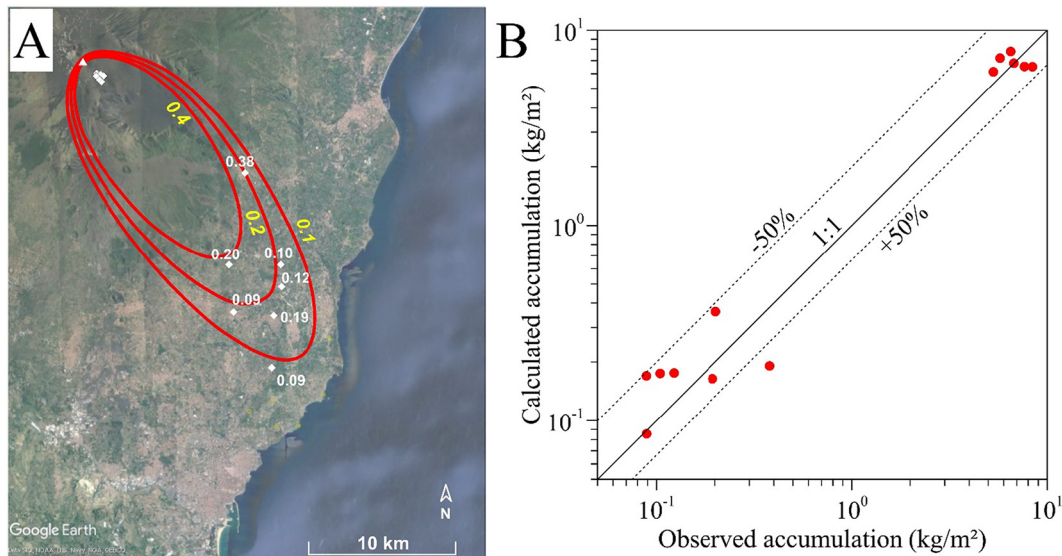


Fig. 4. A) Isomass map (kg/m^2) of the tephra-fallout deposit computed with TEPHRA2 (see also Table 1 for more information on the samples). Isovalues are written in yellow. B) Measured vs modelled ground accumulation (kg/m^2). Dashed lines indicate values of $\pm 50\%$ from the 1:1 line. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

values between 23 and 50 km. The data we obtained for the 23 November 2013, 29 August 2011, 10 April 2011 and 12–13 January 2011 paroxysms of Etna extend this trend to lower heights (7.8 km, 6.4 km, 4.3 km and 7.1 km above vent, respectively) with a new equation that presents a very good coefficient R^2 of 0.94 (Fig. 5B):

$$H_t = 20.3 \times \ln(\lambda_{Md_\phi}) - 41.1 \quad (3)$$

3.3. Whole Deposit Grain-Size Distribution

The application of the Voronoi Tessellation to the 13 collected samples (Table 1) results in a bimodal WDGSD with a gap of sizes between -2 and 1Φ centered on -1Φ (i.e., white histogram in Fig. 6A). We followed the strategy of Alfano et al. (2016) to interpolate the proximal samples with the medial to distal samples (S1 to S7) based on the variation of Md_ϕ values with distance from vent fitted by a Weibull distribution (Fig. 6B). Five equally spaced synthetic samples were created assigning the corresponding Md_ϕ^* value based on the Weibull fit (blue dots in Fig. 6B) and an average sorting σ_ϕ value (i.e., 0.7; Table 2). The resulting WDGSD obtained applying the Voronoi Tessellation on both the collected samples and synthetic data is well sorted with a Md_ϕ value of -0.9 (1866 μm) and a σ_ϕ of 2.1 (blue histogram in Fig. 6A).

Table 2

Grain-size parameters based on mechanical sieving. Md_ϕ and σ_ϕ are from Inman (1952) and x_0 indicates the Rosin-Rammler parameter below which 63% of the grain-size distribution is smaller (see Eq. (1)).

name	Md_ϕ	x_0 (mm)	σ_ϕ
SP1	−3.5	14.6	0.9
SP2	−3.6	14.9	0.8
SP3	−3.7	16.1	0.9
SP4	−3.9	19.0	0.9
SP5	−3.3	12.1	0.7
SP6	−3.4	13.4	0.8
S1	1.8	0.321	0.4
S2	1.7	0.366	0.6
S3	1.4	0.464	0.7
S4	1.7	0.407	1.0
S5	1.8	0.343	0.7
S6	1.9	0.302	0.6
S7	2.2	0.250	0.5

Proximal samples display Md_ϕ values between -3.3 and -3.9 with σ_ϕ between 0.7 and 0.9, in agreement with values observed for 8 Etna paroxysms (Table 1, Fig. 6C–D). Moreover, values of Md_ϕ and σ_ϕ are also in agreement with modelled values derived from the forward simulations of TEPHRA2 using the input parameters obtained from the inversion (Fig. B1 in Appendix B).

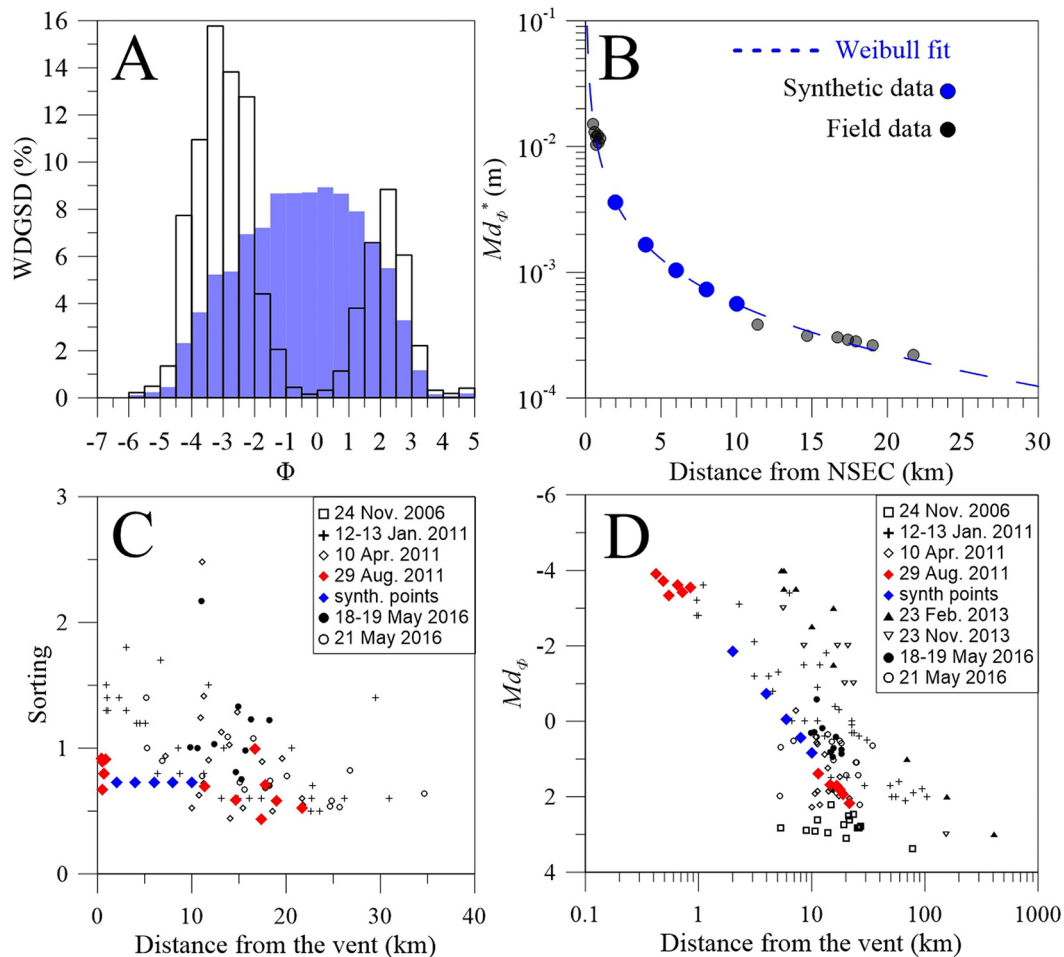


Fig. 6. A) WDGSD retrieved from the Voronoi tessellation based on the 13 collected samples (white histogram) and on the combination of the 13 collected samples with the 5 synthetic data (blue histogram). B) Variation of sample Md_ϕ^* (m) values as a function of the distance from the vent for the 29 August 2011 paroxysm. A Weibull fit is shown by the dashed blue line. The 5 synthetic data (blue circles) are plotted based on the distance from SEC vent derived in Fig. 7A. C) Sorting coefficient σ_ϕ and D) Md_ϕ values (in Φ scale) of ground samples as a function of the distance from the vent for various Etna paroxysms. References: 24 November 2006 (Andronico et al., 2014b); 12–13 January 2011 (Andronico et al., 2014a); 10 April 2011 (Freret-Lorgeril et al., 2021); 23 February 2013 (Poret et al., 2018b); 23 November 2013 (Andronico et al., 2015; Poret et al., 2018a); 18–19 and 21 May 2016 (Edwards et al., 2018). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

3.4. Particle density

Scoria clasts of size between 16 and 36 mm taken from proximal samples (Fig. 2) have densities between 279.3 and 1537.7 kg/m³ with an average value of 555.7 ± 149.4 kg/m³ (Fig. 7). The densities of ash and small lapilli clasts from 63 to 2400 μ m are between 703.1 ± 224.6 kg/m³ and 2650 ± 472.8 kg/m³, respectively. The variation of particle densities from fine to coarse-grained samples has a sigmoidal shape, similar to observations made on samples from the 2006 subplinian eruption of Tungurahua volcano (Ecuador) by Eychenne and Le Pennec (2012). Herein, we determine a relationship between the particle size (d in m) and densities ρ_p (in kg/m³) in the following form:

$$\rho_p = 2511 \times \exp(-772.8 \times d) + 451.7 \quad (4)$$

with an excellent R^2 of 0.96. This law is used hereafter to determine the densities of all particles from 10 to 6 Φ in the 1D model of Bonadonna and Phillips (2003) (see Sections 3.5 and 3.6).

3.5. WDGSD, GSD_{SAT} and TGSD

As already mentioned, the TEM_{SAT} is derived from the mass flux computed at 15 km from the vent. This value is associated with particles theoretically comprised between 1 and 20 μ m (i.e., $> 5.5 \Phi$). Our 1D simulations based on the model of Bonadonna and Phillips (2003) and using the particle density calculated in this work, show that, unless falling as part of aggregates and/or gravitational instabilities, particles $< 20 \mu$ m would sediment starting from about 350 km from vent (red arrow in Fig. 8). In addition, the main tephra fraction deposited and sampled on the ground was likely released from the base of the umbrella cloud at distances < 15 km (green arrows in Fig. 8). This shows that, for the largest part, the ash fraction detected by satellite retrievals at the plume/cloud top is not included in the analysis of WDGSD (see sampling distances in Table 1). The TGSD can, therefore, be obtained by combining WDGSD and GSD_{SAT}.

The empirical WDGSD discussed in previous section was fitted by a Rosin-Rammler distribution (red line in Fig. 9A) with best parameters $x_0 = 3474 \mu$ m and $l = 0.837$ (see Eq. (1)), a correlation coefficient R^2 of 0.99 and a Root Mean Square Error (RMSE) of 0.03. This Rosin-Rammler best fit has a Md_ϕ of -1.2 (2235 μ m) and is well sorted with a σ_ϕ of 2.0.

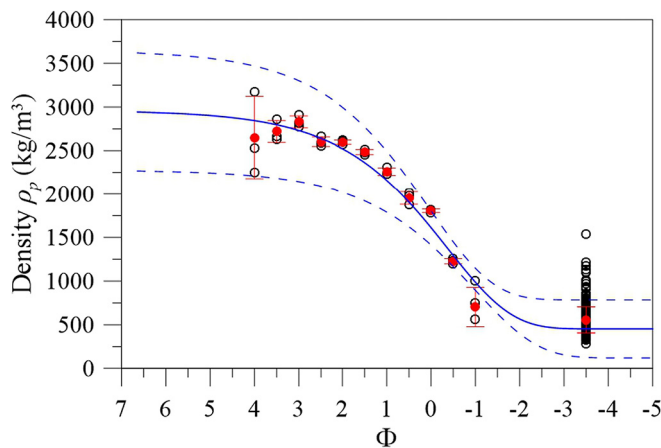


Fig. 7. Variation of clast density ρ_p (black dots) in kg/m³ as a function of particle size (Φ unit). Red dots are the average values obtained for each analyzed fraction (\pm standard deviation). The blue line corresponds to the best fit between density and particle size (see Eq. (4)) with a 95% confidence interval (dashed blue line). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

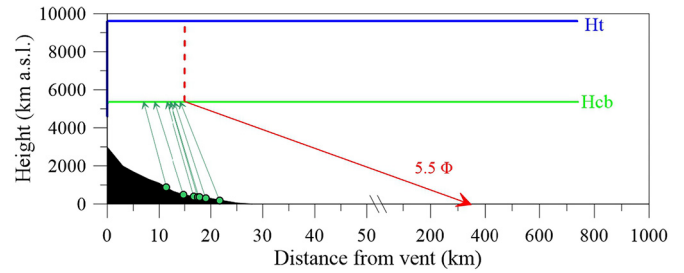


Fig. 8. Plot showing the sedimentation distance of the 5.5 Φ (31–22 μ m) particles released from the base of the umbrella cloud at a distance of 15 km from vent (i.e., distance at which the satellite ash mass flux is computed; red arrow) and release distances of the size fractions corresponding to the Md_ϕ values of ground samples (green circles). Blue and green lines correspond to the cloud top height H_t and base H_{cb} respectively. The dashed red line indicates the transect at 15 km from vent considered for satellite computation of mass flux (Fig. 1B). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Over the whole duration of the event, the average GSD_{SAT} derived from satellite that only considers particle between 10 and 5.5 Φ shows a Md_ϕ value of 7.2 Φ and a σ_ϕ of 0.8 (Fig. 9B). While mass computation methods from deposit also contain particles that sedimented beyond the coastline, the determination of the WDGSD is limited to the analyzed samples and does not extrapolate to non-sampled areas, e.g., distal zones covered by satellite images. Hence, as mentioned above, a sum of both the WDGSD and GSD_{SAT}, considering the ratio between the field and satellite total masses (i.e., 1.9%; see Section 3.1), needs to be done in order to best approximate the TGSD (black histogram in Fig. 9A).

A significant result is the similarity that exists between the wt% of the 1–20 μ m fraction (i.e., the fraction detected by satellite) observed in the TGSD and in the Rosin-Rammler fitted sieve-based WDGSD and TGSD (Fig. 9). Indeed, the 1–20 μ m fraction represents 1.9 wt% of the TGSD, 1.6 wt% of the Rosin-Rammler fitted on the TGSD (Fig. 9A). Moreover, this very fine fraction represents 1.3 wt% of the sieve-based Rosin-Rammler WDGSD. In contrast, the 1–20 μ m fraction of the sieve-based WDGSD not fitted by Rosin-Rammler is 0.03 wt%.

The combined TGSD shows a clear gap in between 3 and 6 Φ . This gap might represent a lack of fine material sampled on ground because of the limited deposit exposure, i.e., presence of the coastline (Figs. 1B and 3), and the physical limitation of the retrievals in the thermal infrared spectral range to discriminate particles coarser than 20 μ m. The Rosin-Rammler best fits suggest that this gap represents 4.6–5.2% of missing material in our dataset.

Finally, 1.5% of the total 5.5–10 Φ fraction (combined ground and satellite data) sedimented within 25 km from vent even though the characteristic sedimentation distance of this fraction is beyond 350 km (Fig. 8). Even though 1.5% is significantly lower with respect to the 46% observed for the same fraction for the Eyjafjallajökull 2010 eruption (up to 56 km from the vent; Bonadonna et al., 2011), we conclude that this fraction has been probably sedimented as part of aggregates and/or gravitational instabilities. Aggregates were not observed in the deposit, but they could have been broken with impact on the ground.

3.6. Mass/area decay at the base of the volcanic cloud

The mass/area decay of various particle sizes at the base of the volcanic cloud is determined along the dispersal axis based on both satellite and 1D model estimates (Fig. 10A). The mass accumulation (kg/m²) from satellite was obtained considering the maximum value of each transect perpendicular to the dispersal cloud axis and by considering all the SEVIRI images from the start of the eruption until 11:00 UTC. The modelled decay is computed with the model of Bonadonna and

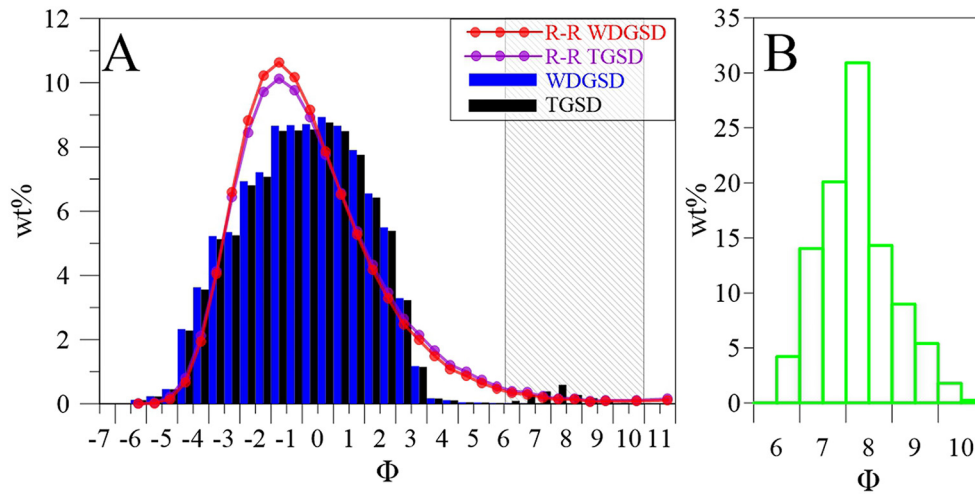


Fig. 9. A) Empirical WDGSD (blue histograms, equivalent to the blue histograms in Fig. 7A) and associated Rosin-Rammler best fit (red line). The empirical WDGSD was combined with GSD_{SAT} from satellite retrievals to obtain the best approximation of TGSD (black histogram); the TGSD Rosin-Rammler best fit is also shown with the purple line. The gray area corresponds to the 10–5.5 Φ area detected by satellite retrievals. B) Satellite-based (computed) average GSD_{SAT} used to derive the TGSD in Fig. 9A (black histograms). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

Phillips (2003) based on the assumption that the TGSD as calculated in previous sections is associated with the average TEM as derived from tephra-fallout deposit analysis (1.4×10^8 kg) combined with the satellite-derived mass (2.5×10^6 kg). It is interesting to notice that the modelled accumulation of the 5.5–10 Φ fraction is twice to three times higher than the same fraction as retrieved from satellite data but presents a similar decay. The overestimation of the modelled fraction could be related to the fact that the 5.5–10 Φ fraction sedimented closer to the vent than expected due to aggregation and/or gravitational instabilities. Such processes are not taken into account in our 1D model. In addition, our 1D model also shows that at 15 km from vent the 3–6 Φ fraction is associated with significant values of mass/area ($>10^{-1}$ kg/m²) even though it is not detected by the satellite retrieval algorithm (Fig. 10B). This might explain the gap between 3 and 6 Φ shown in Fig. 9A. Finally, it is also interesting to notice that around 450 km from vent, when the cloud is very diluted, all the size fractions detectable by the satellite retrievals (5.5–10 Φ) are below 10^{-2} kg/m² (Fig. 10C).

4. Discussion

4.1. Plume height determination

Even though the distribution of the largest clasts around a volcano is typically used to determine the maximum plume height (e.g., Carey & Sparks, 1986; Rossi et al., 2019), such strategy is scarcely used at Etna due to the typical lack of proximal data and the small size of distal tephra clasts (e.g., Coltelli et al., 1998). Here we have shown how the variation of Md_ϕ as a function of the square root of iso Md_ϕ areas (Figs. 2B and 5) can also be used at Etna to retrieve an average plume height of the event based on the value of λ_{Md_ϕ} (using a Weibull function; Bonadonna and Costa, 2012, 2013a, 2013b). In particular, with this study we have extended the empirical relation proposed by Bonadonna and Costa (2013a, 2013b) to average H_t values from 23 km down to about 4 km above vent (i.e., height of 10 April 2011 paroxysm; Freret-Lorgeril et al., 2021). We expect such a trend to help characterize H_t from tephra-fallout deposits of future eruptions, especially in case of missing remote sensing data. We also found that, regardless of the poor exposure of the deposit, the inversion with TEPHRA2 of ground accumulation data provides consistent results for plume height with satellite observations (i.e., between 9 and 11 km a.s.l.; see Appendix B). The height differences between TEPHRA2 and satellite estimates shown in Appendix B are not significant in comparison to the typical uncertainties

observed by satellite and visible imagery, i.e., of about ± 0.5 km (Corradini et al., 2018; Scollo et al., 2019 and references therein). This agreement also confirms that the mass/area values used for TEPHRA2 inversion are related to the spreading of the cloud that was observed in the atmosphere for about 6–7 h (based on satellite images) with a duration of the associated paroxysm of a little more than one hour. In fact, most samples were collected beyond the plume corner (around 1–2 km), being the height above the vent around 6–8 km and the plume being strong (Calvari et al., 2018).

4.2. Insights into TGSDs from the Rosin-Rammler fitting

Based on a large data set, Pioli et al. (2019) have found the Rosin-Rammler formulation (Eq. (1)) as the best strategy to reconstruct the grain-size distribution tails of tephra-fallout deposits. This is particularly interesting given that WDGSDs often lack proximal and/or distal data due to erosion and/or difficult accessibility. Accordingly, using the best x_0 and l parameters (e.g., with Matlab), we were able to reproduce the 29 August 2011 empirical WDGSD with an excellent $R^2 > 0.99$ (with Md_ϕ value of -1.0 ; Fig. 6A). We found that the wt% of the 5.5–10 Φ very fine ash fraction well correlates with the WDGSD and the TGSD (Fig. 9A). This suggests the capacity of the Rosin-Rammler equation to reproduce the tails of WDGSDs (Pioli et al., 2019) that are rarely sampled on ground and to provide a first order estimate of the airborne fraction potentially detected by satellite.

Both Rosin-Rammler parameters, x_0 and l , associated with the empirical WDGSD and the TGSD (resulting from the combination of the WDGSD and the GSD_{SAT}) are in the same range with respect to those found for 49 tephra deposits at other volcanoes analyzed by Pioli et al. (2019) (Fig. 11A). They also show similar trends of x_0 versus Md_ϕ^* (m) and l versus σ_ϕ (Fig. 11B,C). In particular, whereas x_0 intuitively increases as a function of Md_ϕ^* (m) (Fig. 11B), the better sorted the WDGSDs, the smaller the parameter l (Fig. 11C). In particular, the Rosin-Rammler WDGSD of the 29 August 2011 event presents close l value (0.84) with respect to the 12–13 January SEC paroxysm ($l = 0.94$) and the 27 October 2002 South-East Crater strong eruption ($l = 0.98$), for which proximal data (i.e., <5 km from the vent) were also acquired (Andronico et al., 2008, 2014a). Even if other data obtained at Etna present highly variable values of l (blue squares in Fig. 11A-C), it is important to bear in mind that extreme l values are not necessarily associated with poor deposit exposure. As an example,

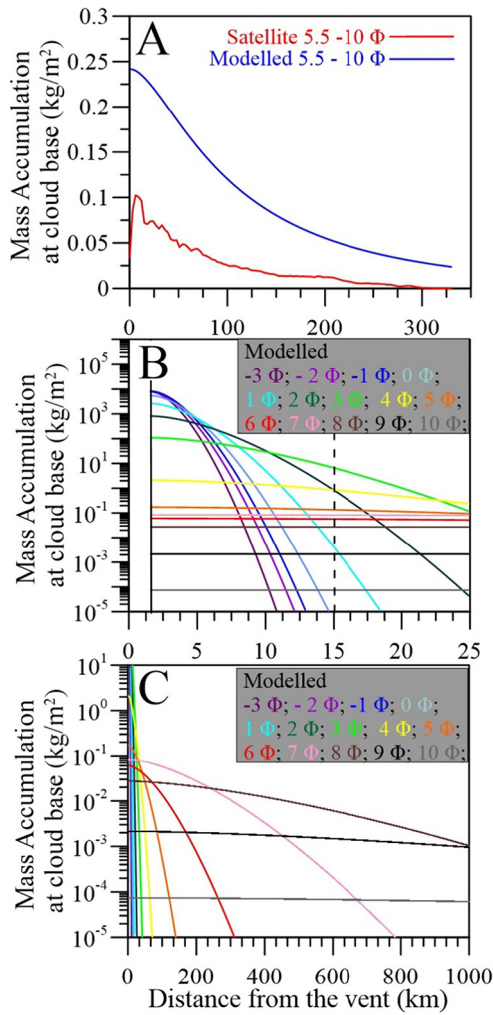


Fig. 10. A) Mass accumulation at the base of the umbrella cloud of all ash fractions as retrieved from satellite images (red line) and based on the 1D sedimentation model of Bonadonna and Phillips (2003) for the 5.5–10 Φ fraction (blue line). Modelled mass accumulation at the cloud base of particle sizes between -3 and 10Φ up to 25 (B) and 1000 km (C) from the vent. The vertical dashed line in B indicates the distance of 15 km from the vent at which satellite mass estimates are computed. The black vertical line in B indicates the plume corner location ($=0.2H$; Bonadonna and Phillips, 2003). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

the highest l of 2.47 (Fig. 11A–C) was derived from the 24 November 2006 deposit that was sampled from proximal (1 km) to distal (80 km) areas at 27 sampling sites (Andronico et al., 2014b).

Various attempts to correlate grain-size parameters (i.e., fine ash content, Md_{ϕ} , σ_{ϕ}) with eruptive style and ESPs exist (e.g., Rust and Cashman, 2011; Costa et al., 2016; Pioli et al., 2019). The percentage of tephra <1 mm, measured along the dispersal axis where the isopach is 10% of the maximum deposit thickness, was also proposed by Walker (1973) as an indicator of explosiveness of an eruption (i.e., the fragmentation index) used to classify explosive eruptions together with the area of pyroclastic dispersal. Nonetheless, all these correlations and classifications of eruptive styles based on grain-size parameters are affected by the large uncertainties associated with both the determination of GSDs and WDGSs and with the determination of ESPs (e.g., MER, plume height); therefore, they often only provide limited and misleading insights. Similarly, the overall 1–20 μm fractions as derived from the Rosin-Rammler best fits of the aforementioned WDGSs/TGSs present a poor correlation with MER (Fig. 12A; Pioli et al., 2019 and references therein). Nonetheless, a better correlation is shown if data are grouped in 3 categories: 1) tephra-fallout deposits older than 1000 years (i.e., most likely depleted in very fine ash due to erosion and/or poor deposit exposure; open squares), 2) recent events rich in very fine ash due to specific eruptive styles and dynamics (e.g., events associated with PDCs), collective sedimentation processes (e.g., aggregation and/or gravitational instabilities) and/or efficient sampling strategy (i.e., sampling carried out during or just after deposit emplacement) and/or good deposit exposure (i.e., large part of the deposit land and accessible for sampling) (green triangles) and 3) all other events of Pioli et al. (2019) dataset including Etna's eruptions (black and light blue circles, respectively) (Fig. 12A). The outlier with very small content in very fine ash (black dot) corresponds to the 1986 coarse-grained basaltic-andesite eruption of Izu-Oshima volcano that forms a small island in Sagami Bay (Japan), and, for which, only the proximal deposit could be analyzed (Mannen, 2006). Moreover, the content of very fine ash increases as a function of satellite-based MERs (Fig. 12B; data from Gouhier et al., 2019). These results suggest that, in case of good deposit exposure and sampling, the content of very fine ash should increase as a function of eruption intensities, with some variation due to specific eruptive dynamics (e.g., presence of co-PDC plumes). It is important to also note that some of the data scatter in Fig. 12A is partly due to the various strategies used in literature to determine MER either from plume height (e.g., Sparks, 1986; Wilson and Walker, 1987; Mastin et al., 2009; Degruyter and Bonadonna, 2012) or from the ratio between deposit mass and eruption duration.

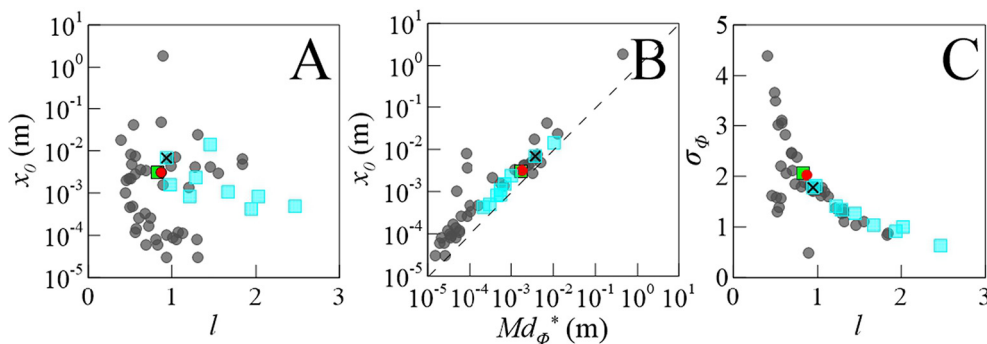


Fig. 11. Variation of Rosin-Rammler parameter x_0 A) as a function of l and B) as a function of Md_{ϕ}^* (m) for all WDGS/TGSs available in literature (see references in Pioli et al., 2019). (C) The Rosin-Rammler parameter l is also plotted as a function of Rosin-Rammler WDGS sorting coefficients. The red dot and green square indicate values for the Rosin-Rammler WDGS and the Rosin-Rammler TGS of the 29 August 2011 paroxysm. Gray dots correspond to the large data set of WDGS obtained from 49 tephra-fallout deposits that also includes 8 Etna deposits (blue squares; Pioli et al., 2019 and references therein). The crossed blue square indicates the 12–13 January 2011 paroxysm, which represents the best studied tephra-fallout deposit at Etna. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

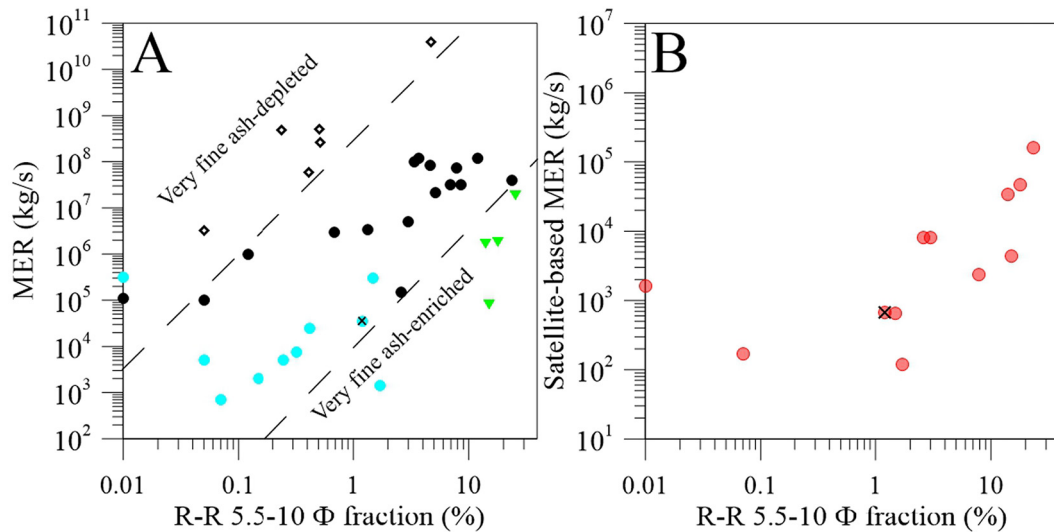


Fig. 12. A) Variation of the 5.5–10 Φ fractions (%) as derived from the Rosin-Rammler fitting of all WDGSs/TGSs available in literature as a function of corresponding eruption MERs (see Pioli et al., 2019 and references therein). Open squares indicate all events that occurred more than 1000 years ago (depleted in very fine ash; i.e., 4 ka Rungwe Pumice eruption, Cotopaxi layer 3 and 5, Pululagua 2450 BP, Tecolote and Baia - Fondi di Baia eruption at Campi Flegrei). Green triangles indicate the events rich in very fine ash (i.e., Mount Saint Helens 1980, Mt. Spurr 1992 and the Eyjafjallajökull 2010). Black and light blue dots correspond to all other events and Etna's eruptions, respectively as reported in the dataset of Pioli et al. (2019). Dashed lines show indicative fields of deposits depleted or enriched in very fine ash fractions. B) MERs computed from satellite method (references in Gouhier et al., 2019) as a function of the 5.5–10 Φ fractions (%) as derived from the Rosin-Rammler fitting of corresponding WDGSs/TGSs (Pioli et al., 2019 and references therein). Black crosses in both figures indicate the 29 August 2011 Etna paroxysm. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

4.3. Multidisciplinary determination of ESPs

The application of the three empirical fitting methods for the integration of EM from tephra-fallout deposits (i.e., Exponential, Power-Law and Weibull) provides similar results due to the relatively good exposure of the deposit in relation to a relatively small paroxysm. In fact, the exponent of the Power-Law fit is >2 (Fig. 3B), suggesting both a small magnitude and limited dispersal of the associated eruptive event and a low impact of distal extreme of integration on the erupted mass estimates (Bonadonna et al., 2015). As a result, we consider the average deposit TEM of $1.4 \pm 0.2 \times 10^8$ kg derived from the Exponential, the Power-Law (with integration limits of 100 and 300 km) and Weibull empirical methods (Fig. 3B) as the best estimate. It should be noted that all these values of TEM are associated with various levels of uncertainty which originate from sampling of mass/area (up to 30%), data contouring, availability of samples and the deposit exposure (up to 70%; e.g., Bonadonna et al., 2015). Nonetheless, this estimate is still in agreement with results from the inversion of ground accumulation data with TEPHRA2 (i.e., $1.8 \pm 0.1 \times 10^8$ kg, see Appendix B) and compares favorably to 4.4×10^8 kg obtained from L-band radar retrievals (Freret-Lorgeril et al., 2018). It is important to bear in mind that, the Exponential and Weibull fits are integrated to infinity, whereas the Power-Law is ideally integrated to the most distal extension of the tephra-fallout deposit. Hence, as stated earlier, our TEM_{DEP} accounts both for the material that sedimented on land (i.e., ~50% of the TEM_{DEP}) and for the material that sedimented beyond the coastline (i.e., >25 km), including the material detected by satellite that eventually has fallen either on the ground or in the sea (Fig. 1B).

TEM_{SAT} and GSD_{SAT} are likely to represent both airborne and deposited very fine ash (either on land or out at sea), even though it is difficult for the two fractions to be quantified. The mass accumulation of the 5.5–10 Φ fraction detected by SEVIRI is shown to decrease as a function of the distance from vent (Fig. 10A). Our estimates show that this very fine ash, if falling as individual particles, would eventually reach the ground at distances beyond 350 km from the vent (Fig. 8) and is, in any case, considered in the deposit-based mass estimates (i.e., TEM_{DEP}) (Fig. 3B). The ratio between TEM_{SAT} and TEM_{DEP} is 1.8%,

1.9% and 2.0% depending on the integration based on the Weibull, Power-law, and Exponential fit, respectively, with a mean of $1.9 \pm 0.1\%$. This means that at 15 km, $1.9 \pm 0.1\%$ of the 5.5–10 Φ fraction is still in the volcanic cloud. Interestingly, this ratio with the TEM_{DEP} computed with TEPHRA2 is similar, with a value of 1.4%. Nonetheless, we cannot exclude that some of the ash retrieved in the satellite images might never sediment. In such a case, instead of using the ratio between both TEM_{SAT} and TEM_{DEP} to determine the proportion of very fine ash that is detected by satellite (see section 3.2), TEM_{SAT} should be divided by TEM_{SAT} + TEM_{DEP}. In the case of the 29 August 2011 paroxysm, the two strategies result in a similar proportion of 1.8–1.9%, respectively. In addition, it is important to note that particles coarser than 5 Φ are likely to remain in the volcanic cloud beyond 15 km as shown by our 1D model (accumulation $>10^{-1}$ kg/m² at the base of the volcanic cloud up to at least 25 km from vent; Fig. 10B and C). If these coarse-ash particles contribute to the cloud thermal signature, they could not be discriminated from 20 μ m particles due to Mie Scattering effects (Prata and Grant, 2001; Stevenson et al., 2015; Gouhier et al., 2019). In such a case, TEM_{SAT} should be considered as minimal values.

Differently with respect to the TEM estimates, the very fine ash fraction detected by satellite retrievals needs to be added to the WDGSd derived from tephra-fallout deposits. In fact, all current strategies used to determine the WDGSd (including the Voronoi Tessellation) only consider the sampled tephra-fallout deposit and do not extrapolate grain-size data to infinity or to the distal extent of the deposit. Our 1D model simulations show that the >5.5 Φ fraction would deposit in the sea (Fig. 8), and, therefore, if it did not fall closer to the vent as aggregates and/or gravitational instabilities, it is not accounted for in the WDGSd. Hence, WDGSd and GSD_{SAT} can be combined based on the relative mass proportion.

It is important to note that our combined satellite-field TEM estimate characterizes the fraction of tephra that is transported in the atmosphere by the buoyant plume and umbrella cloud and deposited away from the eruptive vent (i.e., the SEC). Nonetheless, in the case of lava fountain-fed tephra plume events such as Mt. Etna's paroxysms, lava fountains generate an additional very proximal deposit that builds the

cone (Behncke et al., 2014; Andronico et al., 2014a; De Beni et al., 2015; Spanu et al., 2016; Freret-Lorgeril et al., 2018). Such a very proximal tephra-fallout deposit associated with the lava fountain should be further investigated to better understand the dynamics of lava fountain-fed tephra plumes generated during paroxysmal events at Etna and other volcanoes (e.g., Snee et al., 2021).

Finally, based on the comparison with observations from other Etna paroxysms and modelled Md_{ϕ} values (Figs. 3, 6 and Appendix B), we have shown that using synthetic data to fill sampling gaps in order to determine the WDGSD not only is applicable to large deposits such as that of the 2008–2013 Chait n eruption (Alfano et al., 2016) but also to small deposits such as that of the 29 August 2011 paroxysm of Etna.

5. Conclusions

Quantifying ESPs at very active volcanoes such as Mount Etna is crucial for both real-time ash dispersal forecasting and long-term hazard assessment. The determination of ESPs of weak explosive events associated with poorly exposed tephra-fallout deposits has been proven to be especially challenging. Nonetheless, the characterization of weak explosive events is crucial to an accurate assessment of the whole range of eruption intensity associated with a given volcano. In this work, we combined particle size analyses performed on a sparsely sampled tephra-fallout deposit with numerical modelling and satellite-based observations of a weak paroxysm of Mount Etna (i.e., the 29 of August 2011 event) to investigate the potential of multidisciplinary strategies for the determination of ESPs.

In terms of characterization of small, poorly exposed tephra-fallout deposits we have shown how:

- 1) all empirical integrations used and the inversion of ground accumulation data with TEPHRA2 provide similar estimates of TEM_{DEP} ($1.4 \pm 0.0 \times 10^8$ kg vs $1.8 \pm 0.1 \times 10^8$), which could be due to the limited size of the tephra-fallout deposit;
- 2) plume height of small paroxysms that are not associated with large clasts (especially large lithic clasts) can also be derived from tephra-fallout deposits based on inversion of ground accumulation data as well as on the variation of Md_{ϕ} values with distance from vent;
- 3) poorly-exposed tephra-fallout deposits associated with large sampling gaps related to critical sedimentation regimes should be carefully treated for the determination of WDGSD. Here we show that synthetic data, validated by both the comparison with observed data from other Etna paroxysms and simulations with TEPHRA2, can be used to reconstruct WDGSD that would otherwise result in a bimodal distribution. The resulting WDGSD obtained applying the Voronoi Tessellation on both the collected samples and the synthetic data is well sorted with a Md_{ϕ} value of -0.9 and a σ_{ϕ} of 2.1 . The Rosin-Rammler fitting of such a WDGSD displays similar distribution parameters with respect to other well-sampled WDGSDs at Etna;

In terms of multi-disciplinary determination of ESPs, we have shown how:

- 4) the variation of particle density with particle size shows a sigmoidal distribution that can be used to model particle density in analytical and numerical models;
- 5) the ash fraction detected by satellite retrievals is not included in the analysis of WDGSD. In fact, unless falling as part of aggregates and/or gravitational instabilities, particles $<20 \mu\text{m}$ would sediment starting from about 350 km from vent, while the main tephra fraction deposited and sampled on the ground was released from the base of the umbrella cloud at distances <15 km where satellite data are retrieved for grain-size. As a result,

in order to obtain a TGSD, WDGSD and GSD_{SAT} should be combined considering their relative mass proportion;

- 6) the Rosin-Rammler fitting already shown to best reproduce the tails of WDGSDs (Pioli et al., 2019), is here shown to also best reproduce the tail of the TGSD. In fact, the very fine ash fraction detected by satellite retrievals ($1\text{--}20 \mu\text{m}$) observed in the TGSD ($1.9\text{--}1.2$ wt%) is similar to that observed in the Rosin-Rammler fitted sieve-based WDGSD. In contrast, the $1\text{--}20 \mu\text{m}$ fraction of the sieve-based WDGSD not fitted by Rosin-Rammler is 0.03 wt%;
- 7) the Rosin-Rammler fitting was also used to reconstruct the gap of $3\text{--}6 \Phi$ observed in the TGSD suggesting that this fraction is associated with $3.8\text{--}5.3$ wt% even though it is not represented either in the GSD_{SAT} or in the WDGSD. In fact, our 1D modelling shows that at 15 km where satellite data for mass flux are retrieved, the $3\text{--}6 \Phi$ fraction is associated with significant values of mass/area ($>10^{-1}$ kg/m²) even though it is not detected by the satellite retrieval algorithm; Additionally, such a fraction is not represented in the WDGSD as it is lost at sea;
- 8) the ratio between TEM_{SAT} and TEM_{DEP} is especially important for the determination of TGSD; in case of the 29 August 2011 paroxysm, TEM_{SAT} represents $1.9\% \pm 0.1\%$ of the total deposited mass (i.e., TEM_{DEP}) and 1.4% of the computed TEM_{DEP} with TEPHRA2. It is important to note that TEM_{DEP} accounts both for the material that sedimented on land (i.e., $\sim 50\%$ of the TEM_{DEP} in case of the 29 August 2011 event) and for the material that sedimented beyond the coastline (i.e., >25 km), including the material detected by satellite that eventually has fallen either on the ground or in the sea. So, differently with respect to the determination of TGSD (point 4), the TEM_{SAT} should be considered as mostly contained in TEM_{DEP} even though we cannot exclude that some of the airborne material never sediments and, therefore, is not included in TEM_{DEP} ;
- 9) 1.5 wt% of the $5.5\text{--}10 \Phi$ fraction of the whole TGSD fell on land (within 25 km from vent) and was observed in the deposit even though the characteristic sedimentation distance of this fraction is beyond 350 km; interestingly, our modelled accumulation of the $5.5\text{--}10 \Phi$ fraction is twice to three times higher than the same fraction as retrieved from satellite data but presents a similar decay supporting the possibility that this small fraction was sedimented as part of aggregates and/or gravitational instabilities;
- 10) at 400 km from vent, when the cloud is very diluted, the modelled mass/area at the base of the volcanic cloud of all the size fractions detectable by the satellite retrievals ($5.5\text{--}10 \Phi$) are below 10^{-2} kg/m² indicating that most of the mass sedimented at distances <400 km as suggested by SEVIRI.
- 11) the content in very fine ash as derived from the Rosin-Rammler fits of various WDGSDs/TGSDs tends to increase as a function of both plume-height-derived MER and satellite-derived MER. This suggests an important correlation between fine-ash content and ESPs often attempted but complicated by the large associated uncertainties and that should be further investigated based on more accurate estimates of WDGSDs/TGSDs (i.e., associated with recent and rapidly sampled tephra-fallout deposits and multi-sensor strategies such as the one presented in this paper).

Author contribution statement

FD, LuG and SS were in charge of collecting the samples and of their preliminary analyses. SC and LoG have provided the results of the SEVIRI retrievals. JL has performed the inversion and forward modelling with

TEPHRA2. ER has performed the 1D model simulations. VFL has carried out all WDGSD, TGSD, TEM and particles density estimates. VFL and CB have provided a first interpretation of the results and written the first draft of the paper. All authors have contributed to the editing and finalization of the paper.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Summary of all estimates derived from tephra-fallout deposits of all presented eruptions in this study. See the main text for references. H_t : Plume top height (km above sea level); TEM: Total Erupted Mass (kg); MER: Mass Eruption Rate (kg/s); ΔMd_ϕ , $\Delta\sigma_\phi$ and $\Delta kg/m^2$ are the variation of Md_ϕ , sorting coefficient σ_ϕ (Inman, 1952) and ground accumulation of observed samples, respectively; $\lambda_{Md\phi}$ is the Weibull parameter described by the variation of tephra-fallout deposit Md_ϕ as a function of the square root of the contour areas observed on each deposit (see Fig. 5 and main text for more details); WDGSD: Whole Deposit Grain-Size Distribution

	H_t (km)	TEM (kg)	Volume (m ³)	Duration (min)	MER (kg/s)	ΔMd_ϕ	$\Delta\sigma_\phi$	$\Delta kg/m^2$	$\lambda_{Md\phi}$	WDGSD Md_ϕ	WDGSD σ_ϕ
Cotopaxi L3	23.0	3.8×10^8	0.541	570 ± 195	1.1×10^7	/	/	140–1260	23.5 ± 2.4	/	/
Cotopaxi L5	26.0	2.2×10^{11}	0.23	134 ± 25	2.7×10^7	/	/	48–950	19.2 ± 1.9	/	/
Pululagua	27.0	$4.5 \pm 0.3 \times 10^{11}$	/	37 ± 13	6.2 $\pm 3.8 \times 10^7$	/	/	64–644	30.4 ± 3.0	/	/
Taupo	50.0	6.6×10^{12}	12	360–1020	$1-3 \times 10^8$	/	/	65–826	68.1 ± 6.8	/	/
Fontana (D–F)	27.0–28.0	$2.1-2.8 \times 10^{12}$	2.9–3.8	240–360	1.4×10^8	–3.3–0.9	1.0–2.7	368–1838	/	/	/
Fuego 1974	~14.0	5.2×10^{10}	/	/	/	–2.0–3.0	0.3–1.0	2–114	/	0.05	1.4
Cerro Negro 1992	7.0–7.5	2.35×10^{10}	/	/	/	/	/	10–400	/	–0.1	1.5
Tarawera 1886	33.0	/	0.472	180	1.8×10^8	/	/	30–1000	42.9 ± 4.3	/	/
Etna 122BC	/	3.6×10^{11}	0.4	/	$5-8.5 \times 10^8$	/	/	6–517	/	/	/
16/11/2006	~4.5	7×10^6	/	756	1.5×10^2	/	/	$(0.12-59.5) \times 10^3$	/	~2	~0.8
24/11/2006	~5.0	1.9×10^8	/	/	5×10^3	0.7–3.4	0.4–0.6	$(0.33-1068.5) \times 10^3$	/	~1	/
12–13/01/2011	6.8–13.8	$1.5 \pm 0.4 \times 10^8$	/	100	2.5 $\pm 0.7 \times 10^4$	–3.6–2.1	0.5–1.8	$(0.0-63,542.4) \times 10^3$	$12.0-2.1$	–1.4	/
10/04/2011	6.1–9	$2.5 \pm 1.9 \times 10^7$	/	310 ± 94	3.5 $\pm 0.3 \times 10^4$	–0.3–2.3	0.4–2.5	$(0.11-133.5) \times 10^3$	10.2 ± 1.3	–1	0.9
29/08/2011	9–9.6	1.35×10^8	/	60–70	3.5 $\pm 0.3 \times 10^4$	–3.9–2.2	0.4–1.0	$(0.09-8412) \times 10^3$	9.6 ± 1.1	–0.9	2.1
18–19/05/2016	6.0–6.5	$1.1-6.0 \times 10^7$	/	120	$0.6-3.3 \times 10^3$	–0.6–1.8	0.7–2.2	$(10.7-406.1) \times 10^3$	/	0.35	1.02
21/05/2016	6.0	$4.6-6.5 \times 10^6$	/	130	$5.9-8.3 \times 10^2$	0.3–2.2	0.5–1.4	$(2.8-121.3) \times 10^3$	/	0.64	0.89
22/07/1998	11.0–13.0	$1.7-1.8 \times 10^9$	/	20–40	$0.6-1.2 \times 10^6$	/	/	0.1–80	/	0.8	1.8
23/02/2013	/	$1.3 \pm 0.1 \times 10^9$	/	69 ± 35	1.2×10^6	–3.5–3	/	$(16-17,000) \times 10^3$	/	/	/
23/11/2013	~11.0	$1.3 \pm 0.1 \times 10^9$	/	60	3.1 $\pm 0.5 \times 10^5$	–3–3	/	$(1.4-21,000) \times 10^3$	12.1 ± 1.5	–3.4	1.26

Appendix B. TEPHRA2 estimates. The model is accessible at <https://github.com/geoscience-community-codes/tephra2>

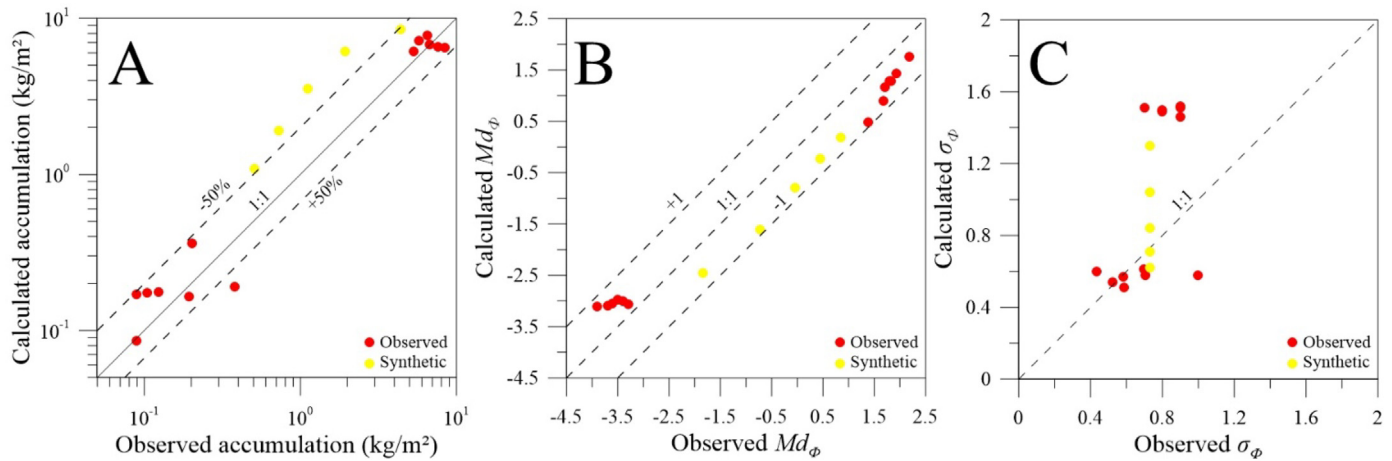


Fig. B1. Observed vs calculated estimates with TEPHRA2 of A) ground accumulation (kg/m^2), B) Md_ϕ and C) σ_ϕ at each sampling sites for the 29 August 2011 tephra-fallout deposit. Dashed lines in A correspond to $\pm 50\%$ from the 1:1 line. Dashed lines in B correspond to $\pm 1 \Phi$ from the 1:1 line. Yellow points correspond to synthetic data (see Fig. 6B of main text). Red points indicate collected samples.

Table B1

Comparison of estimates based on Tephra2 inversion and other strategies: H_t (satellite), TEM_{DEP} (from tephra-fallout deposit based on empirical integration), WDGS (from tephra deposit based on Voronoi tessellation).

	H_t (km a.s.l)	TEM_{DEP} (kg)	WDGS Md_ϕ	WDGS σ_ϕ
TEPHRA2	11.1 ± 1.4	$1.8 \pm 0.1 \times 10^8$	-1.1 ± 0.4	1.8 ± 0.5
Others	9.0–9.6	$1.4 \pm 0.0 \times 10^8$	–0.9	2.1

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