



**HAL**  
open science

## The Mediterranean Sea and the Gulf of Cadiz as a natural laboratory for paleotsunami research: Recent advancements

P.M. de Martini, H.J. Bruins, L. Feist, B.N. Goodman-Tchernov, H. Hadler, J. Lario, G. Mastronuzzi, L. Obrocki, D. Pantosti, Raphael Paris, et al.

### ► To cite this version:

P.M. de Martini, H.J. Bruins, L. Feist, B.N. Goodman-Tchernov, H. Hadler, et al.. The Mediterranean Sea and the Gulf of Cadiz as a natural laboratory for paleotsunami research: Recent advancements. Earth-Science Reviews, 2021, 216, pp.103578. 10.1016/j.earscirev.2021.103578 . hal-03174622

**HAL Id: hal-03174622**

**<https://uca.hal.science/hal-03174622v1>**

Submitted on 24 Nov 2021

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial - NoDerivatives 4.0 International License

## **The Mediterranean Sea and the Gulf of Cadiz as a natural laboratory for paleotsunami research: recent advancements**

De Martini, P.M.<sup>1</sup>, Bruins, H.J.<sup>2</sup>, Feist, L.<sup>3</sup>, Goodman-Tchernov, B.N.<sup>4</sup>, Hadler, H.<sup>5</sup>, Lario, J.<sup>6</sup>, Mastronuzzi, G.<sup>7</sup>, Obrocki, L.<sup>5</sup>, Pantosti, D.<sup>1</sup>, Paris, R.<sup>8</sup>, Reicherter, K.<sup>3</sup>, Smedile, A.<sup>1</sup>, Vött, A.<sup>5</sup>

1 Istituto Nazionale di Geofisica e Vulcanologia, Sezione Roma1, Rome, Italy, [paolomarco.demartini@ingv.it](mailto:paolomarco.demartini@ingv.it), [alessandra.smedile@ingv.it](mailto:alessandra.smedile@ingv.it), [daniela.pantosti@ingv.it](mailto:daniela.pantosti@ingv.it)

2 Ben-Gurion University of the Negev, Jacob Blaustein Institutes for Desert Research, Sede Boker Campus, 84990 Midreshet Ben-Gurion, Israel, [hjbruins@bgu.ac.il](mailto:hjbruins@bgu.ac.il),

3 Division of Earth Sciences and Geography, Neotectonics and Natural Hazards, RWTH Aachen University, Lochnerstraße 4 – 20, 52056 Aachen, Germany, [k.reicherter@nug.rwth-aachen.de](mailto:k.reicherter@nug.rwth-aachen.de) , [l.feist@nug.rwth-aachen.de](mailto:l.feist@nug.rwth-aachen.de)

4 Department of Marine Geosciences, University of Haifa, 199 Abu Khoushy Blvd, Mt. Carmel, Haifa 3498838, Israel, [bgoodman@univ.haifa.ac.il](mailto:bgoodman@univ.haifa.ac.il)

5 Institute of Geography, Natural Hazard Research and Geoarchaeology team, Johannes Gutenberg-Universität Mainz, Germany, [hadler@uni-mainz.de](mailto:hadler@uni-mainz.de), [L.Obrocki2@geo.uni-mainz.de](mailto:L.Obrocki2@geo.uni-mainz.de), [voett@uni-mainz.de](mailto:voett@uni-mainz.de)

6 Facultad de Ciencias, UNED, Spain, [javier.lario@ccia.uned.es](mailto:javier.lario@ccia.uned.es)

7 Dipartimento di Scienze della Terra e Geoambientali, Università degli Studi di Bari, Italy, [giuseppe.mastronuzzi@uniba.it](mailto:giuseppe.mastronuzzi@uniba.it)

8 Université Clermont Auvergne, CNRS, IRD, OPGC, Laboratoire Magmas et Volcans, F-63000 Clermont-Ferrand, France, [raphael.paris@uca.fr](mailto:raphael.paris@uca.fr),

## **Abstract**

After the 2004 Indian Ocean (IOT) and the 2011 Tohoku-oki tsunamis, new research in tsunami-related fields was strongly stimulated worldwide and also in the Mediterranean. This research growth yields substantial advancements in tsunami knowledge.

Among these advancements is the “Paleotsunami” research that has marked particular progress on the reconstruction of the tsunami history of a region. As an integration of the historical documentation available in the Mediterranean and the Gulf of Cadiz areas, geological and geoarchaeological records provide the insights to define the occurrence, characteristics, and impact of tsunamis of the past. Here, we present the recent advancements done for both the onshore and offshore realms.

As for the onshore, we discuss case studies dealing with recent high-resolution works based on: a) direct push *in situ* sensing techniques, applied to identification and characterization of typical paleotsunami deposits features; b) combined XRF-X-CT approach, implemented for the identification of fine-scale sedimentary structures useful for the definition of the causative flow dynamics; c) the geoarchaeological “new field” contribution, with the development of specific diagnostic criteria in search for tsunami impact traces in archaeological strata; d) comparison of multiple dating methods and of different modeling codes for the definition of the potential source for the displacement of boulders of exceptional dimension, identified by 3D size calculation.

As for the offshore advancements, we present case studies focusing on the recognition of tsunami deposits and their sedimentary traces in the geological record from the nearshore, thanks to diver-operated coring equipment, down to the continental slope, by means of vibracorer and long gravity core sampling in deeper areas. The examples provided show a multiproxy approach with a high potential of retrieving a complete record of paleotsunami traces at least during the Holocene. This is based on the combination of multidisciplinary approaches including X-ray imaging, high-resolution measurement of physical properties, X-ray fluorescence data, grain-size analysis, micropaleontology, palynological content, isotopic and optically stimulated luminescence dating methods.

## **Keywords**

Paleotsunami deposits; Mediterranean Sea; High-resolution studies; Geology; Archeology; backwash wave

## **Introduction**

The occurrence of a new tsunami event, how tragic it may be, is a unique opportunity to provide scientists with new critical information that increases knowledge on the different aspects of the phenomenon and to reduce the negative impact of future events on coastal societies.

Geological signatures of the impact of tsunamis of the past are commonly identified in coastal archives based on coring, trenching and subsurface geophysical methods in combination with detailed sedimentological, geomorphological, geochemical, geophysical and microfaunal analysis (e.g. Dominey-Howes et al., 2000, Mastronuzzi et al., 2007, Goodman-Tchernov et al. 2009, De Martini et al., 2010, Vött et al., 2011, Cuven et al., 2013, Pilarczyk et al. 2014, Costa et al., 2016, Chagué-Goff et al., 2017, Finkler et al., 2018a, 2018b, Werner et al., 2018a).

Paleotsunami research was initially developed in the 1980s (Atwater, 1987) in the NW of the United States (Cascadia), in Scotland (Dawson et al., 1988) and in Japan. Because of the youthfulness of the studies on the geological signatures of tsunamis and the relative rarity of large tsunamis, prior to 2004 (the year of the Indian Ocean Tsunami) the number of direct observations of the effects they produced in the natural environment was very limited; this is because most of the attention was focused on the impact on the coastal human environment. Conversely, the enormous amount of detailed observations, collected by various teams and through different techniques, after the 2004 Indian Ocean (IOT) and the 2011 Tohoku-oki tsunamis, clearly highlighted how the occurrence of a tsunami represents a unique opportunity to define the conditions under which tsunami geological records were formed and preserved in geological archives, establish their significance and parametrization. This understanding has a large potential to evaluate the hazard and in perspective to contribute in tsunami risk reduction; in fact, the geological record of past tsunamis is not only a precious source of information on the extent of inundation along the coast and inland and on run-up size, but also on the age of the event(s). These are all useful pieces of information for evaluating past tsunami size, their recurrence at any particular site, and, to potentially define the source location and mechanism when incorporated with modeling approaches. This path of paleotsunami research can be summarized in a few words: studying the present to understand the past, studying the past as a key to the future.

### **1 - Tsunamis in the Mediterranean: historical data, paleotsunami records and potential tsunamigenic sources**

The Mediterranean region is particularly rich in paleotsunami data, and this is due to the availability of historical records that provided precious clues to geologists and archeologists in search for “the direct evidence of tsunamis” (eg. sediments, morphological signatures, destruction layers, etc.; e.g.: Scheffers and Kelletat, 2003; Mastronuzzi, 2010; Anzidei et al., 2014; Scardino et al., 2020).

Tsunamis in the Mediterranean may originate from large earthquakes as well as active volcanoes and significant submarine landslides. In fact, the Mediterranean area is quite a complex active zone of deformation with different geodynamic

processes taking place from west to east and well depicted by the instrumental seismicity record (Figure 1.1).

The Mediterranean area is characterized by NW-SE oriented convergence occurring between the African and Eurasian plates, at the present rate of about 5 mm/yr (DeMets et al., 2010). Important crustal deformation is present in the *Western Mediterranean* where the inner orogenic domain is represented to the west by Rif-Betic belt and the Gibraltar Arc up to the northern Apennines and the Calabrian Arc eastward, while a number of extensional basins form the back-arc domain from the Alboran basin to the west to the Tyrrhenian basin eastward (Faccenna et al., 2014). Convergence is distributed among a number of active tectonic structures with the only subduction zone being still active the Western Ionian Sea under the Calabrian Arc (Faccenna et al., 2014, see Figure 1 for the plate tectonic setting of the Mediterranean).

*The Eastern Mediterranean* is characterized by important active subduction under the Hellenic Arc and under the Cyprean Arc (Angelier et al., 1982, Mascle and Martin, 1990). The Hellenic Arc and Trench system as well as its back-arc region of the Aegean Sea are seismically known to be very active while in central Greece the Corinth Rift, with an extension rate of 5-15 mm/yr (Avallone et al., 2004) is characterized by intense crustal seismicity too.

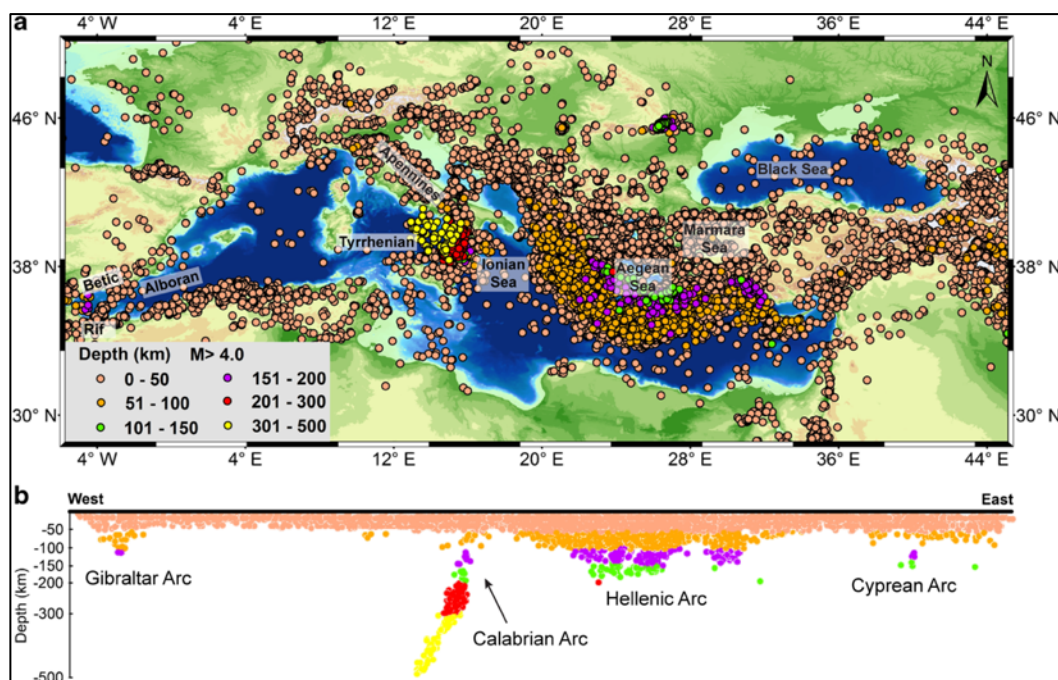


Figure 1.1: a) Instrumental seismicity in the Mediterranean region between 1900 and 2018 (source: <https://earthquake.usgs.gov/earthquakes/search/>) and depth of the earthquakes. B) Color-coded depths of the earthquakes according to a scale in km (modified after Ulutas, 2020).

The *Marmara Sea* is controlled since the Late Pliocene by the very active dextral strike-slip North Anatolian Fault (NAF) and its branches (Yaltirak, 2002). The *Black Sea* region presents along its northern margin an active subduction zone, responsible for the formation of the Caucasus (Robinson, 1997).

Tsunamis may originate from several of these active zones; for example, in the historical past (Soloviev et al., 2000; Maramai et al., 2014), relevant tsunamis were produced by earthquakes originated in the Alboran basin (1522 CE), Ionian Sea (1169, 1693, 1908 CE), Hellenic Arc and Trench system (365 and 1303 C.E.), Cyprean Arc (1222 CE), Levantine Sea (551 CE).

Tsunamis related to active volcanism and landslides represent an even larger challenge in terms of hazard assessment and risk mitigation (Paris, 2015). In fact, not being necessarily anticipated by earthquakes or other defined precursory phenomena, alert cannot be issued and risk reduction can be based only on hazard scenarios, where they exist. For example, the December 22, 2018 flank collapse of Anak Krakatau volcano (Indonesia) that generated a tsunami that devastated the coasts of the Sunda Strait, killing more than 430 people, is a clear example of this criticality (Giachetti et al., 2012, Gouhier & Paris, 2019; Grilli et al., 2019; Takabatake et al., 2019). In most of the cases, tsunami warning systems worldwide are structured to deal with earthquake-generated tsunamis only.

There are several volcanoes in the Mediterranean that could potentially generate tsunamis (e.g. Etna, Vesuvius and Phlegrean Fields coastal volcanoes – Ischia, Linosa, Pantelleria, Eolian islands in Italy and Nisyros-Yali, Santorini-Kolumbo volcanic islands in Greece, Figure 1.3). Stromboli, for instance, is one of the most tsunamigenic volcanoes in the world, with 5 tsunamis described during the 1900s (Maramai et al., 2005a) and a well-documented landslide tsunami in 2002 (e.g. Bonaccorso et al., 2003; Maramai et al., 2005b; Tinti et al. 2006). The recurrent instability of Italian volcanoes such as Stromboli, Etna, Vesuvius or Ischia is well highlighted by a number of submarine debris avalanche deposits (e.g. Pareschi et al., 2006a; Milia et al., 2007; Romagnoli et al., 2009; De Alteriis et al., 2010). Apart from slope instabilities, there are other mechanisms that can be implied in the generation of volcanic tsunamis (Paris et al., 2015): *i.e.*, pyroclastic flows and lahars entering the water, caldera collapse, volcanic earthquake, and underwater explosion. It is thus often difficult to identify the specific process causing tsunamis of the past related to large explosive eruptions, as illustrated by the debates on the 1883 Krakatau tsunami (e.g. Carey et al., 2001; Maeno & Imamura, 2011; Paris et al., 2014) and Late Bronze Age eruption of Santorini (e.g. McCoy & Heiken, 2000; Pareschi et al., 2006b; Bruins et al., 2008; Novikova et al., 2011; Nomikou et al., 2016; Werner et al., 2018b). Nowadays there is a particular concern about tsunami hazard of a future eruption of Kolumbo submarine volcano, 7 km north-east of Santorini (Nomikou et al., 2014). An analogue of the possible future eruption could be found in Ulvrova et al. (2016) that used numerical simulations together with the spatial distribution and composition of tsunami deposits on the coasts of Santorini Island, trying to constrain the source of the tsunami observed during the 1650 CE eruption of Kolumbo.

For what it concerns the hazard posed by submarine landslides in the Mediterranean, the challenge is even larger. The basic information to evaluate this hazard and define tsunami scenarios are landslides location and size along with the estimate of their potential for producing a tsunami and frequency of sliding. An updated distribution of submarine landslides in the Mediterranean basin is provided by the relevant work of Urgeles and Camerlenghi, 2013 (see Figure 2 pp. 2602).

A large variety of tsunami data are available for the Mediterranean region and its connected seas. These include: a) historical reports (see Soloviev et al., 2000;

Maramai et al., 2014, for a unified catalogue “The Euro-Mediterranean Tsunami Catalogue, containing 290 tsunamis generated in the European and Mediterranean seas since 6150 BCE to current days (Figure 1.2), [http://roma2.rm.ingv.it/en/facilities/data\\_bases/52/euro-mediterranean\\_tsunami\\_catalogue](http://roma2.rm.ingv.it/en/facilities/data_bases/52/euro-mediterranean_tsunami_catalogue), this catalogue is the result of a systematic and detailed review of all the regional catalogues available in literature covering the study area, together with about 900 documentary sources, including historical documents, books, scientific reports, newspapers and previous catalogues); b) the Global Historical Tsunami Database run by NOAA for information on tsunami events from 2000 BCE to the present available at [https://www.ngdc.noaa.gov/hazard/tsu\\_db.shtml](https://www.ngdc.noaa.gov/hazard/tsu_db.shtml)); c) onshore and offshore geological signatures (see the ASTARTE Paleotsunami deposits database - North-eastern Atlantic and Mediterranean (NEAM) region - V. 2017.2, available at <https://arcg.is/airGy>, the collected and analyzed data for the Mediterranean area alone can be summarized as follow: 66 sites and 120 geological evidence of tsunamis; 11% of sites exceeds the value of 10 m in elevation a.s.l., while the inundation distance of 17% of the sites overcomes 500 m inland); d) modern instrumental records.

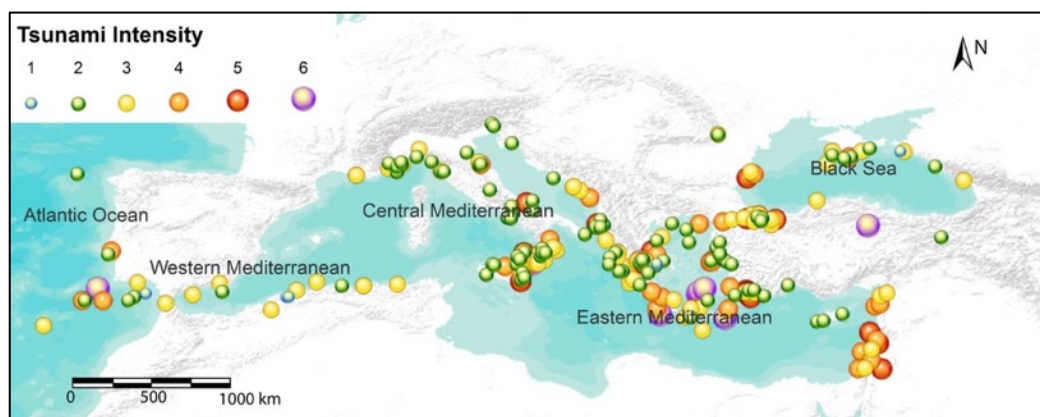


Figure 1.2: Distribution of the historical tsunami occurred in the Mediterranean region and its connected seas, please note that the plotted coordinates of each tsunami event refer to those of the triggering cause, i.e. the coordinates of the tsunami generating earthquake, volcanic eruption or landslide.

As for the potential tsunami seismic sources a review can be found in the European Database of Seismogenic faults (Basili et al., 2013) available at <http://diss.rm.ingv.it/share-edsf/>, (Figure 1.3), and in Petricca and Babeyko, 2019, and references therein.

The 2011 Tohoku-oki disaster highlighted the need for incorporating information from tsunami deposits (Sawai et al., 2012), complex geometry of the source, and quantification of source uncertainties in probabilistic tsunami hazard analysis (PTHA) studies (Selva et al., 2016). Recently, a serious effort was done to provide Probabilistic Tsunami Hazard Map for the NEAM region, see Grezio et al., 2017, and references therein, and for region-wide long-term homogenous PTHA for NEAM area derived purely from seismic sources, more is available at <http://ai2lab.org/tsumapsneam/interactive-hazard-curve-tool/>. Hopefully, in a



near future, tsunami deposits information will be incorporated also in the Mediterranean PTHA, following the Tohoku-oki lessons.

PTHA methods for landslide and volcanic sources of tsunamis are far less established compared to the earthquake sources (Grezio et al., 2017). However, there are efforts in this direction as for example, Paris et al. (2019) proposed a conditional probabilistic hazard analysis for tsunamis generated by volcanic explosions in the submerged part of the Campi Flegrei caldera (Italy).



Figure 1.3: Distribution of seismogenic faults, composite seismogenic sources and active subduction zones (bold black line and gray zone represent fault top trace and fault dipping plane, respectively), modified from the European Database of Seismogenic faults (Basili et al., 2013). The active volcanoes mentioned in the text above are plotted as red squares.

A very critical issue for the management of tsunami hazard and alert in the Mediterranean, is related to its tectonic and geographical arrangement. In fact, in the Mediterranean, all the potential tsunami sources are in the near-field domain and thus, tsunami wave travel times usually do not exceed 1 hour. Because of this, given the limited potential of tsunami alerts in the Mediterranean area, tsunami hazard assessment and prevention actions become the strongest and effective tools for tsunami risk reduction. Under this light, the knowledge of the occurrences of the past, and specifically the history of paleotsunamis that may span back in time several millennia, become a critical contribution to highlight the coastal areas most prone to tsunami inundation, the frequency of these events, and the potential source type (e.g., Okal and Synolakis, 2004).

The Mediterranean area is an ideal open laboratory for paleotsunami research for multiple reasons. First, while it is a relatively small and closed sea it contains nearly all of the tsunami-generating features of an open ocean, including tectonically active plates, volcanoes and submarine landslides. Moreover, the Mediterranean area appears to be ideal also to face the problematic storm vs. tsunami source differentiation, which is very critical in the reconstruction of a



paleotsunami history at a specific site. In fact, its semi-closed morphology coincides with moderate fetch lengths and its climatic conditions are dominated by hot and dry summers, rainy and moderately cold winters. Such conditions disfavor the formation of very large low-pressure systems. The formation of common cyclones in the Mediterranean Sea is known to be favored in front of Morocco coast (Northern Africa), in the Gulf of Genoa (Italy) and around the island of Cyprus (Figure 1.4), while tropical-like cyclones (known as Medicanes or Mediterranean Hurricanes) may primarily generate in the western Mediterranean and in the Ionian Sea, less frequently in the Aegean and in the Levantine seas (Figure 1.4). Medicanes, with diameter usually less than 300 km and an intensity commonly much weaker than tropical cyclones, are usually accompanied by strong winds, heavy rainfall, rare tornadoes and, secondarily, flood (Nastos et al., 2018, and references therein). Even if the Gulf of Cadiz is known to be under the influence of frequent North Atlantic storm waves up to 4 m high, from W-NW (Costa et al., 1992), only a few tempestites are documented onshore (Andrade et al., 2004; Kortekaas and Dowson, 2007). Thus, potential storm/medicane related deposits in the Mediterranean and Gulf of Cadiz coastal areas certainly exists but some physical criteria for attempting distinguishing between storm and tsunami deposits can be used (Morton et al., 2007, and references therein). More recently, reviews of both instrumentally and historically recorded tsunami and storm data also support the position that tsunami traces in the geological record are stand-alone seismo-tectonic signals and cannot be re-interpreted as climate signals (Marriner et al., 2017; Vött et al., 2019b).

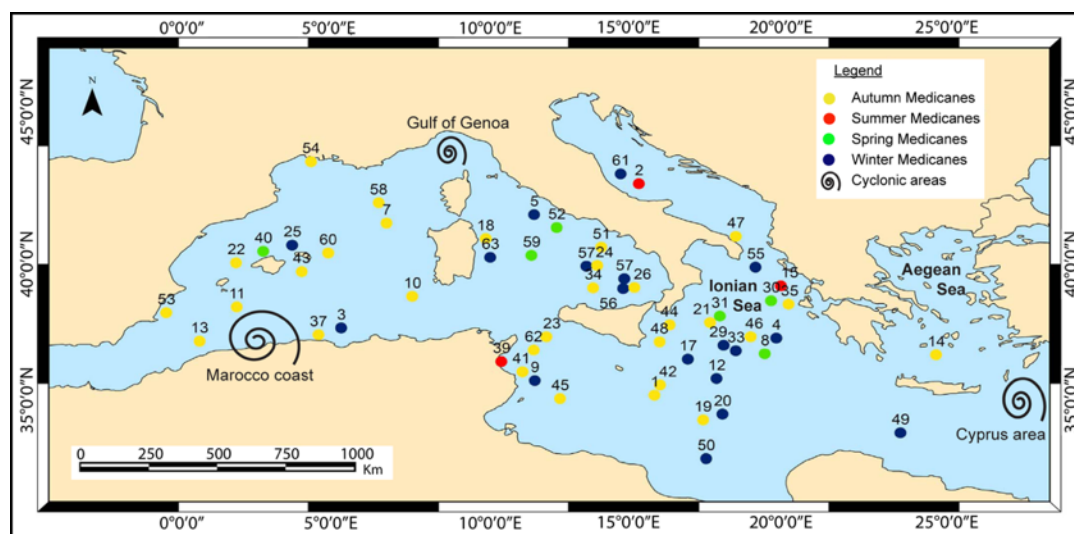


Figure 1.4: Geographical distribution of main cyclonic areas (spiral) and seasonal medicane occurrence for the time period 1969–2014 (yellow color for autumn, red color for summer, green color for spring and blue color for winter) based on 63 numbered cases (modified after Nastos et al., 2018)

## 2 – 2011 Tohoku-oki tsunami lessons

The Mw 9.0 earthquake that occurred on March 11, 2011, offshore Honshu Island (Japan) triggered a tsunami with disastrous effects. Official fatalities were more than 15,899 and missing about 2,529 (at 10 June 2020, National Police Agency,

Japan). Some reinforced concrete buildings and many tsunami barriers were severely damaged (Mori et al., 2011). A large variety of inundation typologies were observed including very different environments from agricultural lands, to coastal rural towns and modern urban cities, interacting with a local coastal morphology being substantially different across the affected area. From optical, thermal, and SAR data (Chini et al., 2013) it was possible to observe a maximum value of inundation of about 6 km, in the central portion of the Sendai plain, with a distance tapering to about 1 km at the edges of the plain, that are characterized by high-relief morphology. Moreover, the general trends of inundation distance and run-up are both “parabolic” with opposite convexity, highlighting an anticorrelation of the two distributions (Chini et al., 2013). In fact, field surveys reported that while on the Sendai Plain the maximum inundation height observed was 19.5 m and the water propagated more than 5 km inland (Goto et al., 2011; Mori et al., 2011), along the ria coast north of Sendai tsunami waves were focused by narrow bays producing the largest run-ups reaching up to 40 m, with maximum values (Goto et al., 2011; Mori et al., 2011). Moreover, run-up heights greater than 10 m and 20 m were found along 425 km and 290 km of coast length, respectively (Mori et al., 2011).

Detailed post-event surveys (Goto et al., 2011) provided a unique opportunity to put solid basis for the interpretation of the tsunami effects on the natural environment, particularly relevant for the study and characterization of paleotsunamis. The 2011 tsunami inundated up to 4.5 km inland along a transect adjacent to Sendai Airport, but the sand-dominated deposit (thicker than 0.5 cm) reached approximately 2.8 km in land (Goto et al., 2011) with a tsunami mud-dominated thin layer traceable up to the inundation limit. Similar observations were also collected 10 km north of the Sendai airport as well as 40 km southward where the tsunami inundated up to 4 km while the sand deposit left terminates at about 3 km (Goto et al., 2011). The tsunami deposits, a massive sand, graded or laminated, often with soil rip-up clasts, generally thin inland with a maximum thickness of about 30 cm in the seaward areas (Goto et al., 2011), in good agreement with the observations collected on modern subaerial sand beds deposited by major tsunamis (Morton et al., 2007). Moreover, in flatter coastal areas like the paddy fields, a centimeter-thick mud drape deposited on top of the tsunami sand probably from suspension load (Goto et al., 2011). Mud sediments were investigated in detail and a high concentration of salt was measured as a result of seawater evaporation (Goto et al., 2011). The preservation of the marine geochemical signature can be problematic especially in sandy deposits (Szczeniński et al., 2007; Chagué-Goff, 2010; Chagué-Goff et al., 2017), but it is effective or at least less problematic in organic-rich sediments (Chagué-Goff, 2010; Chagué-Goff et al., 2017).

In summary, the 2011 Tohoku-oki tsunami showed that the inundation distance (about 5-6 km in the Sendai Plain) was larger than the 2011 tsunami sandy deposit traces (about 3-4 km inland), and that the 2011 geochemical signature left by the tsunami water in the most internal 2 km can be used to trace the inundation limit but appears to degrade quickly mainly because rainfall resulted in the leaching of salts from the sandy sediments (Chagué-Goff et al., 2012). Therefore, this suggests, or better reinforces, the idea that geologists in search for tsunami deposits generally measure a minimum value. It also illustrates very well the possible differences between "paleo" tsunami signatures and the "present-days" events, for

which inundation maps were underestimated in several areas (Mori et al., 2011). In this respect, it is important to emphasize the primary role of detailed geomorphological studies for an effective site selection in paleo-studies.

Another lesson was learned by the many images and videos showing the occurrence of a very important backwash flow towards the ocean. This is certainly a promising area of research, as illustrated by post-event turbidites observed off Japan after the 2011 Tohoku-oki earthquake and tsunami. In particular, Usami et al. (2017) show that turbidity current generated by the 2011 tsunami backwash flow transported sand-size grains from the outer shelf or inner bay to the bathyal slope. This tsunami backwash turbidite was full of foraminifera from the continental shelf, while the earthquake turbidite, deposited below, includes bathyal foraminifera only. Moreover, Arai et al. (2013) evidenced a turbidity current formed by the downslope motion of suspension clouds of seafloor sediments stirred up by the tsunami. More data are needed but these results may suggest that there is a significant potential to test different and new approaches in order to discriminate between bathyal “paleoseismic turbidite” and “paleotsunami turbidite”, hopefully opening a new perspective in terms of tsunami recurrence.

Another lesson derived from the 2011 event include the fact that the so called Jogan 869 CE tsunami was known from both historical (Abe et al., 1990) and field studies (Minoura et al., 2001) and it has been found to be comparable or at least to have strong similarities with the 2011 Tohoku-oki tsunami (Sugawara et al., 2013). Thus, the reconstruction of the tsunami history of the past can well be used to forecast future events.

The gigantic 2011 Tohoku-oki tsunami sheds light on the understanding and estimation of three critical parameters that characterize a tsunami: inundation distance, run-up height and recurrence of inundation at any single site. But it also highlighted how an important backwash flow is with regards to its ability to leave traceable signatures in the offshore realm.

The knowledge of inundation distance, run-up height, and recurrence from the geological record is absolutely unique and critical to undertake the needed actions for tsunami hazard assessment and risk reduction. In fact, paleotsunami research is able to integrate and extend back in time the available historical record filling the possible gaps that are naturally expected in this record (e.g. Goodman-Tchernov et al. 2016, Vött et al., 2018).

The integration of paleotsunami records together with historical accounts and modern modeling yields robust tsunami hazard assessments for future events producing highly valuable scenarios as they are supported by different datasets.

In this paper we suggest that armed with the knowledge gained from research following recent events globally, the Mediterranean area may furnish important new elements not only to a better understanding of the tsunami hazard along its coasts, but also to the global research effort on paleotsunami signatures. This is particularly true given the specific peculiarities of this area, such as the relatively long and robust historical record that can serve as a guide and stimulus for the selection of research areas, and the significant contribution derived from accompanying archeological findings.

In the following, we present several examples of recent multidisciplinary tsunami deposits studies that represent the advancement developed in the Mediterranean area and Gulf of Cadiz, during the past decade or so. Several of these advancements were stimulated by the new information and understanding on onshore inundation and backwash characteristics derived from the 2011 Tohoku-oki tsunami event. In the following, these advancements are subdivided into two categories: onshore and offshore.

### **3- Onshore paleotsunami research advancements**

The onshore advancements are shown by discussing case studies dealing with: (1) recent applications of direct push *in situ* sensing techniques highlighting typical paleotsunami deposits features, with centimeter-scaled stratigraphic data, this high quality-resolution approach is and much less time-consuming with respect to common sediment drilling, allowing large area mapping (potentially useful for inundation distance problem); (2) recent use of X-ray computed tomography (X-CT) defining a detailed characterization of fine-scale sedimentary structures useful to better decipher the link between paleotsunami deposit structures and causative flow dynamics (ideal for tsunami vs storm causative source question); (3) “new” geoarchaeological approach and diagnostic criteria identifying and characterizing tsunami impact and its ramifications in archaeological strata from “normal” site formation processes; (4) new technological efforts for the definition of the potential source (storm vs tsunami) by adopting different modeling codes, 3D size calculation and absolute age of displaced boulders of exceptional size by using Terrestrial Laser Scanning data and Drone photographs as well as multiple dating methods.

#### **3.1 - Chasing paleotsunami signals in coastal sedimentary archives using Direct Push sensing techniques - case study from western Greece**

Advances in paleotsunami research were recently achieved by using direct push (DP) *in situ* sensing techniques in search of tsunami signals (Obrocki et al., 2020; Vött et al., 2019a). DP measurements provide *in situ* stratigraphic data by pushing the DP sensor through soils and sediments (Fig. 3.1.1). DP data allow the differentiation of sediment properties such as composition, grain size and material consolidation on a centimeter-scale. Therefore, DP data help improving the quality and resolution of sediment cores regarding compaction, contamination and loss of sediment effects caused during or by the drilling process and are much less time-consuming than common sediment drilling (Dietrich and Leven 2006).

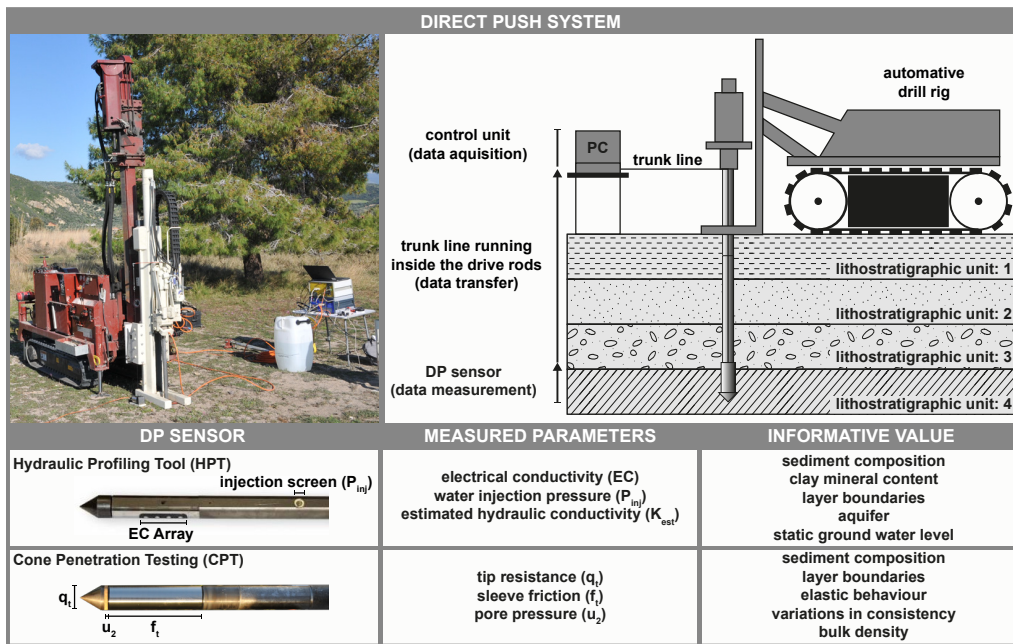


Figure 3.1.1: Illustration of the direct push (DP) system and the hydraulic profiling tool (HPT) and cone penetration testing (CPT) sensors used within the present case study.

Obrocki et al. (2020) identified two palaeotsunamite candidates at the Kaiafa Lagoon (Gulf of Kyparissia, western Peloponnese; Fig. 3.1.2) using hydraulic profiling (DP HPT) and cone penetration testing (DP CPT) tools in combination with stratigraphic information from sediment cores. Radiocarbon dating allowed the allocation of these tsunami layers as candidates for the AD 551 and AD 1303 earthquake and tsunami events, both historically well documented (Guidoboni et al., 1994; Ambraseys, 2009). In contrast, evidence of major storm events was neither found in the local sedimentary record nor it is acknowledged by historical data (cf. Vött et al., 2019b).

Multiple tsunami washover fans are described from the eastern shore of the Kaiafa Lagoon (Koster et al., 2015; Fig. 3.1.2). Typical tsunami features were found within the local stratigraphic record of such fans as, for example, (i) basal erosional unconformities and abrupt increases in grain size, (ii) rip-up clasts and mud caps, (iii) coarsening upward sequences and thinning inland. These parameters were reliably recognized by DP signals collected with DP HPT and DP CPT sensors (Obrocki et al., 2020; Fig. 3.1.1). In search of allochthonous sand sheets, the total sand content based on the same DP dataset were calculated using a simple multivariate linear regression approach (Obrocki et al., 2020).



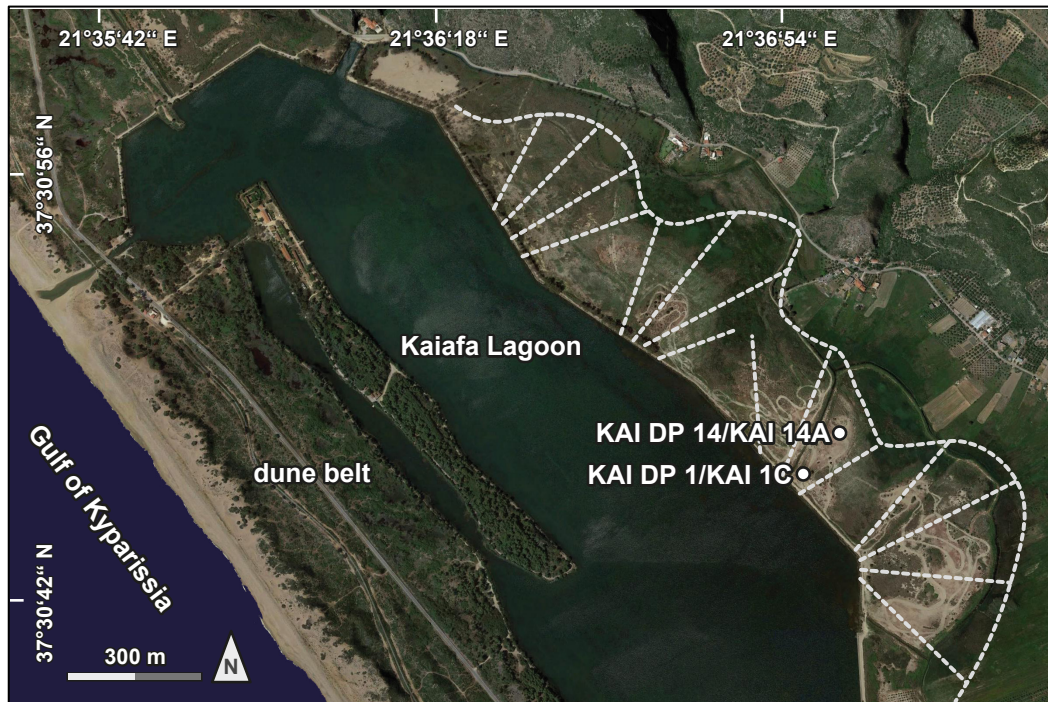


Figure 3.1.2: Topographical overview of the Kaiafa Lagoon (western Greece) with vibracoring and DP sounding sites; for further details, see text. White dashed line depicts tsunami-borne washover fans. Map based on GoogleEarth aerial images (5/5/2015).

In the following, we discuss the large potential of combined DP HPT and DP CPT parameters for the identification of tsunami layers by applying a multivariate linear discriminant analysis (LDA). In a first step, LDA can be used to find the best DP variables and, thus, discriminant functions that help to discriminate between autochthonous and allochthonous (= tsunami) deposits at the Kaiafa Lagoon. In a second step, the linear discriminant functions found are used to identify tsunami deposits in a DP test dataset (Fig. 3.1.3).

*Data collection: Direct push (DP) in situ sensing and sediment analysis*

DP logging was carried out at two sounding sites on the eastern shore of the Kaiafa Lagoon (KAI DP 1, KAI DP 14; Figure 3.1.2) using a Geoprobe 540 MT system adapted to an automotive Nordmeyer RS 0/2.3 drill rig equipped with a Geoprobe HPT K8050 (2.25-inch system) probe for DP HPT measurements and a Geotech probe NOVA for DP CPT logging (Fig. 3.1.1). DP techniques enable the constant measurements of parameters in response to the varying sedimentary conditions with a vertical resolution of 2 cm. During the HPT measurement, water is injected into the underground by a small injection screen at the side of the probe allowing for the acquisition of the water injection pressure ( $P_{inj}$ ) and the estimated hydraulic conductivity ( $K_{est}$ ). Additionally, the HPT probe is equipped with four electrodes in a Wenner array that enables the vertical measurement of the electrical conductivity (EC). The DP CPT system collects constantly the tip resistance ( $q_t$ ), the sleeve friction ( $f_t$ ) and the pore pressure ( $u_2$ ) (Lunne et al., 2002). The tip resistance describes the total force acting on the cone tip divided



by the projected area of the cone, whereas the sleeve friction is defined by the total force acting on the sleeve of the rod divided by the total area of sleeve (Fig. 3.1.1).  $q_t$  and  $f_t$  were further used to calculate the friction ratio ( $R_f$ ) and normalized friction ratio ( $F_r$ ) (Robertson et al., 2016).

In addition to DP measurements, vibracoring was conducted at both DP sounding sites. Sediment-filled plastic liners were opened in the laboratory, cleaned, photo documented and described comprising standard criteria like grain size, sediment colour, carbonate content, organic content, macrofossil content etc. Grain size analyses of fine material (< 2 mm) were carried out using the highly reliable Köhn sieving and pipette method (Köhn, 1929).

#### *Data processing: multivariate linear discriminant analysis (LDA)*

LDA is used to find one or more linear discriminant functions that allows for a dataset classification with maximum inter-group variance relative to the intra-group variance (Bahrenberg et al., 2008). Here, the following variables were used for the discriminant analysis: EC,  $P_{inj}$ ,  $K_{est.}$ ,  $q_t$ ,  $f_t$ ,  $u_2$ ,  $R_f$ ,  $F_r$ .

Data processing was realized in two steps (Fig. 3.1.3): (1) DP data collected at sounding site KAI DP 1 was grouped into six lithostratigraphic units found for vibracore KAI 1C based on sedimentary, geochemical and micro- and macrofossil data, and (2) DP data collected at sounding sites KAI DP 1 and KAI DP 14 were linearly interpolated on a 1 cm depth scale using the 'linterp' function of the 'astrochron' R package (Meyers, 2014).

LDA was then computed for KAI DP 1 (calibration dataset) using the 'lda' function of the 'MASS' R package (Venables and Ripley, 2002), in which the following equation is applied:  $D = b_0 + b_1x_1 + b_2x_2 \dots b_nx_n$

where  $D$  is the discriminant score,  $b_i$  are the discriminant coefficients and  $x_i$  are the values of the discriminating variables. The validity of LDA was determined by the reclassification of dataset KAI DP 1 and the calculation of a reclassification rate. Finally, LDA was used to derive classification rules for the grouping of the test dataset, namely DP data from site KAI DP 14 (Fig. 3.1.3).

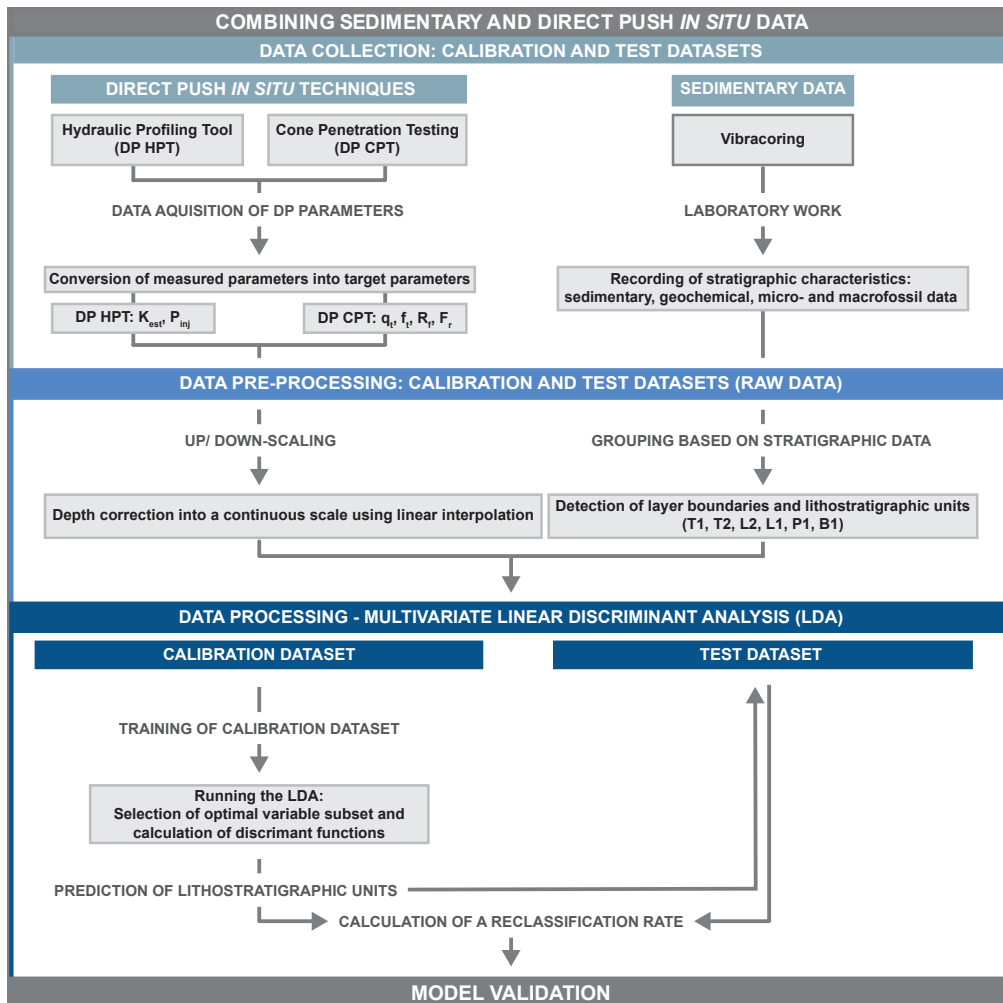


Figure 3.1.3: Combined approach of DP *in situ* and sedimentary data using a multivariate discriminant analysis (LDA) to predict lithostratigraphic units.

### Results and Discussion

Vibracore KAI 1C is characterized by a sequence of basal grey paleo-beach sands (B1), peat (P1), silty-clayey lagoonal sediments (L1), clayey silt with intersecting fine to medium sands (L2), followed by massive tsunami sands (T1) with various features of high-energy transport such as clayey rip-up clasts (T2) (Figure 3.1.4).

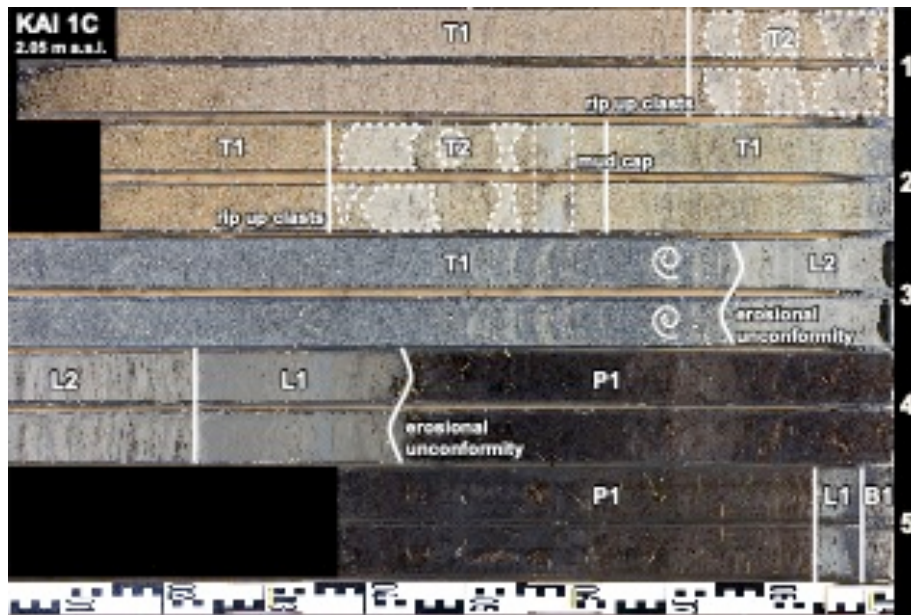


Figure 3.1.4: Photo of vibracore KAI 1C drilled on the eastern shore of the Kaiafa Lagoon. T1 – tsunami sands; T2 – tsunami-borne rip-up clasts, L2 – sandy silt to silty sands, L1 – silty and clayey lagoonal deposits, P1 – peat, B1 – beach sands. Ground surface is to the upper left, final coring depth to the lower right.

LDA was used to verify whether tsunami-related lithostratigraphic units can be identified based on DP HPT and DP CPT data. In total, five discriminant functions were derived with  $u_2$ ,  $f_t$ ,  $q_t$  and  $P_{inj}$  as the most discriminant variables and a Wilks' Lambda of  $6.2 \cdot 10^{-3}$ . The first two discriminant functions (LDA 1, LDA 2) are significant and describe 98.63 % of the variance between the six lithostratigraphic units (B1, P1, L1, L2, T1, T2). The discriminant scores for LDA 1 and LDA 2 document a clear statistical dependence of the DP values on the defined lithostratigraphic units (Figure 3.1.5 A).

The discriminatory profile of LDA 1 (eigenvalue 94.18%) allows to discriminate between high-energy units T1 and T2 with mean values of -5.14 and -4.74, respectively, and between low-energy units P1 and L1 with mean values of 8.30 and 4.62, respectively (Figure 3.1.5 C). The tsunami-induced abrupt change in environmental conditions is marked by an erosional contact at 2.77 m b.s. (Figure 3.1.4). This sharp contact is also depicted by the LDA 1 curve showing a gradual transition from unit L1 to unit L2 followed by an abrupt decrease to minimum discriminant scores of units T1 and T2 (Figure 3.1.5 B).

Lithostratigraphic units B1 and L2 can be well classified by the combined interpretation of LDA 1 and LDA 2 (Figure 3.1.5 A). With the focus on paleotsunami research, results of LDA 1 clearly improves the discrimination between units T1 and B1 (Figure 3.1.5 C). As the beach is part of the tsunami erosion zone, these units are barely discriminable using grain size analysis data alone (Röbke and Vött, 2017).

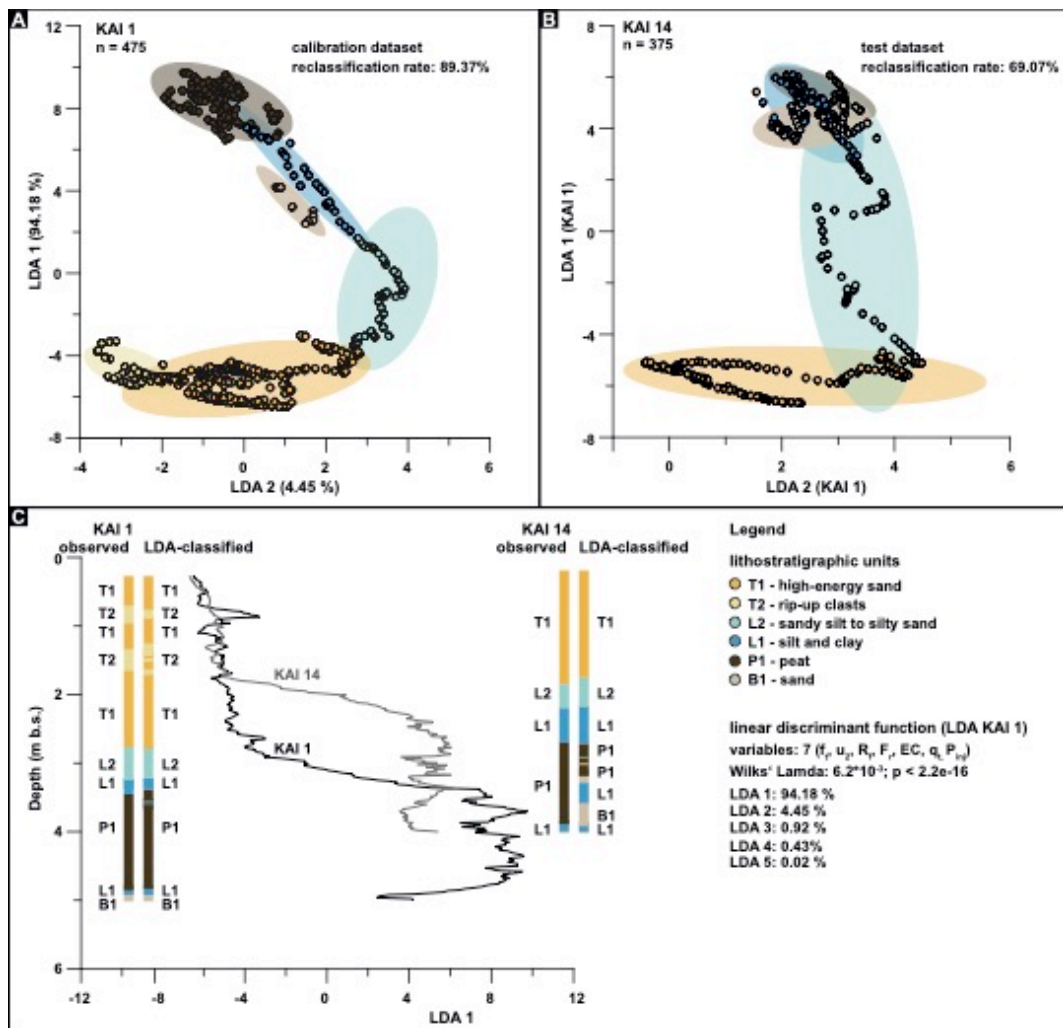


Figure 3.1.5: Results of linear discriminant analysis (LDA) of multivariate DP data with discriminant scores of the first two discriminant functions for the calibration data set KAI DP 1 (A), and the test dataset KAI DP 14 (B). (C) shows vertical discriminatory profiles using the LDA 1 function (KAI DP 1; black, KAI DP 14; grey) and predicted lithostratigraphic records – based on LDA data – compared to ‘real’ ones observed by studying sediment cores.

The high LDA validity is proven by a reclassification rate of 89.37 % for all lithostratigraphic units and 91.58 % for tsunami-related unit T1. Misclassification can be explained by inaccuracies of *a priori* grouping using stratigraphic information. Of the 62 misclassified cases, 14 occurred at layer boundaries between two lithostratigraphic units resulting in an average shift of only 2.80 cm in vertical direction (Figure 3.1.5 C). In contrast to vibracoring, DP *in situ* measurements do not entail the risk of sediment compaction (Obrocki et al., 2020). The depth offset between calibrated and predicted data may be caused by technical reasons. Other 44 falsely grouped cases were observed for the assignment of discriminant scores into unit T1 instead of unit T2 and *vice versa* (Figure 3.1.5 C). The predicted profile shows an increased number of unit T2 layers at almost the same depth but with decreased thickness compared to the calibrated profile. Rip-up clasts are usually small-scale sedimentary features (mm or cm scale, Dawson and Stewart, 2007). It seems that predicted rip-up clasts at

DP sounding site KAI DP 1 represent additional rip-up clast adjacent to the ones found at vibracoring site KAI 1C. The more precise detection and delimitation of unit T2 due to better resolved DP data compared to *a priori* grouped units must be also considered.

Finally, discriminant functions were used to classify the test dataset recorded at sounding site KAI DP 14 c. 130 m northeast of site KAI DP 1 (Figure 3.1.1). The comparison of the predicted dataset and the stratigraphic record of vibracore KAI 14A show a total classification rate of 69.07 %. Difficulties occurred for the grouping of the peat sequence P1 probably due to differences in decomposition resulting in a misclassification into units B1 and L1 (Landva et al. 1983; Missiaen et al., 2015). Figure 3.1.5 C shows that lithostratigraphic units B1, L1 and P1 are in most cases paired rather than separated based on the discriminant scores derived from LDA 1 and LDA 2. Therefore, additional parameters are needed to improve the discrimination between different low-energy units within the test dataset.

In contrast, classification rules derived from the calibration dataset were extremely helpful to differentiate between low-energy and high-energy deposits and to trace tsunami-related sediment layers within the test dataset (Figures. 3.1.5 B & C). This is confirmed by a very good reclassification rate of 93.12 % for unit T1. The LDA 1 discriminatory profiles illustrate the high efficiency of DP data to estimate the thickness of tsunami layers along transects (Figure 3.1.5 C). This allows for the rapid detection of thinning-inland of tsunami-related sand sheets and to estimate the maximum inundation extent.

To sum up, our case study shows that DP is a powerful tool to reliably identify and trace tsunami deposits in their vertical and lateral dimensions. Results of LDA prove a strong statistical relationship between DP HPT and DP CPT data and corresponding lithostratigraphic units. This allows for a clear discrimination between tsunami sediments and other facies within coastal sedimentary sequences. Calibrating DP data by sedimentary data collected at key sites clearly reduces the need of a high number of sediment cores and subsequent laboratory analysis to detect tsunami layers, keeping in mind that DP measurements are considerably quicker to realize and deliver instantaneously analytical *in situ* data. Several DP measurements can be realized within one day only. The use of DP techniques can thus significantly improve the cost and time efficiency of data acquisition within the frame of paleotsunami research.

### **3.2 - Combined XRF-X-CT applied to coastal sedimentary archives, southern Spain case study**

Information gained from historical or Holocene tsunami deposits cannot be combined with quantitative data on the inundation (flow depth and velocity, current direction, runup, etc.). That is why recent events such as the 2011 Tohoku-oki tsunami represent opportunities to better understand the causal links between flow dynamics and deposits. Sedimentary deposits left inland by the 2011 tsunami were extensively studied using different methods: (a) sediment sources estimated from foraminifera, diatom and nannolith assemblages (Pilarczyk et al., 2012; Szczuciński et al., 2012; Tanigawa et al., 2016), heavy minerals (Jagodziński et al., 2012), and geochemistry (Chagué-Goff et al., 2012); (b) current dynamics inferred from anisotropy of magnetic susceptibility

(Schneider et al., 2014); (c) maximum inundation and sand deposition limits identified using grain size analysis (Abe et al., 2012) and geochemical proxies (Chagué-Goff et al., 2012).

*The 1755 Lisbon tsunami example*

Some of these methods were also applied to tsunami deposits in the Mediterranean. The 1755 Lisbon tsunami and its deposits along the Spanish coast represent a relevant case-study. Cuven et al. (2013) used a combination of textural (grain size distribution, AMS), micromorphological (X-ray tomography, thin sections) and geochemical analyses (X-ray microfluorescence, ICP-AES) to provide a comprehensive reconstruction of the different phases of the 1755 tsunami uprush and backwash deposition in the Los Lances Bay (Figure 3.2.1). The uprush-backwash transition is marked by a switch from marine to continental sources of sediments, as evidenced by vertical variations the Ca-Al Spearman coefficient of correlation, higher K (high mud content), and lower Si and Ba concentrations (Cuven et al., 2013).

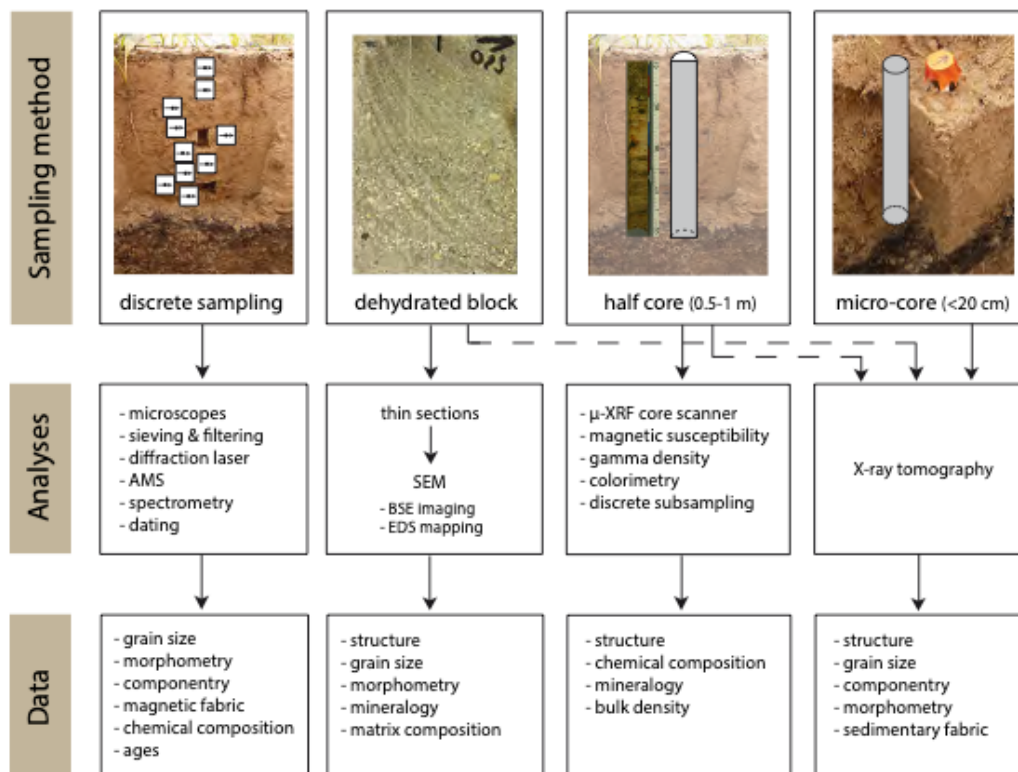


Figure 3.2.1: Summary of the methods used in Cuven et al. (2013) and Falvard & Paris (2017) for characterizing the microtexture, microstructure and chemical composition of tsunami deposit (1755 Lisbon tsunami, southern coast of Spain) using different types of sampling and analyses.

On the same deposits, Falvard and Paris (2017) later focused on X-ray computed tomography (X-CT) to provide a detailed characterization of fine-scale



sedimentary structures (erosive contacts, mudlines, rip-up clasts, lamination), vertical trends of grain size distribution and marine bioclasts, and sedimentary fabric. Falvard & Paris (2017) developed a simple method that automatically calculates grain-size metrics on two-dimensional vertical slice X-ray images, following a given vertical step (defined by the operator) along the sediment core. Lamination observed in the Los Lances deposits correspond to the spaced stratification of upper plane beds. Each new wave initiates a sediment pulse that is recorded by a distinct subunit. The base of each subunit is marked by a laminated mudline, followed by a crudely laminated and inversely graded mixture of mud and sand. The upper part of the subunit is mostly sandy and slightly normally-graded. The peak of maximum grain size thus appears approximately at the middle of the subunit, except when its upper part is truncated by the following wave. The separation of thousands of individual grains on three-dimensional stacks of horizontal slices allows three-dimensional fabric to be characterized at each vertical step. The fabric gives information on the direction and strength of the current, as well as sediment concentration and deposition rate (Paris et al., 2020). X-CT imagery also illustrates the incorporation of mud in the tsunami deposit: scouring and detachment of mud clasts, deformation of these soft clasts in the flow, formation of mud aggregates, mudlines and mud drapes. According to Falvard and Paris (2017), the mudlines that were previously interpreted as short-lived settling phases of clay-to-silt sediment, are in fact poorly sorted (cohesive mud-sand mixture) and crudely laminated. They correspond to the frictional fine-grained region of an erosive traction carpet formed by intense shear stress at the base of a highly concentrated flow. Even if the patterns identified by Falvard and Paris (2017) are probably not valid for all tsunami deposits, their results demonstrate the great potential of X-CT for defining globally applicable laws for the interpretation of tsunami deposits.

### **3.3 - Geoarchaeological Tsunami Deposits research advancements, Crete Island (Greece) case study**

The catastrophic Tohoku-oki tsunami of 11 March 2011 was an extremely powerful event, and many images were taken concerning the impact of this very large tsunami on urban environments, available on the internet. Such visual records are important to comprehend, recognize and decipher paleotsunami deposits in ancient urban centers that are investigated by archaeological excavations.

A progressing tsunami flow can be seen in Figure 3.3.1. The wave demolishes houses, partly or completely, and other human-made constructions situated in its path. The local soil/sediment surface, i.e. the geological substratum, is also partly eroded and transported by the water flow. Both human-made objects and geological/soil materials are chaotically mixed and transported forward by the tsunami. The muddled jumble is dropped as chaotic deposits along the path of its flow, first the heavier objects and gradually lighter materials as the onward and possible backward flows diminish in strength (Figure 3.3.2).



Figure 3.3.1: The front of the destructive tsunami wave flowing through houses and agricultural fields of Sendai, a large coastal city in northeastern Japan. Dark soil/sediment, building debris, and human-made objects are chaotically mixed in the progressing tsunami wave (<http://11reviews.blogspot.com/2011/04/japan-tsunami-march-201103-sendai.html>).



Figure 3.3.2: The coastal town of Minamisanriku in northeastern Japan on March 13, 2011, two days after the tsunami. Destroyed buildings and chaotic deposits, showing mixed debris lying above and inside sediments dumped by the tsunami. © REUTERS/Kyodo, The Atlantic, Alan Taylor, March 13, 2011, 33 photos in focus. (<https://www.theatlantic.com/photo/2011/03/japan-earthquake-aftermath/100023/>)

"The present is the key to the past" is a well-known concept in the geological sciences. Indeed, the recent Tohoku-oki tsunami of 2011, widely covered by the media, will undoubtedly assist to diagnose and interpret ancient tsunami deposits. Concerning the Mediterranean region, there are many ancient coastal cities that may have been hit by tsunamis in the past. However, astonishing little research has been done so far in this field. One of the reasons is that archaeologists, excavating urban sites along the Mediterranean, have usually been trained in archaeological aspects related to architecture (walls, floors) and pottery, but not in the earth sciences. Hence their insights regarding geological, sedimentological, geomorphological and soil features in archaeological excavations are usually limited. It is, therefore, no surprise that comparatively little tsunami evidence has been recognized by "urban" archaeologists. Even in New Zealand, where archaeological remains along the coast have a much more rural character, McFadgen and Goff (2007:263) noticed that "Tsunami evidence has rarely been proposed from archaeological sites, primarily because of a limited understanding of the requisite evidence and environmental context".

The recognition and study of tsunami deposits at archaeological coastal sites can be regarded as a new geoarchaeological field that needs to be developed (Bruins et al., 2008:195). Archaeological formation processes are known to a considerable extent (Schiffer, 1987). Also, a variety of paleo-tsunami signatures have been recognized and described in various deposits worldwide, including the Mediterranean region (Scheffers and Kelletat, 2004; Dominey-Howes et al., 2006). What is usually lacking in archaeological excavations is the ability to differentiate tsunami impact and its ramifications in archaeological strata from "normal" site formation processes.

#### *Geoarcheology at the Minoan town of Palaikastro, Crete (Greece)*

The following examples of ancient tsunami deposits discovered in Greece, in the Minoan town of Palaikastro (Figure 3.3.3), in Crete (Bruins et al., 2008), highlight well the potential of geoarcheology in the paleotsunami research.

The tsunami occurred during the Late Minoan IA period, at some time during the middle part of the 2nd millennium BCE, was caused by the huge volcanic eruption at the Santorini islands in the Aegean Sea, situated about 155 km northeast of Palaikastro located along the coast of northeastern Crete (Bruins et al., 2008, 2009; Bruins and Van der Plicht, 2014). In the course of the long-standing archaeological excavations at Palaikastro, the possibility of tsunami impact was first raised during 1986, when excavations of Building 2 (located in the "Exc" area of Figure 3.3.3) revealed puzzling destruction features (MacGillivray et al., 1987: 150-151). However, it was not clear whether a previous earthquake or a possible tsunami were responsible for the levelling of Building 2. About a decade later, the archaeologists Driessen and Macdonald (1997: 90) wrote regarding Building 2: "In searching for tsunami evidence, we are mostly clutching at straws. Further work is clearly required in this area". Subsequent field research by Bruins, in cooperation with MacGillivray, deputy-director of the investigations, led to detailed discoveries of tsunami deposits, found to be widespread along the Promontory and East Beach of the Palaikastro site (Figure 3.3.3).



Figure 3.3.3: The bay of Palaikastro and the surrounding landscape, looking southward. Ancient tsunami deposits and impact features were discovered in relation to archaeological stratigraphic phases and ancient buildings along the Promontory (Figure 3.3.4), the East Beach (Figure 3.3.5) and the main site of archaeological excavations, indicated by “Exc” (Photo by H.J. Bruins, taken from Mount Kastri).

Most of these deposits are characterized by a chaotic mixture of building remains, human-made objects, particularly broken pottery sherds, as well as various geological materials, including local soil and sometimes even reworked volcanic tephra from the Minoan Santorini eruption. The term “geoarchaeological tsunami deposits” was coined by Bruins in view of this mixed facies of geological materials and human-made objects. Their genesis may be paralleled with modern examples, such as visible in Figures 3.3.1 and 3.3.2 regarding the 2011 Tohoku-oki tsunami in Japan.

The following tsunami signatures were recognized in the geoarchaeological tsunami deposits (Figure 3.3.4) at Palaikastro (Bruins et al., 2008): (1) Erosional contact with the underlying strata; (2) Volcanic ash intraclasts in the lower part of the deposit; (3) Remains of building foundations and reworked building stone material in the lower part of the deposit; (4) Individual marine shells; (5) Marine micro-fauna; (6) Imbrication of rounded beach pebbles, building debris, ceramic sherds and even bones; (7) Multi-modal chaotic composition.





Figure 3.3.4: A cliff section along the East Beach at Palaikastro (Crete), showing various tsunami signatures, including a sharp unconformity (erosional contact) between the well-sorted sediments in the lower part of the picture, overlain by chaotic geoarchaeological tsunami deposits. For scale, the white stick, standing on the erosional contact, is 100 cm long. The lower part of the chaotic tsunami deposit shows remains of a building foundation to the left and many dislocated building stones. Higher up in the section occur smaller pieces of stones, beach pebbles, pottery sherds, marine shells and sometimes volcanic ash intraclasts, all mixed with fine particles of local reddish soil (Photo by H.J. Bruins).

An important diagnostic characteristic of the multimodal chaotic deposits is the predominant horizontal orientation of pottery sherds, stone slabs and even animal bones (Figure 3.3.5). This is quite uncommon in archaeological layers, as most site formation processes tend to randomize the spatial orientation of artefacts (Schiffer, 1987). Another feature visible in Figure 3.3.5, indicative of tsunami deposition, is a discontinuous layer of well-rounded beach pebbles, which were also swept along with the tsunami flows and deposited in between horizontally oriented stacked ceramic sherds. Obviously, such facies cannot be explained by “normal” archaeological site formation processes. Pottery sherds and other flat building debris are well suited, due to their shape, to obtain imbrication patterns by tsunami flows.

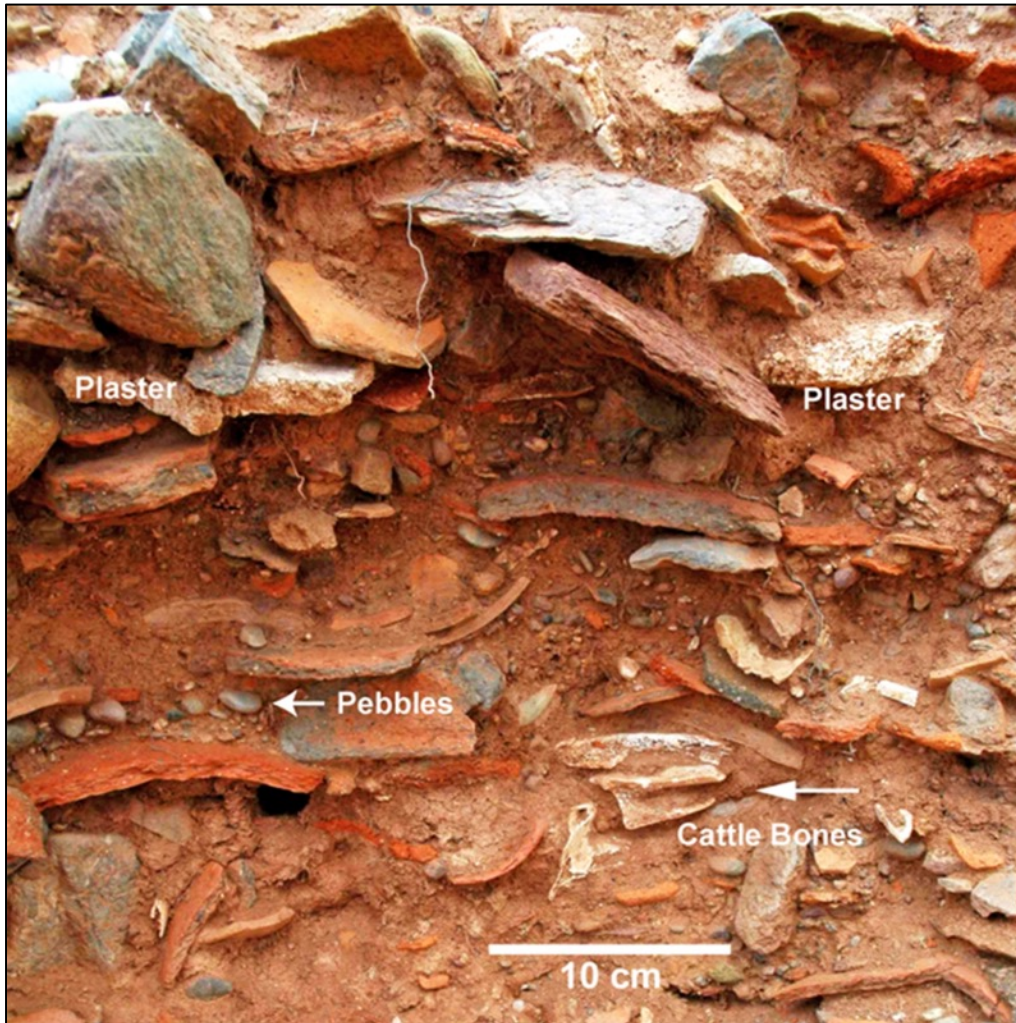


Figure 3.3.5: Chaotic geoarchaeological tsunami deposits at Palaikastro, Crete (Greece), visible at the Promontory (Bruins et al., 2008, pp198, Fig. 7). These multimodal chaotic deposits show a mixture of human-made objects originating from destroyed buildings, including stone slabs, wall plaster, pottery sherds, food remains (cattle bones), all mixed with local reddish soil and beach pebbles. Many of the objects show imbrication patterns, deposited as “roof tiles” on each other, indicative of powerful water flows (Photo by H.J. Bruins).

Considering the elevation of these deposits on the East Beach cliff and the lower sea level in Minoan times, the tsunami wave(s) must have been higher than 9 m to reach the buildings on these elevated parts with sufficient destruction impact. The backwash flow was probably less powerful on the elevated Promontory and East Beach cliff, while the remaining wall remnants of destroyed Minoan buildings helped to retain debris (compare also with Figure 3.3.2).

Although we tried during the field research to find traces of the runup level of the tsunami caused by the Minoan Santorini eruption, we were unable to detect related diagnostic signatures. The surrounding landscape, as visible in Figure 3.3.3, is on the one hand characterized by the rather steep slopes of Mount Petsofas, and on the other hand by agricultural fields on the lower slopes and in the broad valley. Both erosion and agricultural activities have probably erased runup features beyond recognition. One has to realize that, unlike modern



tsunami events, in which runup can be studied immediately, palaeotsunami research may not be able to detect all elements of a tsunami event in the landscape, after ca 3.500 years, in this case. Indeed, as noted by Bruins et al (2008) a detailed stratigraphic recording and preservation mechanism is required with sufficient resolution and memory capability in space and time, which may be lacking in the natural landscape. Ancient human settlements and archaeological excavations along the coast may sometimes provide such a detailed stratigraphic framework, in which buried tsunami signatures may become both exposed and recognizable. Micromorphology studies under the petrographic microscope of undisturbed in situ samples of tsunami deposits at Palaikastro formed an important part of the research (Bruins et al., 2008). The technique for taking samples of unconsolidated soils and sediments was developed first in the field of soil science (Osmond and Bullock, 1970), expanding in recent years into a large number of micromorphology specializations (Nicosia and Stoops, 2017). Undisturbed samples of the Palaikastro tsunami deposits were impregnated by epoxy in a specialized lab in order to facilitate the making of thin sections. Thus, their micromorphology could be studied, showing important key constituents on the micro-scale, including coralline algae, foraminifera and inter-tidal zone components, such as pieces of cemented beach-rock and rounded beach sand, which occur mixed with local reddish soil in these multimodal deposits.

The geoarchaeological tsunami layer is visible over a wide area (ca 200 m) along the Promontory and East Beach cliffs, albeit with local variations. These chaotic deposits cannot be explained by sea storms. Morton et al. (2007) made a comparative study of sandy tsunami and storm deposits. They noticed that sea storm inundation is a gradual process, involving many waves, which transport sediment primarily as bed load by traction, leaving numerous discrete layers. The differences between sedimentation by tsunamis and sea storms in the Mediterranean were presented and discussed for many locations by Vött et al. (2019). Concerning the above deposits at Palaikastro, it is clear that sea storms cannot account for imbrication of large stone objects and pottery sherds, as shown in Figure 3.3.5. Moreover, these deposits are multi-modal, containing a range of particles of different sizes from clay to building stones. Tsunami action is capable to smash, erode, transport, mix and dump such a variety of objects and particle sizes, as can be seen in Figures 3.3.1 and 3.3.2 regarding the 2011 Tohoku-oki tsunami in Japan. Anyway, powerful storms like hurricanes do not occur in the Mediterranean region, which is the main geographic focus of this article.

#### **3.4 - Boulders deposits in the Mediterranean Sea: examples from its western and central parts**

The presence along the Mediterranean coast of large boulders of exceptional size - compared to those normally accumulated inland by storm surges - was firstly reported by Pirazzoli et al (1999) along the coast of the Euboean Gulf, in Greece. Since then, numerous studies have reported anomalous boulders accumulations in several sites of the long coastal perimeter of the Mediterranean (i.e.: Mastronuzzi and Sansò, 2000; 2004; Kelletat and Schellmann, 2002; Whelan and Kelletat, 2002; Scicchitano et al., 2007; Mastronuzzi et al., 2007; Vött et al., 2008; 2019b; Scheffers et al., 2008; Maouche et al., 2009; Barbano et al., 2010; Alvarez-Gómez et al., 2011; Mastronuzzi and Pignatelli, 2012; Shah-Hosseini et al., 2013;

Lario et al., 2017; Roig-Munar et al., 2018; 2019). The particular conformation of the basin and the articulation of some coastal tracts could amplify storm waves providing them particular energy which may also be able to move boulders of exceptional sizes producing phenomena of morphological convergence (Mastronuzzi and Sanso, 2004, 2006; Mastronuzzi et al., 2006; Biolchi et al., 2016; 2019; Piscitelli et al., 2017). The interpretation of the morphogenetic process responsible for the accumulation of these unusual deposits does not find a consensus among scientists, although it is well known that large boulders have been accumulated by recent tsunamis. Eyewitnesses permitted to ascribe the accumulation of mega boulders to the impact of both 2004 IOT and 2011 Tohoku-oki tsunamis (i.e.: Goto et al., 2007; 2011; Wallis et al., 2018; Iwai et al., 2019 and references therein) however, very strong storms may cause the accumulation of large boulders too (Switzer and Burston, 2010, Cox et al., 2019). Because of this and in absence of a direct evidence, the first question to answer about boulders is about their origin: paleo-storms (tempestite) or paleo-tsunamis (tsunamite)?

In recent years, studies on large boulders focused on two main objectives that are to identify their origin and to define the hazard and vulnerability of the territory with respect to their genesis. Important results derive from studies at different sites along the Mediterranean coast: specifically (i) along the eastern coasts of Spain and of Balearic Islands (Alvarez-Gómez et al., 2011; Lario et al., 2017; Roig-Munar et al., 2018b; 2019a), both faced to the Algerian coast known as a seismic and tsunamigenic area, (ii) along the coast of Puglia (Mastronuzzi e Pignatelli, 2012), in the southernmost part of Italy, placed between the seismo- and tsunamigenic areas of the Apennine and Dinaric, Albanid and Hellenid chains, and (iii) on the island of Malta (Biolchi et al., 2016) in the center of Mediterranean basin.

### *Spanish Coast*

The Mediterranean Spanish Coast has very few geological records of EWE (Extreme Wave Events) such as tsunamis or storm surges. Some historical tsunamis have been reported in the region, however, their impact at the coast was negligible. During the past years some boulders deposits, usually with ridge morphology, have been studied (Figure 3.4.1).

Tsunami generation models show that in the Iberian Peninsula, Murcia is the province potentially affected by the greatest tsunamis, mainly generated by large seismic events in Northern Algeria. The Balearic Islands of Ibiza and Minorca also may be affected by the highest maximum wave heights passing the four-meters mark locally. (Alvarez-Gómez et al., 2011). The hazard decision matrix for tsunamis on the Spanish Mediterranean coast indicates that a 6.0-6.5 Mw earthquake less than 40 km offshore and at a depth of less than 100 km would have the potential to generate a locally destructive tsunami (IOC, 2011).

Different approaches to calculate the minimum height of the tsunami wave or storm wave capable of moving the boulders found along the coastline have been applied using formulas for a JBB scenario by Engel and May (2012) and for a submerged scenario (Nott, 2003). These models show that storms in this part of the Mediterranean are unlikely to generate these boulder ridges.



Figure 3.4.1: Location of imbricated boulder deposits in the Mediterranean Spanish Coast. Yellow points and lines show sites with boulders cited in text.

Large boulders have been found on marine cliffs in more than 20 sites on Majorca Island (Roig-Munar et al., 2019a). Deposits consist of large imbricated boulders of up to 25 t located on platforms that form the rocky coastline of the island, several tens of meters from the edge of the cliff, up to 15m above sea level. They are mainly located on the eastern and southern coasts of Majorca (Figure 3.4.1), where wave height and energy are low compared with those from the N and NE. Roig-Munar et al. (2019a) used the equations of Nott (2003) modified by Engel and May (2012), for discriminating between tsunami and storm waves, as well as the Transport Figure (TF) equation of Scheffers and Kelletat (2003). The research concludes that storms on Majorca cannot move any boulder with  $TF > 1000$ , which agrees with observations in the field that show that the biggest storm waves ever recorded in Majorca have not been able to move any of the observed boulders. Results also show that tsunami wave heights  $< 3.5$  m in the eastern and southern sectors or  $< 2$  m in the northern sector, can move boulders with  $TF > 1000$  at sea level. Therefore, tsunami appears to be the more likely mechanism for transportation and deposition of the boulder field (Roig-Munar et al., 2019a). Interestingly, the study sites are located at places where numerical models of tsunami simulation from earthquakes on the North African coast predict tsunami impact on Majorca (Alvarez-Gómez et al., 2011). Results from radiocarbon data and rate of growth of dissolution pans show a range between 1590 to 1895 CE for the ages of the boulders emplacement at the studied localities, (Kelletat et al., 2005; Scheffers et al., 2008; Roig-Munar, 2016; Roig-Munar et al., 2019a). Documentary sources also confirm a very large tsunami affecting the SE coast of Majorca in 1756, with a inundation distance up to 2 km and a run-up up to 45 m.

Although some storm waves might play a role in their dislodging, the distribution of the boulder sites along the island, the degree and direction of imbrication and the run-up necessary for their placement suggest transport from northern African tsunami waves that hit the coastline of Majorca. (Roig-Munar et al., 2019a).

Almost 25 sites with large imbricated boulders deposits have been found on platforms along the rocky coastline of Minorca Island (Figure 3.4.1). Boulders weight is up to 229 t and they are located several tens of meters from the edge of the cliff and up to 15m above the sea level (Gomez-Pujol and Roig-Munar, 2013; Roig-Munar et al., 2018a; Martín-Prieto et al., 2019).

Using the same methodology applied for Majorca Is., Roig-Munar et al. (2018a) analyzed 3000 large boulders at the studied sites and selected 720 with  $TF > 1000$ . Weight, height above sea level and distance from the edge of the cliff indicate that these boulders were dislodged and positioned by the action of tsunami waves, however, some of these were also displaced by storm waves. The numerical models of tsunami simulation from Álvarez-Gómez et al. (2011) show that earthquakes generated at the northern African coast could strongly affect the south-eastern and western areas of Minorca whereas the prevailing and strongest wind comes from the north. However, there are some boulders in the northern part of the island that cannot be explained by storm waves and could be emplaced because of the refraction of a tsunami wave originating from the northern African coast or the occurrence of submarine landslides off the Catalan platform or at the Liguro-Provençal basin platform (Roig-Munar et al., 2018a).

Large boulder deposits were found at seven sites on Ibiza and Formentera Islands (Figure 3.4.1), with a weight up to 43 t; they were located several tens of meters from the edge of the cliff, up to 11 m above sea level and several kilometers away from any inland escarpment (Roig-Munar et al., 2019b). Again, the same modeling approach was applied and results show that at least at four sites the storm conditions would not be sufficient to move the studied boulders. The analysed deposits show well-imbricated boulders with an average direction of the boulders that comply with tsunami model directions coming from Algeria (Alvarez-Gómez et al., 2011). Data from fauna incrustated in some boulders show a radiocarbon dating 1875 CE while measures of dissolution basin pools in Formentera boulders give an age of 1792 CE (Roig-Munar et al., 2019b), close to the historical tsunami that also affected other Balearic Island.

Lario et al. (2017) studied coastal deposits of blocks and megablocks in Cope (Figure 3.4.1), Murcia (SE of Iberian Peninsula). There, blocks greater than 1 m and located up to 4 m a.s.l. were piled up by an EWE, either a tsunamis or great storms.

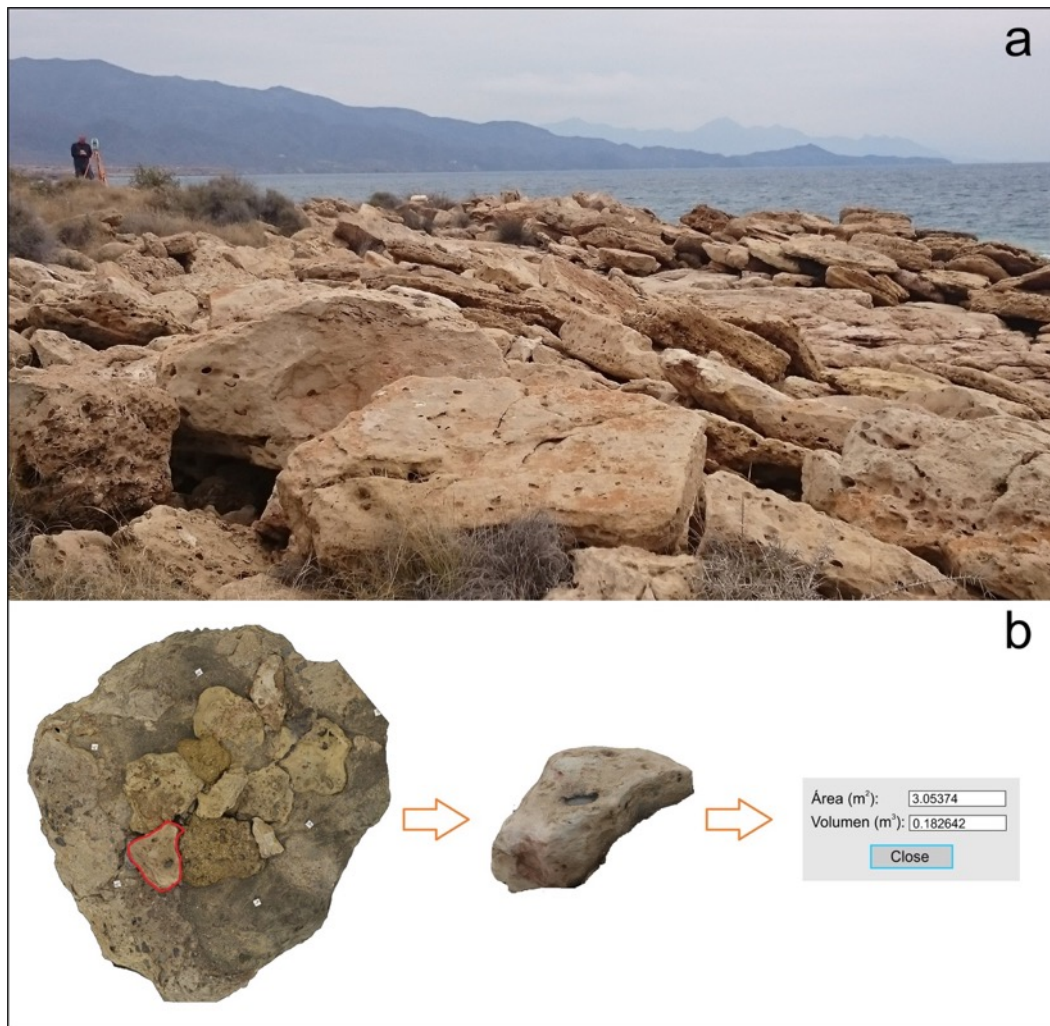


Figure 3.4.2: a) Imbricated boulder ridge in Cabo Cope (Murcia, SE Spain). Maxima altitude is 4 m asl. b) Boulder photogrammetry to calculate area and volume of each boulder using Agisoft Photoscan software.

An unmanned aerial vehicle (drone) was utilized to perform aerial photogrammetry, which complemented terrestrial photogrammetry of higher resolution (Figure 3.4.2). A 3D model of the terrain was developed in order to obtain the dimensional and spatial parameters of these blocks (Figure 3.4.3). Different models were used to calculate potential tsunami wave heights and scale of storms capable of generating such deposits (Nott, 2003; Engel and May, 2012). The needed wave heights for tsunamis are 2.7-3.9 m whereas for storms are 10.9 to 15.6 m. While the cyclonic storms on the coast of Murcia are of low entity, the study of maximum wave height measured in the buoy of Cabo de Palos in the period 1985-2012 shows that on 28 February 1995 ([www.puertos.es](http://www.puertos.es)) waves exceeded 4.5 m heights at least four times, with maxima of 5.6 m. In the area of Cabo Cope, the models also carried out for the period 1958-2017 shows a maximum wave height of 3.73 m reached in October 1973. With this data we can conclude that the wave heights of more than 10 m necessary to be able to mobilize the larger blocks composing this coastal bar, did not occurred during large storms hitting in the area in instrumental times. This supports the hypothesis that only a



tsunami wave can produce this boulder ridge. Dating this type of deposits is still a challenge, and in this case no useful dating materials were found.

In 1522 a  $M > 6.5$  earthquake hit Almería and affected large areas of the western Mediterranean, causing more than 1.000 deaths. The epicentral area is located offshore on the platform of the Gulf of Almería. The earthquake triggered several underwater landslides that could have produced a tsunami (Reicherter and Becker-Heidmann, 2009). No supporting evidence that this earthquake triggered the tsunami associated with the Cope deposits was found, however, that is the only one documented in the area.

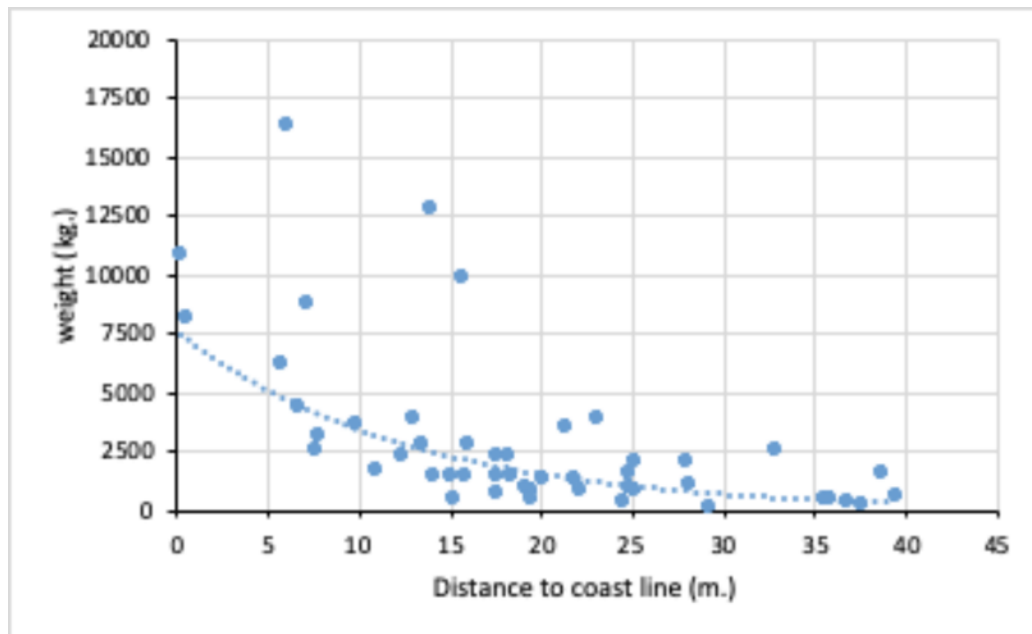


Figure 3.4.3: Relationship between distance to the shoreline and the weight of the blocks (Cabo Cope (Murcia, SE Spain)).

Six sites were identified in North Castellon, Valencia Region (Figure 3.4.1) with boulders ridges with sedimentary characteristics typical of tsunamis flows (Roig-Munar et al., 2018b). Boulders have an average weight of 1.5 t and are located at an average distance from the coast line of 15.5 m and at 2.3 m a.s.l. Using the same approach used for the previous sites the results show that local storms are not capable of moving the boulders to the present position and a tsunami wave seems to be the main cause. Also, according to the Alvarez-Gomez et al. (2011) models, the average boulder orientation shows that the coast was affected by tsunami waves likely generated in north Algeria passing through the Ibiza-Majorca channel. There is no data about the age of the events.

### *Italian Coast*

The coasts of Italian peninsula were repeatedly hit by extreme marine events that detached boulders from the adlittoral/intertidal/infralittoral area scattering them inland above the present sea level and at different distances from the coastline (Fig. 3.4.4A). In some cases eyewitnesses reported the potential of the recent strongest storms to detach, transport and accumulate boulders of different



size. Surveys performed in the recent past highlighted boulders accumulation (never exceeding 1 t.) due to storm waves in many places along the Adriatic and Ionian coast of Apulia as well as along the Ionian coasts of Sicily (Mastronuzzi and Sansò, 2004; Mastronuzzi et al., 2006; Barbano et al., 2010). In the following some cases-studies with boulder accumulation that was attributed to the impact of the tsunami are described. This is based on a multidisciplinary approach including (i) high-definition digital methodologies survey of the ground surface, (ii) search for historical sources, (iii) determination of age and (iv) hydrodynamic modelling by means of methods developed on the basis of Nott's equations (2003) (Pignatelli et al., 2009; Scicchitano et al., 2012).

Main conclusion of these studies (Figure 3.4.4) is that the coasts of Southern Italy were repeatedly hit by tsunamis generated by local offshore earthquakes or resulting from large landslides detached along the continental slope, possibly triggered by strong local inland earthquakes.

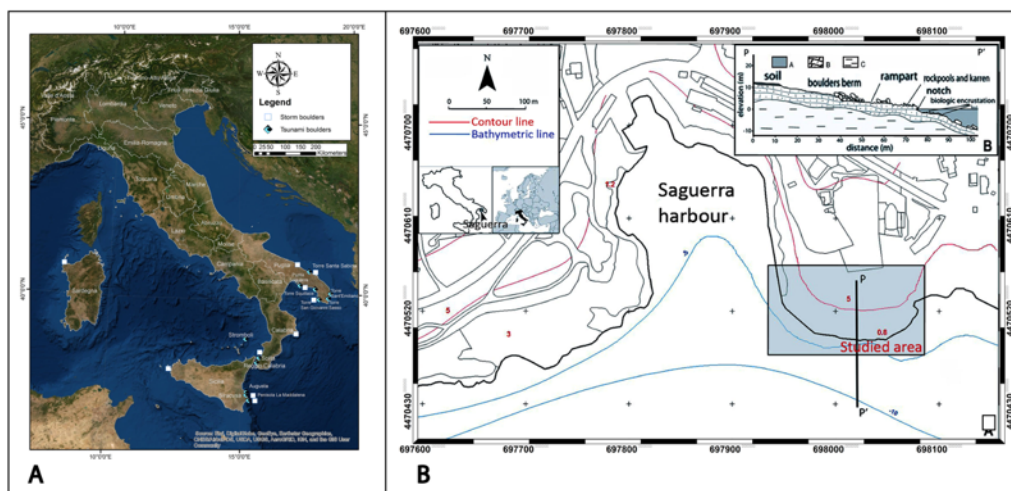


Figure 3.4.4: (A) Location of main boulders fields accumulation along the Italian coasts, and (B) the area of Punta Saguerra near Taranto

### *Apulian Ionian Coasts*

Many sites along the Ionian rocky coasts of Apulia are characterised by the presence of boulders of unusual size, isolated, sparse in fields or arranged in berms some kilometers long. Some of them were studied in detail using an interdisciplinary approach integrating digital surveys with archaeological and historical analyses and AMS age determinations. The results highlighted that the earthquake-generated tsunami occurred on February 20, 1743 emplaced boulders on the coasts of the Otranto Channel (Mastronuzzi et al., 2007) and of the Southern Adriatic (Mastronuzzi and Sansò, 2004).

In the innermost vertex of the Taranto Gulf, many rocky coasts show isolated boulders of different size. Punta Saguerra is a rocky headland gently sloping toward the sea marked by an about 200m-long berm of well cemented bio-algal calcarenite boulders (Figure 3.4.4B). This berm is composed of about one hundred boulders, up to 30 tons, which locally are arranged in rows of imbricated patterns (Figure 3.4.5). These are separated from the coastline by a quasi-flat surface at 2–5 m above present sea level (a.p.s.l). This area is marked by small coalescent

karstic potholes, giving place to pinnacle-like forms (spitzkarren) separated inward by gradually widening flat depressions.

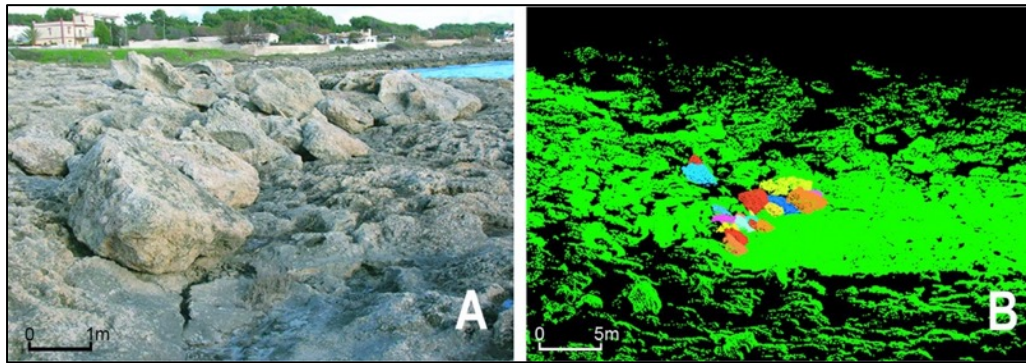


Figure 3.4.5: Boulders at Punta Saguerra near Taranto along the Ionian coasts of Apulia (SE Italy) (A) and the result of the survey (B) performed by mean terrestrial Laser Scanner;  $^{14}\text{C}$  age determinations and historical documents suggest that boulders were scraped and scattered inland by the 1836 tsunami.

In order to define the berm depositional mechanism, a detailed study of its morphology was performed using a digital 3D approach. Digital surveys have been performed using a Terrestrial Laser Scanner Leica Scanstation 2 able to scan the landscape up to about 300 m capturing up to 50,000 points per second. The post processing phase permitted the reconstruction of the 3D view as results of the overlapping of more scans (referred to the Global Navigation Satellite System of Apulia region by means of a GPS Leica 1230 in differential modality in Real Time Kinematic). The 3D digital model provided precise details of each boulder: position, size and distance from the shoreline and, indirectly, shape, weight and long axis azimuth. Applying the most recent hydrodynamic equation derived by Nott's theory (Nott, 2003; Pignatelli et al., 2009; Scicchitano et al., 2012), these data have been used to reconstruct the minimum wave able to detach and transport boulders. Comparing these values to the historical and recent sea-wave data for the area as recorded by IIM (Istituto Idrografico della Marina) and by RON (Rete Ondametria Nazionale) buoys (IIM, 1982; Inghilesi et al., 2000) it is possible to conclude that the extreme event responsible for this berm accumulation was a tsunami.

Many boulders are characterized by bio-encrustations of *Vermetids* spp and holes of *Lithophaga lithophaga*; others are characterised by solution potholes that indicate alternatively their intertidal/adlittoral/infralittoral provenience. Because  $^{14}\text{C}$  age determinations on *L. lithophaga* could present aging effects (Quarta et al., 2019), AMS analyses were performed only on *Vermetid* spp. sampled on boulder surfaces. Radiocarbon dating yielded an age of about 450 years, using a  $\Delta R$  derived from the nearest site with a value of  $121 \pm 60$  (Reimer, 2013), it is possible to attribute to this event an age in the first half of the XIX century. Chronicles from the Archiepiscopal Curia of Rossano, a city on the Ionian coast of Calabria facing the Taranto Gulf, describe the effects of strong waves that flooded the coast in coincidence of a strong earthquake on April 24, 1836. This supports the case of an earthquake triggered tsunami Based on historical reports, the earthquake epicenter is located inland, near the town of Rossano, if this is

correct,, this tsunami is likely related to a shaking induced submarine landslide along the continental slope.

### *Malta Island*

Boulders accumulations at five sites along the eastern gently sloping rocky coasts of Malta were surveyed by an international team of researchers (Biolchi et al., 2016). The survey included underwater surveys by scuba divers (Self-Contained Underwater Breathing Apparatus) and aerial surveys by UAV (Unmanned Aerial Vehicle). The Sicily–Malta channel is exposed to strong storms characterized by long fetch from NW/NE/E winds. The Malta archipelago is separated from Sicily by the Malta scarp, which is one of most active seismic structures of the area, responsible for the genesis of earthquakes and related tsunamis in the past. Even if earthquakes were felt in Malta (February 20, 1743 for example), no historical documents report the impact of tsunami on any of the three islands of the archipelago.

The sites were studied with the aim to define the mechanism that emplaced 77 boulders along the coast. Direct surveys of the boulders, age determination and hydrodynamic analysis were performed. AMS age determinations on *Vermetid*, *Serpulid* and *Chthamalidae* spp that bioencrust some boulders provide useful chronological references (Figure 3.4.6).

The geochronological data set highlighted some age-clusters that suggest different times of emplacement of these boulders. Interestingly, some of them include the ages of strong tsunamis occurred in this region of the Mediterranean, such as CE 963, 1329, 1743, 1693. However, the relation between hydrodynamic analyses (consistency between boulder sizes and actual wave features) and radiocarbon ages (some results suggest a “modern” age) seem to suggest that numerous boulders, and specifically the smaller ones, can have been scraped and scattered inland by strong storm waves.

These results underline the uncertainties to be faced in the interpretation of these signatures: on the one side, there is age correlation with tsunamis of the past for the emplacement of boulders, on the other, it is likely that boulder emplacement is related to large storms. In any case the possibility of a combined effect of storms and tsunami highlight the potential tsunami hazard for the archipelago too and call for more studies on the subject.

The geologic signatures of the 1743 and 1693 tsunamis were recognized along the coasts of the central Mediterranean Sea thanks to different evidence (washover fans, mega boulders, marine layer sandwiched in lagoonal peat layers). If these considerations seem to exclude the possibility that other tsunamis, even previous to the listed ones, may have hit the Maltese archipelago, on the other hand it must be remembered that the energy of a tsunami that impacts along the coast can cancel the evidence of previous impacts and leave only his record (Mastronuzzi et al., 2013).

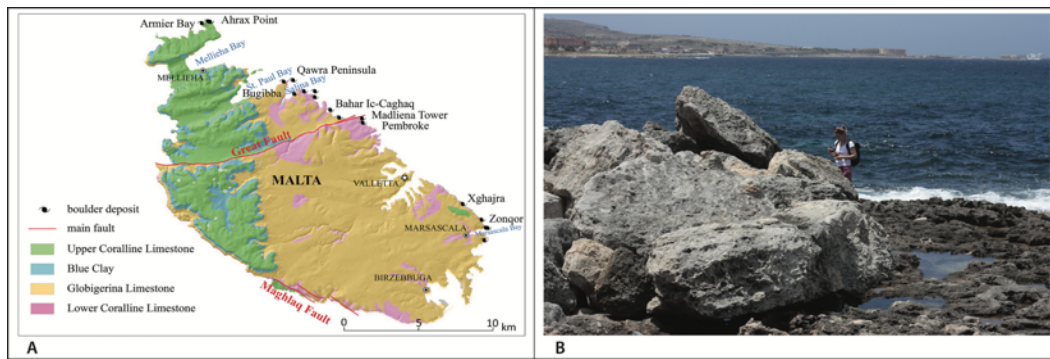


Figure 3.4.6 - (A) Location of the studied sites in Malta and, (B) boulders accumulation at Armier Bay (Photo via Stefano Furlani, University of Trieste); clusters of the 14C age determinations performed on bio-encrustations from the boulders of this site suggest these boulders may be emplaced by different historical tsunamis generated by historical earthquakes located in Eastern Sicily.

#### 4 - Offshore paleotsunami research: case studies from offshore Israel, Italy and Portugal

Offshore tsunami deposits and their sedimentary fingerprint in the geological record, although considered to offer a higher potential in terms of preservation and spatial coverage (Rhodes et al., 2006; Dawson & Stewart, 2007; Goff et al., 2012), remain less known and explored. After the 2004 Indian Ocean and 2011 Tohoku-oki tsunamis, offshore surveys were conducted to collect material, from the inner continental shelf (0 to 30 m below sea level (b.s.l.)) and over the shelf break (>100 m b.s.l.). These studies highlighted that tsunami outflow (backwash waves) had produced intense erosion, sediment transport and deposition offshore and that in the marine realm the tsunami-related units are subjected to the action of currents, waves and bioturbation immediately after their deposition (Paris et al., 2010; Feldens et al., 2012; Sakuna et al., 2012; Ikehara et al., 2014; Tamura et al., 2015; Yoshikawa et al., 2015; Seike et al., 2017).

In recent decades the offshore shallow shelf environment (from the coastline to about 30 m water depth) has shown very fruitful for preserving archives of past tsunami events, both from field evidence and modeling (Weiss and Bahlburg, 2006). While terrestrial coastal deposits are, in relative terms, easily accessible, they are also more exposed to post-deposition erosion, clean-up, coastal development (human modification), and other forms of disturbance that can nearly or entirely prevent their long-term preservation (Szczuciński, 2012). These disturbances are commonplace both in modern and ancient contexts (Dey et al., 2014). In contrast, preservation in the shallow offshore environment has shown promising as a potential sediment archive, particularly at depth, below typical storm wave base. Working in the offshore environment requires the collection of cores or samples from research vessels or by diving. This is one of the reasons that studies in this realm are very limited despite their high preservation potential. Cost and accessibility to collection sites is often challenging, and collecting long cores (>1 m) from sandy or coarser upper shelf sediments sometimes requires the use of less known coring techniques, such as diver-operated coring equipment (Figure 4.1A modified from Goodman Tchernov et al., 2016). However, large

research vessels equipped with winches and bathymetric and sediment echosounding systems can carry long gravity corers and vibracoring equipment (up to ca. 100 m water depth because of energy supply by cable from board).

For all these reasons (costly, work-intensive, ship mission, preservation, long-shore currents, etc.), the overall number of studies from this zone is significantly lower than those on land. But the potential of retrieving a complete record is much higher.

In the following, we introduce three case studies from offshore of Israel, southern Italy and southern Portugal that show how the offshore geological record is a very promising archive of past tsunamis.

#### *Offshore of Caesarea and Ashkelon (Israel)*

In Israel, studies offshore of Caesarea and Ashkelon have shown stratigraphic sequences of typical condition Nile-derived sands interbedded with anomalous coarse layers (Figure 4.1 B) that exhibit a mixed origin from shallow and deep sediments, indicating multi-directional high-velocity transport events (e.g. Goodman-Tchernov et al., 2009; Hoffmann et al., 2018; Reinhardt et al., 2006; Tyuleneva et al., 2018). On the basis of geophysical studies that reconstruct the changing bathymetric surfaces caused by these events (Goodman-Tchernov and Austin Jr., 2015), these layers were confirmed as high-energy deposits characterized by erosional surfaces. The ages of these layers coincide with the ages of documented earthquakes, or in some cases, tsunamis descriptions from historical accounts (Papadopoulos et al., 2014; Salamon et al., 2011). Initially, the lack of correlatable tsunami-derived terrestrial coastal deposits raised questions regarding the validity of the interpretation of tsunami deposits within the offshore record. However, a review of archaeological reports from land excavations in Caesarea suggests that rather than a lack of tsunami deposits there is instead a lack of interpreted tsunami deposits (Dey et al., 2014). Layers that could be clearly correlated with the offshore deposits were found in the descriptions from the excavations, these were previously defined as construction fill, refuse, dredge deposits (Neev et al., 1987), or left uninterpreted (Holum et al., 2008). Reassessment of these descriptions and records showed that they corresponded characteristically (with some variation as expected due to their different position along the coast), chronologically, and stratigraphically with the abovementioned offshore tsunami deposit sequences. In some cases, unexplained phenomena that confused the interpretations of the archaeologists such as 'fully articulated human skeleton found without any signs of deliberate burial or disturbance' found within allochthonous shell beds, and buried-while-alive edible vascular plants, took on a more logical meaning within the context of the tsunami history of the site (Dey et al., 2014).



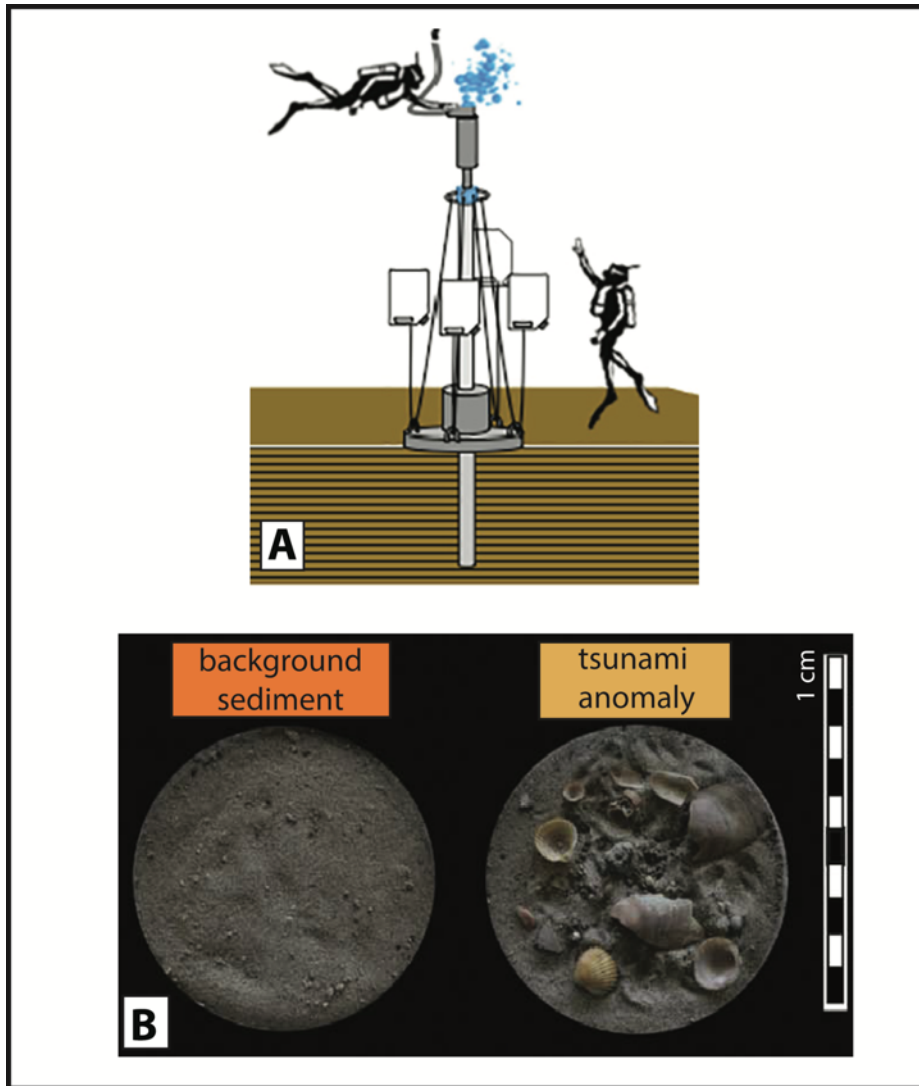


Figure 4.1: A) inner shelf results collected thanks to diver-operated coring equipment (image modified from Goodman Tchernov et al., 2016) and, B) photographs of sediments with typical condition Nile-derived sands (background sediment) interbedded with anomalous coarse layers that exhibit a mixed origin from shallower and deeper sediments (tsunami anomaly) (modified from Tyuleneva et al., 2018).

#### *Offshore of eastern Sicily (Italy)*

The Bay of Augusta in southern Italy was hit by several well documented historical tsunamis and was the focus of both onshore (De Martini et al., 2010) and offshore (Smedile et al., 2011, 2012, 2020) tsunami studies. As for the offshore studies, particular attention was posed in collecting additional insight into the expression of tsunami deposits in the middle offshore environment. In a first study, Smedile et al. (2011) presented results from a 6.7 m long, fine sediment core (MS06) sampled at a water depth of 72 m, 2.3 km offshore the Augusta Bay (SE Sicily). Research involving X-ray imaging, high-resolution measurement of physical properties, grain-size analysis, micropaleontology, isotopic dating methods ( $^{210}\text{Pb}$ ,  $^{137}\text{Cs}$  and  $^{14}\text{C}$ ) and tephrochronology was carried out, looking for any subtle



anomaly that could represent a proxy for tsunami occurrence. Results obtained from the studied core highlighted 12 anomalous intervals, marked by peaks with a high percentage of displaced epiphytic foraminifera (taxa that are normally living in the infralittoral zone on vegetated and coarse substrates, such as the seagrass *Posidonia oceanica*) and by sandy component increments, which were interpreted as the primary effect of tsunami backwash waves. The tsunami interpretation is also supported by the matching between historical tsunami accounts, both local and basin-wide ones (those of CE 1908, 1693, 1169, 365 Crete and the ca. 3.6 ka B.P. Santorini tsunami), and the age windows of five studied layers. Moreover, the ages of seven events, identified in the offshore sediments, show a positive correlation with the ages of tsunami deposits found onshore along the Augusta Bay coastline (De Martini et al., 2010), reinforcing the tsunami mechanism hypothesis.

In a recent study performed in the same area of MS06, Smedile et al. (2020) published new paleotsunami results from four one-meter-long gravity cores (MG cores), collected along a transect at 60 to 110 m water depth. In this new study, the evidence of tsunami deposits are based, as for core MS06, on sedimentology and displaced benthic foraminifera but these are reinforced by X-ray fluorescence data (Figure 4.2). Two erosional surfaces are recognized coupled with grain size increase, abundant *Posidonia oceanica* seagrass remnants and a significant amount of *Nubecularia lucifuga*, an epiphytic sessile benthic foraminifera considered to be transported from the inner shelf (see the plate in Figure 4.2). The occurrence of Ti/Ca and Ti/Sr increments, coinciding with peaks in organic matter (Mo inc/coh) suggests terrestrial run-off coupled with an input of organic matter. By indirect age-estimation methods using  $^{210}\text{Pb}$  profiles and the comparison of the Volume Magnetic Susceptibility data collected on MG cores and MS06 core, the two horizons were attributed to two distinct historical tsunamis (CE 1542 and CE 1693). The Augusta Bay offshore studies suggests two main “advances” for the offshore investigations: i) sedimentology and micropaleontology (displaced benthic foraminifera) remain the most informative techniques amongst all those applied, and can well benefit from the X-ray fluorescence (XRF) results, ii) the compared study of multiple cores collected at different depths along the shelf provided important insights about the dynamics of tsunami backwash and its imprint in the offshore stratigraphic sequence not easily gained from a single core.

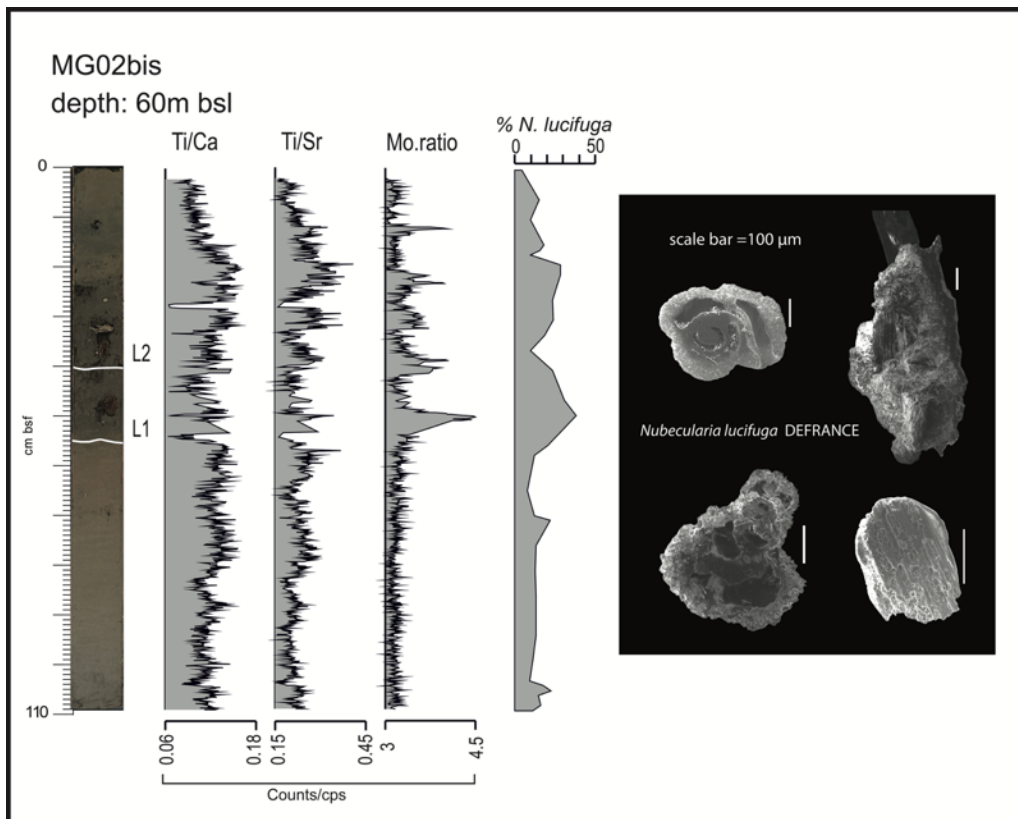


Figure 4.2: Outer shelf tsunami recovery from a gravity core (modified from Smedile et al., 2020) showing two erosional surfaces (L1 and L2) recognized thanks to grain size increase, abundant *Posidonia oceanica* seagrass remnants and a significant amount of *Nubecularia lucifuga* (on the right a small plate with some SEM photos), an epiphytic sessile benthic foraminifera considered to be transported from the inner shelf. The occurrence of Ti/Ca and Ti/Sr increments, coinciding with peaks in organic matter (Mo inc/coh) suggests terrestrial run-off coupled with an input of organic matter.

### *Offshore of Portugal*

In CE 1755 a strong earthquake-generated tsunami caused major inundations, sediment transport and transfer, as well as significant changes of the coastline along the Portuguese coasts and in capital Lisbon. To date, mainly terrestrial coastal archives have been used to characterize various sedimentary parameters of tsunami layers. The sedimentological characteristics and paleoecological effects of the CE 1755 Lisbon tsunami and older tsunami backwash deposits have been retrieved and studied on the southwestern Portuguese shelf during the RV METEOR expedition M152/1 in November 2018, the work is still ongoing (Reicherter et al., 2019).

Before performing offshore sampling, bathymetry and sediments were mapped, as on the Portuguese shelves post-glacial deposits are only preserved in pockets (Figure 4.3). The mapping helped to find drill places with larger sediment thicknesses.

Most of the coring stations of expedition RV Meteor M152/1 off the Portuguese Algarve coast were located within areas of fine-grained surface sediments. Gravity

cores were collected in water depths larger than ca. 100 m (-to 800 m) along the Portimão canyon. Up to 6 m long sediment cores in plastic liners were retrieved with vibracoring on the Algarve continental shelf in water depth of less than ca. 100 m. At some sites duplicate and triplicate cores were collected to obtain more sediment, *i.e.*, sample material for laboratory analyses and Optically Stimulated Luminescence (OSL) dating. Additional Van Veen grab samples were taken from the seafloor surface in a transect parallel to the coast and along the continental slope extending a western transect. The aim was to identify the CE 1755 Lisbon tsunami and also older potential tsunami layers. The following scientific questions were the focus of expedition M-152:

- Do variations in thickness, sediment composition and preservation of a CE 1755 Lisbon tsunami deposit exist on the Portuguese continental shelf and slope?
- Are events predating the CE 1755 tsunami recorded along the Algarve coast; and can information on timing, frequency and magnitude of these be obtained, thus extending the time-window of observation and recurrence interval?
- What effects (e.g. on benthic life) after offshore tsunami backwash deposition can be observed?
- What are sedimentological characteristics for the identification of historic and prehistoric tsunami deposits on the shelf?

These scientific questions are addressed by means of a multiproxy approach including analyses on the grain-size, compositional, geophysical and geochemical properties (inorganic and organic), micropaleontological and palynological content and suitable dating techniques (radiocarbon, OSL).

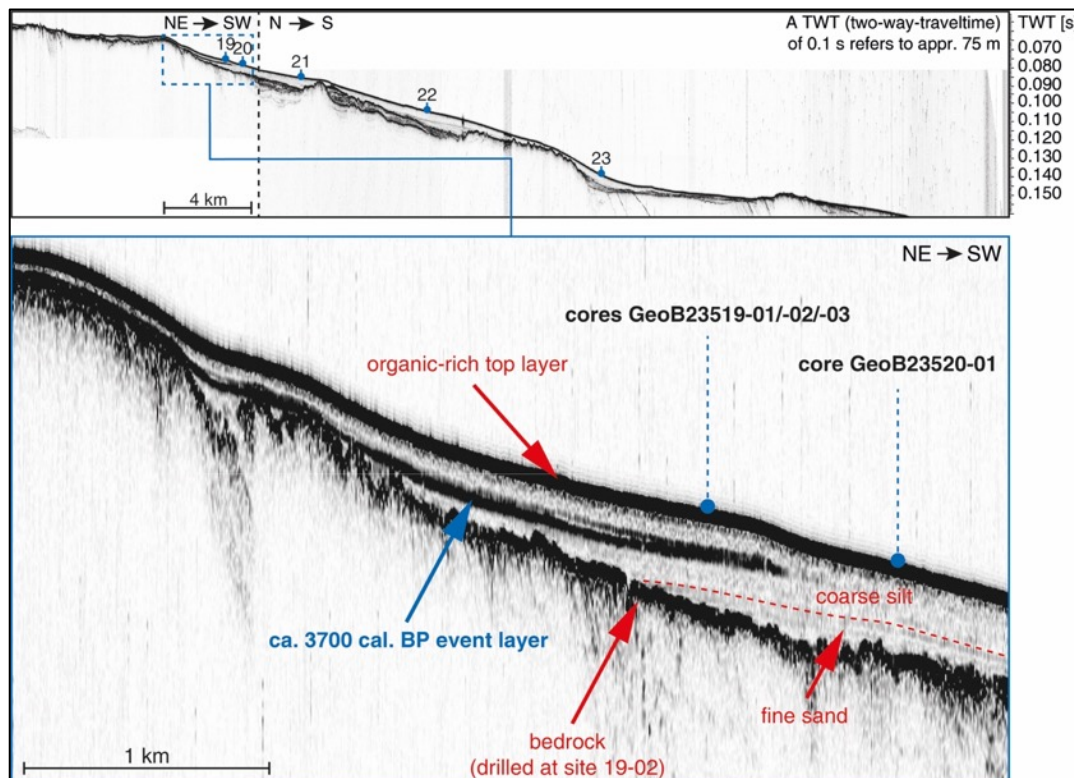


Figure 4.3: Parasound profiles of the western Algarve transect including coring stations. A two-way-travel time (TWT) of 0.1 s refers to approximately 75 m. The strong reflector resembles a ca. 3700-year-old tsunami backwash layer (see Figure 4.4).

Major results from this expedition, although preliminary, are: 1) the shallow offshore area in specific settings on the southern Portuguese shelf provides a reliable sedimentary archives for tsunami events of the past, at least during the Holocene, 2) unknown tsunami events in prehistoric times via sedimentary features off southwest Portugal were identified (Figure 4.4), and, 3) the tsunami catalogue of Portugal was extended back in time and this represent a contribution to the risk and hazard assessment of western Europe (Lario et al., 2011; Reicherter et al., 2019).

In conclusion, the results shown for Portugal, as well as those shown off Israel and Sicily indicate the shelf to be a suitable archive for paleotsunami deposits due to good preservation favored by mild hydrodynamic conditions. Well-preserved tsunami backwash signatures may only be found in specific offshore geomorphological settings, such as small depositional basins with limited bioturbation as this can destroy especially thin deposits. Even though offshore backwash deposits have been neglected in tsunami research, in future studies, these should be considered (in the right geomorphological setting) as a reliable archive of tsunami events and a possible extension of the onshore record.

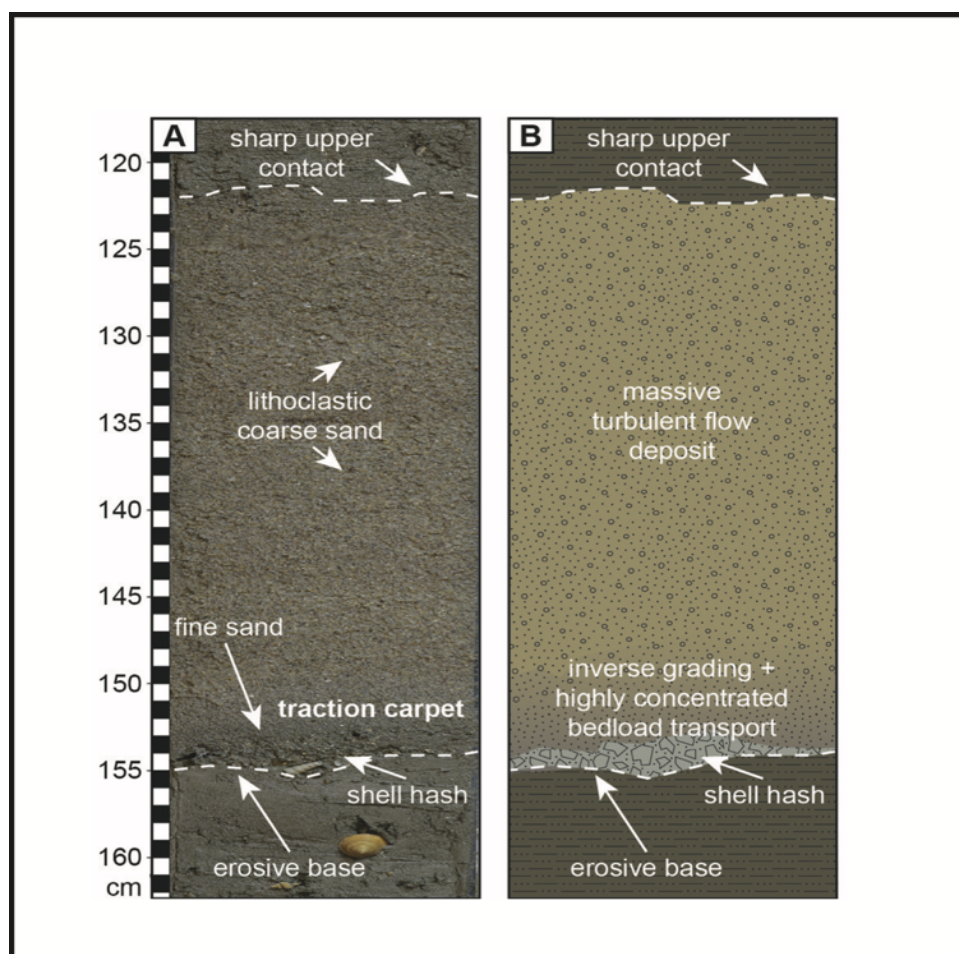


Figure 4.4: Outer shelf (65 m water depth) example of a ca. 3700-year-old tsunami layer (backwash) off southern Portugal (Reicherter et al., 2019). Sedimentary characteristics differ significantly from onshore deposit.

## Conclusions

"The present is the key to the past" is a guiding concept in the geological sciences and ten years after the 2011 Tohoku-oki tsunami, we present some examples of recent multidisciplinary tsunami deposit studies both onshore and offshore trying to underline the recent progress made in the Mediterranean area and in the Gulf of Cadiz, taking advantage of the lessons learned from this giant Japan event.

"The past is a key for the future", this is a guiding concept too that is at the basis of the hazard and risk evaluation that can be derived by the research of paleotsunamis signatures presented in this review. In fact, reconstructing the past history of tsunamis in the Mediterranean is a precious input to estimate the potential impact of future tsunamis in the region.

### *Onshore high-resolution studies: Direct Push technique*

The case study from western Greece showed that direct push (DP) sensing techniques applied to coastal sedimentary archives yield promising results and advantages. For instance, DP techniques help to identify and trace, on a centimeter scale, tsunami deposits both in their vertical and lateral dimensions, with a clear discrimination between tsunami sediments and other facies, and coupled with a prompt analytical *in situ* data visualization. Also, DP approaches allow for the reduction of the number of sediment cores and laboratory analyses needed. In this regard, the 2011 Tohoku-oki event enables the paleotsunami research community to re-evaluate the paleo-reconstruction of important tsunami criteria such as the inundation distance and more in general the areas prone to tsunami inundation. However, the observed 2011 mismatch between the water penetration distance and the sand dominated deposit left inland (about 70%) together with the difficult preservation of the inundation geochemical signature shed new light on the possible differences between "paleo" tsunami signatures and traces of "actual" large events, asking for a stronger effort in terms of multidisciplinary high-resolution studies to be applied for a better evaluation of the local tsunami hazard. This consideration supports the idea that researchers in search for paleotsunami deposits will likely observe a minimum value in terms of inundation distance and probably also in terms of recurrence at the site. This is coherent with the Jogan 869 CE tsunami lesson, known from both historical and field studies but not fully interpreted in its characteristics before the 2011 Tohoku-oki event.

### *Onshore high-resolution studies: combined XRF-X-CT approach*

Another significant contribution may be derived from the combined XRF-X-CT high resolution approach allowing the automatic calculation of grain-size metrics on two-dimensional vertical slice X-ray images along the core and the identification of scouring and detachment of mud clasts as well as formation of mud aggregates, mudlines and mud drapes on a millimeter scale, proved to be quite promising in

furnishing crucial information on the characterization of flow dynamics, and potentially adding an extra parameter to be used for the well-known problem of the differentiation between storm- and paleotsunami-related deposits, question that may affect also the Mediterranean region.

#### *Onshore high-resolution studies: new approaches on boulders*

The recent Mediterranean studies of boulders of exceptional size displaced tens of meters from the edge of the cliff and up to tens of meters above sea level on coastal platforms showed that a great effort has been done in order to obtain precise dimensional and spatial parameters of these blocks with high resolution 3D model of the terrain (using drones, DGPS and terrestrial Laser Scanner surveys) and a reliable age of the emplacement period (e.g. the use and integration of multiple approaches as radiocarbon dating and rate of growth of dissolution pans). This was combined with the application of recently available different codes to calculate the minimum height of the tsunami wave or storm wave capable of moving these boulders at the present position.

#### *The geo-archeological “new field” contribution*

An important effort for the compilation of paleotsunami deposits databases has been recently achieved in the North-East Atlantic and Mediterranean area. This region is known for being quite rich in historical data on tsunamis of the past, organized in relevant databases, but what it is still missing is the contribution from the studies, conducted mainly by archeologists, in numerous ancient coastal sites. In this sense the geo-archaeological approach to the tsunami research, is quite new and it clearly needs to be further developed favoring a strong collaboration between the geo- and archeo-research communities together with the integration of their different and often complementary approaches and techniques.

The (geo-)archaeological contribution in the Crete Island (Greece) case study is based on the recognition and study of tsunami deposits at the Minoan town of Palaikastro and resulted in the development of the ability to differentiate tsunami impact deposits in archaeological strata from “normally expected” site sediments. This result was obtained thanks to macroscale identification of predominant horizontal orientation of pottery sherds, stones and bones potentially related to the imbrication patterns due to a tsunami flow and to micromorphology studies of undisturbed *in situ* samples showing clear characteristics of multi-modal deposits, unexpected from commonly known archaeological site deposits formation processes.

#### *Offshore studies: multidisciplinary and multiproxy approaches*

Another important advancement from 2011 event is the observation of the occurrence of significant backwash flow likely capable of leaving traceable signatures in the offshore realm. This element stimulated new ideas about the approaches to be used offshore, particularly on the continental shelf areas while little research has been done so far in the deeper basins in order to discriminate between “paleoseismic turbidite” and “paleotsunami turbidite.”



As for the offshore research and the tsunami sedimentary traces in the geological record, we present some recent studies underlining the potential of the offshore sedimentological record as an archive of past tsunamis. In Israel, investigations on the shallow shelf environment discovered anomalous coarse layers exhibiting a peculiar content from shallower and deeper sediments, suggesting tsunami related multi-directional high-velocity transport. In Italy, the compared study of multiple cores, collected at different depths on the middle offshore environment, was accomplished thanks to a multidisciplinary approach (from X-ray imaging to micropaleontology and multiple dating methods) identifying several layers characterized by terrestrial run-off coupled with an input of organic matter and interpreted as tsunami related. Similar results were obtained on the southwestern Portuguese shelf (down to the shelf break water depth) by applying a multiproxy approach including analyses on compositional, geophysical and geochemical properties and multiple dating techniques looking for the sedimentological characteristics and paleoecological effects of the CE 1755 Lisbon tsunami and older backwash events.

Finally, it is important to notice that most if not all the paleotsunami studies presented share a common effort for multidisciplinary and multiproxy approaches. In fact, it is now considered from the whole community as essential to get stronger and more robust evidence for tsunami events of the past. At the same time the high-resolution investigations are providing unprecedented details on the grain size distribution and on thin sedimentary structures that can substantially help in the reconstruction of flow dynamics of the causative tsunami event. There are also two “new” fields of investigation that until now received limited attention: geoarchaeology and offshore studies that are presented here with their approaches. This highlights that in the near future we can expect new paleotsunami data from these new promising approaches. In conclusion all the case studies, new methodological aspects, and data presented on paleotsunami records may and should contribute to a robust and modern tsunami hazard assessment. This will also positively affect the production of highly valuable inundation scenarios as they are supported by different (modeling and field data) approaches.

## References

- Abe, H., Sugeno, Y., Chigema, A., 1990. Estimation of the height of the Sanriku Jogan 11 Earthquake-tsunami (A.D. 869) in the Sendai Plain. *Journal of the Seismological Society of Japan (Zishin)* 43, 513–525.
- Abe, T., Goto, K., Sugawara, D., 2012. Relationship between the maximum extent of tsunami sand and the inundation limit of the 2011 Tohoku-oki tsunami on the Sendai Plain, Japan. *Mar. Geol.* 282, 142-150.
- Álvarez-Gómez, J.A., Aniel-Quiroga, I., González, M. & Otero, L. 2011. Tsunami hazard at the Western Mediterranean Spanish coast from seismic sources. *Nat. Hazards Earth Syst. Sci.* 11, 227-240.
- Ambraseys, N. 2009. *Earthquakes in the Mediterranean and Middle East. A multidisciplinary study of seismicity up to 1900*, Cambridge University Press, Cambridge.
- Andrade, C., Freitas, M.C., Moreno, J., Craveiro, S.C., 2004. Stratigraphical evidence of Late Holocene barrier breaching and extreme storms in lagoonal sediments of Ria Formosa, Algarve, Portugal. *Mar. Geol.* 210, 339-362.
- Anzidei, M., Esposito, A., Benini, A., 2014. Evidence of active subsidence at Basiluzzo island (Aeolian islands, southern Italy) inferred from a Roman age wharf. *Quat. Int.* 332, 143-150, doi: 10.1016/j.quaint.2014.03.019.
- Arai, K., Naruse, H., Miura, R., Kawamura, K., Hino, R., Ito, Y., Inazu, D., Yokokawa, M., Izumi, N., Murayama, M., Kasaya, T., 2013. Tsunami-generated turbidity current of the 2011 Tohoku-Oki earthquake. *Geology* 41, 1195–1198. <http://dx.doi.org/10.1130/G34777.1>
- Atwater, B., 1987. Evidence for Great Holocene Earthquakes Along the Outer Coast of Washington State. *Science* 236, 942-944. DOI: 10.1126/science.236.4804.942
- Avallone, A., Briole, P., Agatza-Balodimou, A.M., Billiris, H., Charade, O., Mitsakaki, C., Nercessian, A., Papazissi, K., Paradissis, D., Veis, G., 2004. Analysis of eleven years of deformation measured by GPS in the Corinth Rift Laboratory area, *Comptes Rendus. Geoscience* 336, 301-311.
- Bahrenberg, G., Giese, E., Mevenkamp, N. & Nipper, J. 2008. *Statistische Methoden in der Geographie. Band 2: Multivariate Statistik*, Gebrüder Borntraeger Verlagsbuchhandlung, Berlin, Stuttgart.
- Barbano, M.S., Pirrotta, C., Gerardi, F., 2010. Large boulders along the south-eastern Ionian coast of Sicily: storm or tsunami deposits? *Mar. Geol.* 275, 140–154.
- Basili R., Kastelic V., Demircioglu M. B., Garcia Moreno D., Nemser E. S., Petricca P., Sboras S. P., Besana-Ostman G. M., Cabral J., Camelbeeck T., Caputo R., Danciu L., Domac H., Fonseca J., García-Mayordomo J., Giardini D., Glavatovic B., Gulen L., Ince Y., Pavlides S., Sesetyan K., Tarabusi G., Tiberti M. M., Utkucu M., Valensise G., Vanneste K., Vilanova S., Wössner J., 2013. The European Database of Seismogenic Faults (EDSF) compiled in the framework of the

Project SHARE. <http://diss.rm.ingv.it/share-edsf/>, doi: 10.6092/INGV.IT-SHARE-EDSF

- Biolchi, S., Furlani, S., Antonioli, F., Baldassini, N., Causon Deguara, J., Devoto, S., Di Stefano, A., Evans, J., Gambin, T., Gauci, R., Mastronuzzi, G., Monaco, C., Scicchitano, G., 2016. Boulder accumulations related to extreme wave events on the eastern coast of Malta. *Nat. Hazards Earth Syst. Sci.* 16, 737-756.
- Biolchi S., Furlani S., Devoto S., Scicchitano G., Korbarclvic T., Šepić V.J. 2019. The origin and dynamics of coastal boulders in a semi-enclosed shallow basin: A northern Adriatic case study. *Mar. Geol.* 411, 62-77.
- Bonaccorso, A., Calvari, S., Garfi, L., Lodato, L., Patanè, D., 2003. Dynamics of the December 2002 flank failure and tsunami at Stromboli volcano inferred by volcanological and geophysical observations. *Geophys. Res. Lett.* 30, 1941.
- Bruins, H.J., MacGillivray, J.A., Synolakis, C.E., Benjamini, C., Keller, J., Kisch, H.J., Klügel, A., Van der Plicht, J., 2008. Geoarchaeological tsunami deposits at Palaikastro (Crete) and the late Minoan IA eruption of Santorini. *J. Archaeol. Sci.* 35, 191–212.
- Bruins, H.J., Van der Plicht, J., MacGillivray, J.A., 2009. The Minoan Santorini eruption and tsunami deposits in Crete (Palaikastro): Geological, archaeological, 14C dating and Egyptian chronology. *Radiocarbon* 51, 397–411.
- Bruins, H.J., Van der Plicht, J., 2014. The Thera olive branch, Akrotiri (Thera) and Palaikastro (Crete): Comparing radiocarbon results of the Santorini eruption. *Antiquity* 88, 282–287.
- Carey, S., Morelli, D., Sigurdsson, H., Bronto, S., 2001. Tsunami deposits from major explosive eruptions: An example from the 1883 eruption of Krakatau. *Geology* 29, 347-350.
- Costa, C.N.L., Rita, M.M., 1992. Wave Climatology of the Portuguese Coast (Clima de agitação marítima na costa portuguesa). Technical Report Rel. PT-OM-03/92, Instituto Hidrográfico, Lisbon.
- Cox, R., O'Boyle, L., Cytrynbaum J., 2019. Imbricated Coastal Boulder Deposits are Formed by Storm Waves, and Can Preserve a Long-Term Storminess Record. *Sci. Rep.* 9, 10784. <https://doi.org/10.1038/s41598-019-47254-w>
- Chagué-Goff, C., 2010. Chemical signatures of palaeotsunamis: a forgotten proxy? *Mar. Geol.* 271, 67–71.
- Chagué-Goff, C., Andrex, A., Szczuciński, W., Goff, J., Nishimura, Y., 2012. Geochemical signatures up to the maximum inundation of the 2011 Tohoku-oki tsunami – implications for the 869 AD Jogan and other palaeotsunamis. *Mar. Geol.* 282, 65-77.
- Chagué-Goff, C., Szczuciński, W., Shinozaki, T. 2017. Applications of geochemistry in tsunami research: A review. *Earth Sci. Rev.* 165, 203- 244.
- Chini, M., Piscini, A., Cinti, F.R., Amici, S., Nappi, R., De Martini P.M., 2013. The 2011 Tohoku-Oki (Japan) tsunami inundation and

- liquefaction investigated by optical, thermal and SAR data, *IEE Geosci. Remote Sens. Lett* 10, 347-351.
- Costa, P.J.M., Costas, S., González-Villanueva, R., Oliveira, M.A, Roelvink, D., Andrade, C., Freitas, M.C., Cunha, P.P., Martins, A., Buylaert, J.-P. & Murray, A. 2016. How did the AD 1755 tsunami impact on sand barriers across the southern coast of Portugal? *Geomorphology* 268, 296-311. <https://doi.org/10.1016/j.geomorph.2016.06.019>.
- Cuven, S., Paris, R., Falvard, S., Miot-Noirault, E., Benbakkar, M., Schneider, J.L., Billy, I., 2013. High-resolution analysis of a tsunami deposit: case-study from the 1755 Lisbon tsunami in southwestern Spain. *Mar. Geol.* 337, 98-111.
- Dawson, A.G., Long, D., Smith, D.E., 1988. The Storegga slides: evidence from eastern Scotland for a possible tsunami. *Mar. Geol.* 82, 271-276. [https://doi.org/10.1016/0025-3227\(88\)90146-6](https://doi.org/10.1016/0025-3227(88)90146-6)
- Dawson, A.G., Stewart, I., 2007. Tsunami deposits in the geological record. *Sediment. Geol.* 200, 166-183.
- De Alteriis, G., Insinga, D.D., Morabito, S., Morra, V., Chiocci, F.L., Terrasi, F., Lubritto, C., Di Benedetto, C., Pazzanese, M., 2010. Age of submarine debris avalanches and tephrostratigraphy offshore Ischia Island, Tyrrhenian Sea, Italy. *Mar. Geol.* 278, 1-18.
- DeMets, C., Gordon, R.G., Argus, D.F., 2010. Geologically current plate motions. *Geophys J. Int.* 181, 1-80.
- De Martini, P.M., Barbano, M.S., Smedile, M.S., Geradi, F., Pantosti, D., Del Carlo, P., Pirrotta, C. 2010. A unique 4000 year long geological record of multiple tsunami inundations in the Augusta Bay (eastern Sicily, Italy). *Mar. Geol.* 276: 42-57.
- Dey, H., Goodman-Tchernov, B., Sharvit, J., 2014. Archaeological evidence for the tsunami of January 18, A.D. 749: A chapter in the history of early Islamic Qâysariyah (Caesarea Maritima). *J. Rom. Archaeol.* 27, 357-373. <https://doi.org/10.1017/S1047759414001287>.
- Dietrich, P., Leven, C., 2006. Direct Push-Technologies, in: Kirsch R. (Eds.), *Groundwater Geophysics*. Springer, Berlin, Heidelberg.
- Dominey-Howes, D., Cundy, A., Croudace, I., 2000. High energy marine flood deposits on Astypalaea Island, Greece: possible evidence for the AD 1956 southern Aegean tsunami. *Mar. Geol.* 163, 303-315.
- Dominey-Howes, D.T.M., Humphreys, G.S., Hesse, P.P., 2006. Tsunami and palaeotsunami depositional signatures and their potential value in understanding the late-Holocene tsunami record. *Holocene* 16, 1095-1107.
- Driessen, J., Macdonald, C.F., 1997. *The Troubled Island, Minoan Crete before and after the Santorini eruption*. Université de Liège, Aegaeum 17.
- Engel, M., May, S. M. 2012. Bonaire's boulder fields revisited: Evidence for Holocene tsunami impact on the Lee-ward Antilles. *Quat. Sci. Rev.* 54, 126-141.
- Faccenna, C., Becker, T., Auer, L., Billi, A., Boschi, L., Brun, J.P., Capitanio, F.A., Funicello, F., Horvath, F., Jolivet, L., Piromallo, C., Royden, L., Rossetti, F., Serpelloni, E., 2014. Mantle dynamics in the

- Mediterranean. *Rev. Geophys.*, 52, 283–332, doi:10.1002/2013RG000444.
- Fago P., Pignatelli C., Piscitelli A., Milella M., Venerito M., Sansò P., Mastronuzzi G., 2014. The WebGIS on Italian tsunamis: an useful reference tool. *Mar. Geol.*, 355, 369–376.
- Falvard, S., Paris, R., 2017. X-ray tomography of tsunami deposits: towards a new depositional model of tsunami deposits. *Sedimentology* 64, 453–477.
- Feldens, P., Schwarzer, K., Sakuna, D., Szczuciński, W., Sompongchaiyikul, P., 2012. Sediment distribution on the inner continental shelf off Khao Lak (Thailand) after the 2004 Indian Ocean Tsunami. *Earth Planets Space*, 64, 875–887.
- Finkler, C., Fischer, P., Kaika, K., Rigakou, D., Metallinou, G., Hadler, H., Vött, A., 2018a. Tracing the Alkinoos Harbor of ancient Kerkyra, Greece, and reconstructing its palaeotsunami history. *Geoarchaeology* 33, 24–42.
- Finkler, C., Baika, K., Rigakou, D., Metallinou, G., Fischer, P., Hadler, H., Emde, K. and Vött, A., 2018b. Geoarchaeological investigations of a prominent quay wall in ancient Corcyra – implications for harbour development, palaeoenvironmental changes and tectonic geomorphology of Corfu island (Ionian Islands, Greece). *Quat. Int.* 473, 91–111.
- Giachetti, T., Paris, R., Kelfoun, K., Ontowirjo, B., 2012. Tsunami hazard related to a flank collapse of Anak Krakatau volcano, Sunda Strait, Indonesia. *Geol. Soc. Lond. Spec. Publ.* 361, 79–89.
- Goff, J., Chagué-Goff, C., Nichol, S., Jaffe, B., Dominey-Howes, D., 2012. Progress in palaeotsunami research. *Sediment. Geol.*, 243–244, 70–88.
- Gómez-Pujol, L., Roig-Munar, F.X., 2013. Acumulaciones de grandes bloques en las crestas de los acantilados del sur de Menorca (Illes Balears): observaciones preliminares. *Geo-Temas*, 14, 71–74.
- Goodman-Tchernov, B.N., Dey, H.W., Reinhardt, E.G., McCoy, F., Mart, Y., 2009. Tsunami waves generated by the Santorini eruption reached Eastern Mediterranean shores. *Geology* 37, 943–946. <https://doi.org/10.1130/G25704A.1>
- Goodman-Tchernov, B.N., Austin Jr., J., 2015. Deterioration of Israel's Caesarea Maritima's Ancient Harbor Linked to Repeated Tsunami Events Identified in Geophysical Mapping of Offshore Stratigraphy. *J. Archaeol. Sci. Reports* 444–454.
- Goodman Tchernov, B., Katz, T., Shaked, Y., Qupty, N., Kanari, M., Niemi, T., Agnon, A., 2016. Offshore Evidence for an Undocumented Tsunami Event in the 'Low Risk' Gulf of Aqaba-Eilat, Northern Red Sea. *PLoS One* 11, e0145802. <https://doi.org/10.1371/journal.pone.0145802>
- Goto, K., Chagué-Goff, C., Fujino, S., Goff, J., Jaffe, B., Nishimura, Y., Richmond, B., Suguwara, D., Szczuciński, W., Tappin, D.R., Witter, R., Yulianto, E., 2011. New insights of tsunami hazard from the 2011 Tohoku-oki event. *Mar. Geol.* 290, 46–50.

- Gouhier, M., Paris, R., 2019. SO<sub>2</sub> and tephra emissions during the December 22, 2018 Anak Krakatau flank-collapse eruption. *Volcanica* 2, 91-103.
- Grezio A., Babeyko, A., Baptista, M.A., Behrens, J., Costa, A., Davies, G., Geist, E.L., Glimsdal, S., González, F.I., Griffin, J., Harbitz, C.B., LeVeque, R.J., Lorito, S., Lövholt, F., Omira, R., Mueller, C., Paris, R., Parsons, T., Polet, J., Power, W., Selva, J., Sörensen, M.B., Thio, H.K., 2017. Probabilistic Tsunami Hazard Analysis (PTHA): multiple sources and global applications. *Rev. Geophys.* 55, 1158-1198.
- Grilli, S., Tappin, D., Carey, S., Watt, S., Ward, S., Grilli, A., Engwell, S., Zhang, C., Kirby, J., Schambach, L. and Muin, M., 2019. Modelling of the tsunami from the December 22, 2018 lateral collapse of Anak Krakatau volcano in the Sunda Straits, Indonesia. *Sci. Rep.* 9, 11946.
- Guidoboni, E., Comastri, A., Traina, G., 1994. Catalogue of ancient earthquakes in the Mediterranean area up to the 10th century. Volume 1. ING-SGA, Bologna.
- Hoffmann, N., Master, D., Goodman-Tchernov, B., 2018. Possible tsunami inundation identified amongst 4-5th century BCE archaeological deposits at Tel Ashkelon, Israel. *Mar. Geol.* 396, 150–159. <https://doi.org/10.1016/j.margeo.2017.10.009>
- Holum, K.G., Stabler, J.A., Reinhardt, E.G., 2008. Caesarea reports and studies: excavations 1995-2007 within the Old City and the ancient harbor, in: *Caesarea Reports and Studies: Excavations 1995-2007*. Oxford.
- Ikehara, K., Irino, T., Usami, K., Jenkins, R., Omurad, A. and Ashi, J., 2014. Possible submarine tsunami deposits on the outer shelf of Sendai Bay, Japan resulting from the 2011 earthquake and tsunami off the Pacific coast of Tohoku. *Mar. Geol.* 358, 120–127.
- Inghilesi, R., Corsini, S., Guiducci, F. and Arseni, A., 2000. Statistical analysis of extreme waves on the Italian coasts from 1989 to 1999. *Boll. di Geofis. Teor. ed Appl.* 41, 315-337.
- Istituto Idrografico della Marina, 1982. *Il vento e lo stato del mare lungo le coste italiane e dell'Adriatico*, Istituto Idrografico della Marina, III.
- Iwai S., Goto K., Ishizawa T., 2019. A gigantic boulder transported by the 2011 Tohoku-oki tsunami. *Island Arc* 28, e12321. <https://doi.org/10.1111/iar.12321>
- IOC. 2011. ICG/NEAMTWS: Seventh Session Paris, France 23–25 November 2010. Intergovernmental Oceanographic Commission, Reports of Governing and Major Subsidiary Bodies, 45 p.
- Jagodziński, R., Sternal, B., Szczuciński, W., Chagué-Goff, C., Sugawara, D., 2012. Heavy minerals in the 2011 Tohoku-oki tsunami deposits- insights into sediment sources and hydrodynamics. *Sediment. Geol.* 282, 57-64.
- Kelletat, D., Scheffers, A., Scheffers, S., 2005. Tsunami-eine unterschätzte Naturgefahr? *Geowissenschaftliche Mitteilungen*, 5-13.



- Köhn, M., 1929. Korngrößenbestimmung mittels Pipettanalyse. *Tonindustrie-Zeitung* 55, 729-731.
- Kortekaas, S., Dawson, A.G., 2007. Distinguishing tsunami and storm deposits: an example from Martinhal, SW Portugal. *Sediment. Geol.* 200, 208-221. <https://doi.org/10.1016/j.sedgeo.2007.01.004>.
- Koster, B., Vött, A., Mathes-Schmidt, M., Reicherter, K., 2015. Geoscientific investigations in search of tsunami deposits in the environs of the Agoulinitza peatland, Kaiafas Lagoon and Kakovatos (Gulf of Kyparissia, western Peloponnese, Greece). *Z. Geomorphol., N.F., Supplementary Issue* 59/4, 125-156. [http://dx.doi.org/10.1127/zfg\\_suppl/2014/S-00192](http://dx.doi.org/10.1127/zfg_suppl/2014/S-00192)
- Landva, A.O., Korpijaakkoo, E.O., Pheeney, P.E., 1983. Geotechnical classification of peats and organic soils, in: Jarrett, O.M. (Ed.), *Testing of Peats and Organic Soil*, American Society for Testing and Materials, ASTM STP 20, pp. 37-51.
- Lario, J., Zazo, C., Goy, J.L., Silva, P.G., Bardají, T., Cabero, A., Dabrio, C.J., 2011. Holocene paleotsunami catalogue of SW Iberia. *Quat. Int.* 242, 196-200. <https://doi.org/10.1016/j.quaint.2011.01.036>
- Lario, J.; Spencer, C.; Bardaji, T., Marchante, A., 2017. Eventos de oleaje extremo en la costa del sureste peninsular: bloques y megabloques como indicadores de tsunamis o tormentas extremas. *Geo-Temas* 17, 227-230.
- Lunne, T., Robertson, P.K. & Powell, J.J.M., 2002. *Cone Penetration Testing in Geotechnical Practice*. Taylor and Francis Group, London, New York.
- MacGillivray, J.A., Sackett, L.H., Driessen, J., Smyth, D., 1987. Excavations at Palaikastro 1986. *The Annual of the British School of Archaeology at Athens* 82, 135-154 (Plates 19-24).
- Maeno, F., Imamura, F., 2011. Tsunami generation by a rapid entrance of pyroclastic flow into the sea during the 1883 Krakatau eruption, Indonesia. *J. Geophys. Res.* 116, B09205.
- Maramai, A., Graziani, L., Tinti, S., 2005a. Tsunamis in the Aeolian Islands Southern Italy: a review. *Mar. Geol.* 214, 11-21.
- Maramai, A., Graziani, L., Alessio, G., Burrato, P., Colini, L., Cucci, L., Nappi, R., Nardi, A., Vilaro, G., 2005b. Near- and far-field survey report of the 30 December 2002 Stromboli (Southern Italy) tsunami. *Mar. Geol.* 215, 93-106, [doi:10.1016/j.margeo.2004.11.009](https://doi.org/10.1016/j.margeo.2004.11.009)
- Martin-Prieto, J.A.; Roig-Munar, F.X.; Rodrigues-Perea, A., Gelabert, B. 2019. Nova troballa de blocs de tsunami a les costes rocoses de sa Punta de sa Miloca-Corral Fals (Sud de Menorca, illes Balears). *Nemus. Revista de l'Ateneu de Natura* 9, 7-14.
- McCoy, F.W., Heiken, G., 2000. Tsunami generated by the Late Bronze age eruption of Thera (Santorini), Greece. *Pure Appl. Geophys* 157, 1227-1256.
- McFadgen, B.G., Goff, J.R., 2007. Tsunamis in the New Zealand archaeological record. *Sed. Geol.* 200, 263-274.

- Marriner, N., Kaniewski, D., Morhange, C., Flaux, C., Giaime, M., Vacchi, M. & Goff, J., 2017. Tsunamis in the geological record: Making waves with a cautionary tale from the Mediterranean. *Sci Adv.* 3, 31700485, <http://dx.doi.org/10.1126/sciadv.1700485>.
- Masclé, J., Martin, L., 1990. Shallow structure and recent evolution of the Aegean Sea: a synthesis based on continuous reflection profiles. *Mar. Geol.* 94, 271-299.
- Mastronuzzi G., Sansò P., 2000. Boulders transport by catastrophic waves along the Ionian coast of Apulia (Southern Italy). *Mar. Geol.* 170, 93-103.
- Mastronuzzi G., Sansò P., 2004. Large Boulder Accumulations by Extreme Waves along the Adriatic Coast of southern Apulia (Italy). *Quat. Int.* 120, 173-184.
- Mastronuzzi, G., Pignatelli, C. & Sansò, P., 2006. Boulder Fields: A Valuable Morphological Indicator of Paleotsunami in the Mediterranean Sea. *Z. Geomorphol., NF Suppl.-Bd.* 146: 173-194.
- Mastronuzzi G., Sansò S., 2006. Coastal geomorphology and tsunami vulnerability. The case study of Apulia region (Italy). *Geogr. Fis. e Din. e Quat.* 29, 83-91.
- Mastronuzzi G., Pignatelli C., Sansò P., Selleri G., 2007. Boulder accumulations produced by the 20th February 1743 tsunami along the coast of southeastern Salento (Apulia region, Italy). *Mar. Geol.* 242, 191-205.
- Mastronuzzi G., 2010. Tsunami in Mediterranean sea. *Egyptian J. Environ. Change* 2, 1-9.
- Mastronuzzi G., Pignatelli C., 2012. The boulders berm of Punta Saguerra (Taranto, Italy): a morphological imprint of the Rossano Calabro tsunami of April 24, 1836? *Earth Planets Space*, 64, 829-842
- Mastronuzzi G., Brückner H., De Martini P.M., Regnaud H., 2013. Tsunami: from the open sea to the coastal zone and beyond, in: Mambretti S. (Ed.), *Tsunami: From Fundamentals to Damage Mitigation*. WIT Press, Southampton, pp. 1-36.
- Meyers, S.R., 2014. *Astrochron: a R Package for Astrochronology* Version 0.6.5.
- Milia, A., Raspini, A., Torrente. M.M., 2007. The dark nature of Somma-Vesuvius volcano: Evidence from the ~3.5 ka B.P. Avellino eruption. *Quat. Int.* 173-174, 57-66.
- Minoura, K., Imamura, F., Kuran, U., Nakamura, T., Papadopoulos, G.A., Takahashi, T., Yalçiner, A., 2000. Discovery of Minoan tsunami deposits. *Geology* 28, 59-62, [https://doi.org/10.1130/0091-7613\(2000\)028<0059:DOMTD>2.0.CO;2](https://doi.org/10.1130/0091-7613(2000)028<0059:DOMTD>2.0.CO;2).
- Missiaen, T., Verhegge, J., Heirman, K., Crombé, P., 2015. Potential of cone penetration testing for mapping deeply buried palaeolandscapes in the context of archaeological surveys in polders. *J. Archaeol. Sci.* 55, 147- 187.
- Mori, N., Takahashi, T., The 2011 Tohoku Earthquake Tsunami Joint Survey Group, 2012. Nationwide post event survey and analysis of

- the 2011 Tohoku earthquake tsunami. *Coast Eng. J.* 54, 1250001-1-1250001-27. <https://doi.org/10.1142/S0578563412500015>
- Morton, R.A., Gelfenbaum, G., Jaffe, B.E., 2007. Physical criteria for distinguishing sandy tsunami and storm deposits using modern examples. *Sed. Geol.* 200, 184–207.
- Nastos, P.T., Karavana Papadimoua, K., Matsangourasa I.T., 2018. Mediterranean tropical-like cyclones: Impacts and composite daily means and anomalies of synoptic patterns. *Atmos. Res.* 208, 156–166. <http://dx.doi.org/10.1016/j.atmosres.2017.10.023>
- Neev, D., Bakler, N., Emery, K.O., 1987. *Mediterranean Coasts of Israel and Sinai: Holocene Tectonism from Geology, Geophysics, and Archaeology.* Taylor and Francis, New York.
- Nicosia, C. and Stoops, G., 2017. *Archaeological Soil and Sediment Micromorphology,* Wiley, Blackwell, Oxford.
- Nomikou, P., Carey, S., Bell, K., Papanikolaou, D., Bejelou, K., Cantner, K., Sakellariou, D., Perros, I., 2014. Tsunami hazard risk of a future volcanic eruption of Kolumbo submarine volcano, NE of Santorini Caldera, Greece. *Nat. Hazards* 72, 1375-1390.
- Nomikou, P., Druitt, T.H., Hübscher, C., Mather, T.A., Paulatto, M., Kalnins, L.M., Kelfoun, K., Papanikolaou, D., Bejelou, K., Lampridou, D., Pyle, D.M., Carey, S., Watts, A.B., Weiss, B., Parks, M.M., 2016. Post-eruptive flooding of Santorini caldera and implications for tsunami generation. *Nat. Commun.* 7, 13332.
- Nott, J., 2003. Waves, coastal boulder deposits and the importance of the pretransport setting. *Earth Planet Sci Lett* 210, 269-276.
- Novikova, T., Papadopoulos, G.A., McCoy, F.W., 2011. Modelling of tsunami generated by the giant Late Bronze Age eruption of Thera, South Aegean, Greece. *Geophys. J. Int.* 186, 665-680.
- Obrocki, L., Vött, A., Wilken, D., Fischer, P., Willershäuser, T., Koster, B., Lang, F., Papanikolaou, I., Rabbel, W. & Reicherter, W., 2020. Tracing tsunami signatures of the ad 551 and ad 1303 tsunamis at the Gulf of Kyparissia (Peloponnese, Greece) using direct push in situ sensing techniques combined with geophysical studies. *Sedimentology* 67, 1274-1308. <https://doi.org/10.1111/sed.12555>.
- Okal, E.A., Synolakis, C.E., 2004. Source discriminants for near-field tsunamis, *Geophys. J. Int.* 158, 899–912.
- Osmond, D.A., Bullock, P., 1970. *Micromorphological Techniques and Applications.* Harpenden: Agricultural Research Council, Technical Monograph No. 2.
- Papadopoulos, G.A., Gràcia, E., Urgeles, R., Sallares, V., De Martini, P.M., Pantosti, D., González, M., Yalciner, A.C., Mascle, J., Sakellariou, D., Salamon, A., Tinti, S., Karastathis, V., Fokaefs, A., Camerlenghi, A., Novikova, T., Papageorgiou, A., 2014. Historical and pre-historical tsunamis in the Mediterranean and its connected seas: Geological signatures, generation mechanisms and coastal impacts. *Mar. Geol.* 354, 81–109. <https://doi.org/10.1016/j.margeo.2014.04.014>

- Pareschi, M.T., Boschi, E., Mazzarini, F., Favalli, M., 2006a. Large submarine landslide offshore Mt. Etna. *Geophys. Res. Lett.* 33, L13302.
- Pareschi, M.T., Favalli, M., Boschi, E., 2006b. Impact of the Minoan tsunami of Santorini: simulated scenarios in the eastern Mediterranean. *Geophys. Res. Lett.* 33, L18607.
- Paris, R., Fournier, J., Poizot, E., Etienne, S., Mortin, J., Lavigne, F., Wassmer, P., 2010. Boulder and fine sediment transport and deposition by the 2004 tsunami in Lhok Nga (western Banda Aceh, Sumatra, Indonesia): a coupled offshore-onshore model. *Mar. Geol.* 268, 43–54.
- Paris, R., Wassmer, P., Lavigne, F., Belousov, A., Belousova, M., Iskandarsyah, Y., Benbakkar, M., Ontowirjo, B., Mazzoni, N., 2014. Coupling eruption and tsunami records: the Krakatau 1883 case-study, Indonesia. *Bull. Volcanol.* 76, 814.
- Paris, R., 2015. Source mechanisms of volcanic tsunamis. *Philos. Trans. Royal Soc A* 373, 20140380.
- Paris, R., Ulvrova, M., Selva, J., Brizuela, B., Costa, A., Grezio, A., Lorito, S., Tonini, R., 2019. Probabilistic hazard analysis for tsunamis generated by subaqueous volcanic explosions in the Campi Flegrei caldera, Italy. *J. Volcanol. Geotherm. Res.* 379, 106-116.
- Paris, R., Falvard, S., Chague', C., Goff, J., Etienne, S., Doumalin, P., 2020. Sedimentary fabric characterised by X-ray tomography: a case study from tsunami deposits on the Marquesas Islands, French Polynesia. *Sedimentology* 67, 1207-1229.
- Petricca, P., Babeyko, A.Y., 2019. Tsunamigenic potential of crustal faults and subduction zones in the Mediterranean. *Sci. Rep.* 9, 4326. <https://doi.org/10.1038/s41598-019-40740-1>
- Pignatelli C., Sansò P., Mastronuzzi G., 2009. Evaluation of tsunami flooding using geomorphologic evidence. *Mar. Geol.* 260, 6-18.
- Pilarczyk, J.E., Horton, B.P., Witter, R.C., Vane, C.H., Chagué-Goff, C., Goff, J., 2012. Sedimentary and foraminiferal evidence of the 2011 Tōhoku-oki tsunami on the Sendai coastal plain, Japan. *Sed. Geol.* 282, 78-89.
- Piscitelli A., Milella M., Shah-Hosseini M., Morhange C., Mastronuzzi G., 2017. Numerical approach to the study of coastal boulders: the case of Martigues, Marseille, France. *Quat. Int.* 439, 52-64.
- Quarta G., Fago P., Calcagnile L., Cipriano G., D'Elia M., Moretti M., Scardino G., Valenzano E., Mastronuzzi G., 2019. <sup>14</sup>C age offset in the Mar Piccolo sea basin in Taranto (southern Italy) estimated on *Cerastoderma glaucum* (POIRET, 1789). *Radiocarbon* 61, 1387-1401. DOI:10.1017/RDC.2019.38
- Reicherter, K., Becker-Heidmann, P., 2009. Tsunami deposits in the western Mediterranean: remains of the 1522 Almería earthquake? *Geol. Soc. Lond. Spec. Publ.* 316, 217-235.
- Reicherter, K., Vött, A., Feist, L., Costa, P.J.M., Schwarzbauer, J., Schütttrumpf, H., Jens, H., Raeke, A., Huhn-Frehers, K., 2019. Lisbon 1755, Cruise No. M152/1, 02.11. - 14.11.2018, Funchal (Portugal) - Hamburg (Germany). Electronic report, METEOR-

- Berichte M152, Gutachterpanel Forschungsschiffe, Bonn, Germany, doi: 10.2312/cr\_m152
- Reimer, P.J., 2013. IntCal13 and Marine13 radiocarbon age calibration curves 0-50000 years Cal BP. *Radiocarbon* 55, 1869-1887.
- Reinhardt, E.G., Goodman, B.N., Boyce, J.I., Lopez, G., van Hengstum, P., Rink, W.J., Mart, Y., Raban, A., 2006. The tsunami of 13 December A.D. 115 and the destruction of Herod the Great's harbor at Caesarea Maritima, Israel. *Geology* 34, 1061. <https://doi.org/10.1130/G22780A.1>
- Rhodes, B., Tuttle, M., Horton, B., Doner, L., Kelsey, H., Nelson, A., Cisternas, M., 2006. Paleotsunami Research. *Eos Trans. Am. Geophys. Union* 87, 205-209.
- Robertson, P.K., 2016. Cone penetration test (CPT)-based soil behaviour type (SBT) classification system – an update. *Can. Geotech. J.* 53, 1910-1927.
- Robinson, A.G., 1997. Regional and petroleum geology of the Black Sea and surrounding region. *American Association of Petroleum Geology Memoir*, 68, p. 385.
- Röbke, B.R., Vött, A., 2017. The tsunami phenomenon. *Prog Oceanogr.* 159, 296-322.
- Roig-Munar, F.X., 2016. Blocs de tempesta i tsunami a les costes rocoses de les Illes Balears. Anàlisi geomorfològica i morfomètrica. Ph.D. Thesis, Departament de Geodinàmica i Geofísica. Universitat de Barcelona, Spain, 410 pp.
- Roig-Munar, F. X., Vilaplana, J. M., Rodríguez-Perea, A., Martín-Prieto, J. Á., & Gelabert, B. 2018a. Tsunamis boulders on the rocky shores of Minorca (Balearic Islands). *Nat. Hazards Earth Syst. Sci.*, 18, 1985-1998.
- Roig-Munar, F. X., Forner, E., Martín-Prieto, J. Á., Segura, J., Rodríguez-Perea, A., Gelabert, B., & Vilaplana, J. M. 2018b. Presència de blocs de tsunamis i tempestes a les costes rocoses de la serra d'Irta (el Baix Maestrat, País Valencià). *Nemus. Revista de l'Ateneu de Natura* 8, 7-21.
- Roig-Munar, F. X., Rodríguez-Perea, A., Vilaplana, J.M., Martín-Prieto, J. & Gelabert, B. 2019a. Tsunami boulders in Majorca Island (Balearic Islands, Spain). *Geomorphology* 334, 76-90.
- Roig-Munar, F.X.; Rodríguez-Perea, A.; Martín-Prieto, J.A.; Gelabert, B.; & Vilaplana, J.M. 2019b. Tsunami boulders on the rocky coasts of Ibiza and Formentera (Balearic Islands). *J. Mar. Sci. Eng.* 7, 327.
- Romagnoli, C., Kokelaar, P., Casalbore, D., Chiocci, F.L., 2009. Lateral collapses and active sedimentary processes on the northwestern flank of Stromboli volcano, Italy. *Mar. Geol.* 265, 101-119.
- Sakuna, D., Szczuciński, W., Feldens, P., Schwarzer, K., Khokiattiwong, S., 2012. Sedimentary deposits left by the 2004 Indian Ocean tsunami on the inner continental shelf offshore of Khao Lak, Andaman Sea (Thailand). *Earth Planets Space* 64, 931-943.
- Salamon, A., Rockwell, T., Guidoboni, E., Comastri, A., 2011. A critical evaluation of tsunami records reported for the Levant Coast from



- the second millennium BCE to the present. *Isr. J. Earth Sci.* 58, 327–354. <https://doi.org/10.1560/IJES.58.2-3.327> A
- Sawai, Y.Y., Namegaya, Y., Okamura, Y., Satake, K., Shishikura, M., 2012. Challenges of anticipating the 2011 Tokohu earthquake and tsunami using coastal geology. *Geophys. Res. Lett.* 39, L21309.
- Scardino G., Piscitelli A., Milella M., Sansò P., Mastronuzzi G., 2020. Tsunami fingerprints along the Mediterranean coasts. *Rend. Lincei Sci. Fis. Nat.* 31,319–335. DOI: 10.1007/s12210-020-00895-w
- Scheffers, A., Kelletat, D., 2003. Sedimentological and geomorphologic tsunami imprints worldwide – a review. *Earth Sci. Rev.* 63, 83–92.
- Scheffers, A., Kelletat, D., Vött, A., May, S., Scheffers, S., 2008. Late Holocene tsunami traces on the western and southern coastlines of the Peloponnesus (Greece). *Earth Planet. Sci. Lett.* 269, 271–279.
- Schiffer, M.B., 1987. *Formation Processes of the Archaeological Record*. Albuquerque: University of New Mexico Press.
- Schneider, J.L., Chagué-Goff, C., Bouchez, J.L., Goff, J., Sugawara, D., Goto, K., Jaffe, B., Richmond, B., 2014. Using Magnetic Fabric to Reconstruct the Dynamics of Tsunami Deposition on the Sendai Plain, Japan — The 2011 Tohoku-Oki Tsunami. *Mar. Geol.* 358, 89–106.
- Scicchitano, G., Pignatelli, C., Spampinato, C. R., Piscitelli, A., Milella, M., Monaco, C., Mastronuzzi, G., 2012. Terrestrial Laser Scanner techniques in the assessment of tsunami impact on the Maddalena peninsula (South-eastern Sicily, Italy), *Earth Planets Space*, 64, 889–903.
- Seike, K., Kobayashi, G., Kogure, K., 2017. Post-depositional alteration of shallow-marine tsunami-induced sand layers: A comparison of recent and ancient tsunami deposits, Onagawa Bay, northeastern Japan. *Island Arc*. e12174. <https://doi:10.1111/iar.12174>.
- Selva, J., Tonini, R., Molinari, I., Tiberti, M. M., Romano, F., Grezio, A., Melini, D., Piatanesi, A., Basili, R., Lorito, S., 2016. Quantification of source uncertainties in Seismic Probabilistic Tsunami Hazard Analysis (SPTHA), *Geophys. J. Int.* 205, 1780–1803.
- Shah-Hosseini M., Morhange C., De Marco A., Wante J., Anthony E. J., Sabatier F., Mastronuzzi G., Pignatelli C., Piscitelli A., 2013. Coastal boulders in Martigues, French Mediterranean: evidence for extreme storm waves during the Little Ice Age. *Zeitschrift für Geomorphol.* 57, 181–199.
- Smedile, A., De Martini, P.M., Pantosti, D., Bellucci, L., Del Carlo, P., Gasperini, L., Pirrotta, C., Polonia, A., Boschi, E., 2011. Possible tsunami signatures from an integrated study in the Augusta Bay offshore (Eastern Sicily-Italy). *Mar. Geol.* 281, 1–13. <https://doi.org/10.1016/j.margeo.2011.01.002>
- Smedile, A., De Martini, P.M., Pantosti, D., 2012. Combining inland and offshore paleotsunamis evidence: The Augusta Bay (eastern Sicily, Italy) case study. *Nat. Hazards Earth Syst. Sci.* 12, 2557–2567. <https://doi.org/10.5194/nhess-12-2557-2012>

- Smedile, A., Molisso, F., Chagué, C., Iorio, M., De Martini, P.M., Pinzi, S., Collins, P., Sagnotti, L. and Pantosti, D., 2020. New coring study in Augusta Bay expands understanding of offshore tsunami deposits (Eastern Sicily, Italy). *Sedimentology* 67, 1553-1576. doi: 10.1111/sed.12581.
- Soloviev, S.L., Solovieva, O.N., Go, C.N., Kim, K.S., Shchetnikov, N.A., 2000. *Tsunamis in the Mediterranean Sea 2000 B.C.–2000 A.D. Advances in Natural and Technological Hazards Research*, Kluwer Academic Publishers, London.
- Sugawara, D, Goto, K., Jaffe, B.E., 2013. Numerical models of tsunami sediment transport - Current understanding and future directions. *Mar. Geol.* 352, 295-320. <http://dx.doi.org/10.1016/j.margeo.2014.02.007>
- Switzer, A.D., Burston, J.M., 2010. Competing mechanisms for boulder deposition on the southeast Australian coast. *Geomorphology* 114,42-54. doi:10.1016/j.geomorph.2009.02.009
- Szczuciński, W., Niedzielski, P., Kozak, L., Frankowski, M., Ziola, A., Lorenc, S., 2007. Effects of rainy season on mobilization of contaminants from tsunami deposits left in a coastal zone of Thailand by the 26 December 2004 tsunami. *Environ. Geol.* 53, 253–264.
- Szczuciński, W., 2012. The post-depositional changes of the onshore 2004 tsunami deposits on the Andaman Sea coast of Thailand. *Nat. Hazards* 60, 115–133. <https://doi.org/10.1007/s11069-011-9956-8>
- Szczuciński, W., Kokociński, M., Rzeszewski, M., Chagué-Goff, C., Cachão, M., Goto, K., Sugawara, D., 2012. Sediment sources and sedimentation processes of 2011 Tohoku-oki tsunami deposits on the Sendai Plain, Japan — Insights from diatoms, nannoliths and grain size distribution. *Sed. Geol.* 282, 40-56.
- Takabatake, T., Shibayama, T., Esteban, M., Achiari, H., Nurisman, N., Gelfi, M., Tarigan, T.A., Kencana, E.R., Fauzi, M.A.R., Panalaran, S., Harnantyari, A.S., Kyaw, T.O., 2019. Field survey and evacuation behavior during the 2018 Sunda Strait tsunami. *Coast Eng J.* 61, 423-443, DOI: 10.1080/21664250.2019.1647963
- Tamura, T., Sawai, Y., Ikehara, K., Nakashima, R., Hara, J., Kanai, Y., 2015. Shallow marine deposits associated with the 2011 Tohoku-oki tsunami in Sendai Bay, Japan. *J Quat. Sci.*, 30, 293-297.
- Tanigawa, K., Sawai, Y., Namegaya, Y., 2016. Diatom assemblages within tsunami deposits from the 2011 Tohoku-oki earthquake along the Misawa coast, Aomori Prefecture, northern Japan. *Mar.Geol.* 396, 6-15.
- Tinti, S., Maramai, A., Armigliato, A., Graziani, L., Manucci, A., Pagnoni, G., Zaniboni, F., 2006. Observations of physical effects from tsunamis of December 30, 2002 at Stromboli volcano, southern Italy. *Bull. Volcanol.* 68, 450–461, <https://doi.org/10.1007/s00445-005-0021-x>
- Tyuleneva, N., Braun, Y., Katz, T., Suchkov, I., Goodman-Tchernov, B., 2018. A new chalcolithic-era tsunami event identified in the

- offshore sedimentary record of Jisr al-Zarka (Israel). *Mar. Geol.* 396, 67–78  
<https://doi.org/https://doi.org/10.1016/j.margeo.2017.07.008>
- Ulutaş, E., 2020. The May 11 Paphos, Cyprus, earthquake: implications for stress regime and tsunami modelling for the Eastern Mediterranean shorelines. *Arab. J. Geosci.* 13, 970. <https://doi.org/10.1007/s12517-020-05943-1>
- Urgeles, R., Camerlenghi, A., 2013. Submarine landslides of the Mediterranean Sea: Trigger mechanisms, dynamics, and frequency-magnitude distribution, *J. Geophys. Res. Earth Surf.* 118, 2600–2618. doi:10.1002/2013JF002720
- Ulvrova, M., Paris, R., Nomikou, P., Kelfoun, K., Leibbrandt, S., Tappin, D.R., McCoy, F.W., 2016. Source of the tsunami generated by the 1650 AD eruption of Kolumbo submarine volcano (Aegean Sea, Greece). *J. Volcanol. Geotherm. Res.* 321, 125-139.
- Usami, K., Ikehara, K., Jenkins, R.G., Ashi, J., 2017. Benthic foraminiferal evidence of deep-sea sediment transport by the 2011 Tohoku-oki earthquake and tsunami, *Mar. Geol.* 384, 214-224, <https://doi.org/10.1016/j.margeo.2016.04.001>.
- Venables, W.N., Ripley, B.D., 2002. *Modern Applied Statistics with S*. Springer, New York.
- Vött, A., Brückner, H., May, M., Lang, F., Herd, R., Brockmüller, S., 2008. Strong tsunami impact on the Bay of Aghios Nikolaos and its environs (NW Greece) during Classical-Hellenistic times. *Quat. Int.* 181, 105 -122, <https://doi.org/10.1016/j.quaint.2007.02.017>.
- Vött, A., Lang, F., Brückner, H., Gaki-Papanastassiou, K., Maroukian, H., Papanastassiou, D., Giannikos, A., Hadler, H., Handl, M., Ntageretzi, K., Willershäuser, T., Zander, A., 2011. Sedimentological and geoarchaeological evidence of multiple tsunamigenic imprint on the Bay of Palairos-Pogonia (Akarnania, NW Greece). *Quat. Int.* 242, 213-239.
- Vött, A., Hadler, H., Koster, B., Mathes-Schmidt, M., Röbbke, B.R., Willershäuser, T., Reicherter, K., 2018. Returning to the facts: Response to the refusal of tsunami traces in the ancient harbour of Lechaion (Gulf of Corinth, Greece) by ‘non-catastrophists’ – Reaffirmed evidence of harbour destruction by historical earthquakes and tsunamis in AD 69-79 and the 6th cent. AD and a preceding pre-historical event in the early 8th cent. BC. *Zeitschrift für Geomorphol.* N.F. 61, 275-302. <http://dx.doi.org/10.1127/zfg/2018/0519>
- Vött, A., Bruins, H.J., Gawehn, M., Goodman-Tchernov, B.N., De Martini, P.M., Kelletat, D., Mastronuzzi, G., Reicherter, K., Röbbke, B.R., Scheffers, A., Willershäuser, T., Avramidis, P., Bellanova, P., Costa, P.J.M., Finkler, C., Hadler, H., Koster, B., Lario, J., Reinhardt, E., Mathes-Schmidt, M., Ntageretzi, K., Pantosti, D., Papanikolaou, I., Sansò, P., Scicchitano, G., Smedile, A. & Szczuciński, W., 2019b. Publicity waves based on manipulated geoscientific data suggesting climatic trigger for majority of tsunami findings in the Mediterranean – Response to ‘Tsunamis in the geological record:

- Making waves with a cautionary tale from the Mediterranean' by Marriner et al. (2017). *Zeitschrift für Geomorphol. N.F., Supplementary Issue*. doi: 10.1127/zfg\_suppl/2018/0547
- Vött, A., Willershäuser, T., Rübke, B.R., Obrocki, L., Fischer, P., Hadler, H., Emde, K., Eder, B., Gehrke, H.J. & Lang, F. (2019a): Major flood events recorded in the Holocene sedimentary sequence of the uplifted Ladiko and Makrisia basins near ancient Olympia (western Peloponnese, Greece). *Zeitschrift für Geomorphol.* 62, 1-53. DOI: 10.1127/zfg\_suppl/2018/0499.
- Wallis, S.R., Fujiwara, O., Goto, K., 2018. Geological studies in tsunami research since the 2011 Tohoku earthquake. *Geol. Soc. Spec. Publ.* 456, 39-53.
- Weiss, R., Bahlburg, H., 2006. A Note on the Preservation of Offshore Tsunami Deposits. *J. Sediment. Res.* 76, 1267-1273. <https://doi.org/10.2110/jsr.2006.110>
- Werner, V., Baika, K., Fischer, P., Hadler, H., Obrocki, L., Willershäuser, T., Tzigounaki, A., Tsigkou, A., Reicherter, K., Papanikolaou, I., Emde, K. & Vött, A., 2018a. The sedimentary and geomorphological imprint of the AD 365 tsunami on the coasts of southwestern Crete (Greece) – Examples from Sougia and Palaiochora. – *Quat. Int.* 473, 66-90.
- Werner, V., Baika, K., Tzigounaki, A., Reicherter, K., Papanikolaou, I., Emde, K., Fischer, P. & A. Vött, 2018b. Mid-Holocene tectonic geomorphology of northern Crete deduced from a coastal sedimentary archive near Rethymnon and a Late Bronze Age Santorini tsunamite candidate. *Geomorphology* 326, 167-189. <http://doi.org.com/10.1016/j.geomorph.2018.09.017>
- Yaltirak, C., 2002. Tectonic evolution of the Marmara Sea and its surroundings, *Mar. Geol.* 190, 493-529, DOI: 10.1016/S0025-3227(02)00360-2.
- Yoshikawa, S., Kanamatsu, T., Goto, K., Sakamoto, I., Yagi, M., Fujimaki, M., Imura, R., Nemoto, K., Sakaguchi, H., 2015. Evidence for erosion and deposition by the 2011 Tohoku-oki tsunami on the nearshore shelf of Sendai Bay. *Japan. Geo Mar. Lett.* 35, 315-328. DOI 10.1007/s00367-015-0409-3.