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2	Melting behavior of SiO₂ up to 120 GPa
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12	ABSTRACT
13	The structure of liquid silicates is commonly described as a statistical mixture of various atomic
14	entities with relative abundances that can vary with pressure, temperature and composition.
15	Unfortunately, this view remains largely theoretical due to scarce experimental reports on the silicate
16	melt structure, in particular under pressure. We performed X-ray diffraction of the SiO2 end-member
17	to probe the melting curve up to \sim 120 GPa and 7000 K, and the melt structure up to \sim 80 GPa. We
18	confirm the steep increase of the melting curve above ~14 GPa when stishovite becomes stable over
19	coesite in subsolidus conditions, with a slope of about 80 K/GPa. Then, around 45 GPa and 5400 K,
20	the melting curve flattens significantly, an effect most likely reflecting the densification of the SiO_2
21	melt structure. The signal of diffuse X-ray scattering is compatible with a change of the Si
22	coordination number from 4 to 6 along the melting curve, in agreement with previous works reporting
23	a similar evolution during the cold compression of SiO2-glass. Because of the limited pressure range
24	(within 10 to 20 GPa) in which the melting curve changes its slope, we speculate a difficult
25	coexistence of tetrahedral SiO_4 and octahedral SiO_6 units in SiO_2 melt at high pressures.
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28	KEYWORDS
29	SiO ₂ -silica, melting diagram, melt structure, lower mantle pressures
30	
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I. INTRODUCTION

40 Although SiO_2 is rarely present as a pure phase in mantle rocks, it is a dominant component of 41 terrestrial planets. Because SiO_2 is highly refractory, its melting curve significantly affects the melting diagram of silicates present in the crust and the mantle. At subsolidus conditions, SiO₂ undergoes a 42 major transition from coesite to stishovite, at a pressure of ~14 GPa for a temperature of 3000 K 43 44 (Zhang et al. 1993). The transition involves a change of the Si coordination number from tetrahedral "SiO₄" to octahedral "SiO₆" units, accompanied with a density change of \sim 30% (Akaogi et al. 2011). 45 46 This transition occurs in all silicates with a transition pressure that depends on composition. The 47 structural changes are major and induce a divergence of physical and chemical properties between the 48 shallower and deeper parts of planetary mantles.

The change of the Si coordination number is also expected to occur in silicate melts (Sanloup et al. 49 2013). For SiO_2 , the experimental investigation on the melt is strongly hampered by a steep increase of 50 51 the melting curve at pressures higher than the coesite to stishovite transition: The melting temperature 52 was reported to increase from 3000 to 3900 K between ~14 GPa and ~22 GPa (Shen and Lazor 1995). Such a steep Clapeyron slope $(dT/dP=\Delta V/\Delta S)$ denotes a large volume of fusion between stishovite and 53 54 a low density melt (LDM) containing Si in 4-fold coordination. In contrast, the Clapevron slope of the 55 melting curve is almost flat below ~14 GPa, which implies a negligible density contrast between 56 coesite and LDM (Zhang et al. 1993). The brutal change of slope of the melting curve at ~14 GPa shows that LDM is maintained at pressures above the Si coordination change in the solid state. 57

58 The pressure-induced Si coordination change in amorphous SiO_2 was investigated experimentally at room temperature (Benmore et al. 2010; Lin et al. 2007; Murakami and Bass 2010; Sato and 59 Funamori 2010) and theoretically (Meade et al. 1992; Stixrude and Karki 2005; Takada et al. 2016; 60 Usui and Tsuchiya 2010). In both cases, the transformation to an octahedral-based SiO_2 structure is 61 clearly demonstrated. At high temperatures, the formation of a high-density melt (HDM), with a 62 63 similar structure to the high-density glass potentially, would induce a smaller density contrast between stishovite and HDM, compared to stishovite and LDM. This would induce the flattening of the 64 Clapeyron slope of the melting curve with increasing pressure. This line of reasoning agrees with the 65 report of a moderate increase of the melting temperature from 5000 K to 5800 K with increasing 66 pressure from 40 to 120 GPa (Usui and Tsuchiya 2010). Still, large inconsistencies remain. For 67 68 example, the melting temperature was reported from 4400 K (Luo et al. 2002) to 5250 K (Usui and Tsuchiya 2010) at a pressures around 50 GPa. In addition, neither the change in Clapeyron slope of the 69 70 SiO₂ melting curve at increasing pressure, nor the change in Si coordination number in the melt have 71 been properly documented.

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II. MATERIALS AND METHODS

75 2.1 Experimental design

We first refined the melting curve of SiO_2 up to ~120 GPa using the laser-heated diamond anvil 76 77 cell (LH-DAC) coupled with in situ X-ray diffraction. Samples consisted of small shards of pure SiO₂ 78 glass, typically of 20-50 µm across and 10-15 µm thick. The glass was loaded in the pressure chamber between two KCl layers of similar thickness. The use of a KCl makes the X-ray analysis more 79 difficult, due to intense KCl peaks; however, a critical advantage is to minimize the axial thermal 80 gradients within the sample. In addition, soft KCl layers favor homogenous sample pressure at high 81 temperature. The amount of water trapped in the pressure chamber was minimized by performing all 82 sample-loading operations under N₂-flux in a glove bag. Samples were heated alternatively using a 83 84 CO₂ laser on one side or two fiber-lasers on both sides of the DAC. We found that SiO₂ absorbs 85 relatively well the 1 µm radiation of fiber-lasers after the sample is already heated to high temperatures using the CO_2 . In all cases, laser spots were more than 20 μ m in diameter. 86

87 In a second series of experiments, we probed the structure of SiO_2 melt at high pressures. We did not use a pressure medium in this case to prevent artifacts that could arise from its diffuse scattering, if 88 the pressure medium eventually melts. Yet, a small grain of KCl was loaded in the gasket hole for the 89 measurement of the nominal pressure. The sample geometry consisted in two thin disks of SiO_2 quartz 90 91 single crystal in direct contact with the diamonds, between which a little bit of Si-metal powder was 92 distributed. The role of Si-metal is to couple with fiber-lasers in order to initiate the sample heating before the annealed SiO_2 could eventually absorb the laser radiation by itself. We preferred the 93 94 addition of Si to Pt powder that would have also provided an internal pressure calibrant, because Pt 95 melts at lower temperature than SiO_2 at high pressures; its signal of diffuse scattering would have hampered the study of SiO₂-melt. The presence of a minor excess of Si in SiO₂ melt should not affect 96 97 significantly the melt structure, nor the mechanism of the pressure-induced Si coordination change. 98 Upon laser heating, SiO₂ eventually melts at mid-distance between the two diamond culets. Then, 99 further increase of the laser power induces a thicker layer of molten SiO₂. In such a sample experiencing large axial temperature gradient, some SiO₂ always remain solid at the contact with cold 100 101 diamonds. More importantly, the temperature gradient within the molten part of the sample should be 102 minor, due to turbulent convection in the high-pressure melt.

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104 **2.2 Determination of pressure and temperature in the laser heated diamond anvil cell**

We used the equation of state of KCl (Dewaele et al. 2012) to determine the nominal pressure at 300K before and after laser heating. Then, the sample pressure at high temperature was determined using two different cross-calibrated methods: (i) the P-V-T equation of state of stishovite (Wang et al. 2012) and (ii) the nominal pressure corrected by $X_{Th}\alpha KT$ to account for the increase of pressure in the laser spot (X_{Th} is an adjustable parameter, α and K are thermal expansion and bulk modulus, respectively, of coesite (Akaogi et al. 2011) at nominal pressures below 14 GPa and stishovite (Wang et al. 2012) for pressures above). We used samples loaded in KCl and heated below the melting curve to calibrate the value of X_{Th} . For this sample geometry, the two methods give pressures within +/- 5 GPa for X_{Th} values increasing linearly from 35% to 90%, when temperature is increased from 4000 to 7000 K, respectively. At the melting temperature, the second method is preferred, because the remaining fraction of stishovite grains may encounter a temperature lower than the SiO₂-melt, because they are located closer to the diamonds. The second method was also used to determine the pressure for DAC loadings without KCl pressure medium.

Temperatures up to 7000 K were determined based on spectroradiometric measurements (Fig. 1). 118 To make these measurements accurate, a critical parameter of the optical set-up is the use of reflective 119 120 Schwarzfield-type telescopes and the absence of any refractive device along the optical path, in order 121 to prevent chromatic aberration (Schultz et al. 2005). Temperature measurements are performed with a 122 precision of ± 100 K and a reproducibility better than ± 30 K. When not using KCl insulating layers in the DAC, the temperature uncertainty could be twice larger, due to steeper axial temperature gradient. 123 We note that the measured temperature corresponds to the hottest part of the sample, where thermal 124 125 emission is the most intense.

The sample behavior was monitored *in situ* using an X-ray beam much smaller than the hot-spot on the sample (see below). The fluorescence (in the visible range of wavelength) of the irradiated KCl pressure medium, or irradiated Re-gasket, was used to adjust the spectrometer position, in order to perform the temperature measurement at the position of the X-ray beam. The spectrometer alignments were carefully checked before and after each melting experiments.

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132 **2.3 X-ray diffraction measurements**

At the ID27 beamline of the ESRF (Grenoble, France), we used a monochromatic X-ray beam of 133 134 0.3738 Å wavelength focused by two Kirkpatrick-Baez mirrors onto a spot of less than $2x2 \ \mu m^2$. 135 Diffraction patterns were recorded using a MAR-160 CCD detector and integrated using DIOPTAS (Prescher and Prakapenka 2015). Acquisition times extended to 10 sec and 300 sec for solid and 136 molten SiO₂ samples, respectively. For experiments dedicated to the determination of the melting 137 curve, volumes and relative abundances of SiO₂ and KCl were refined using the XRDUA code (De 138 139 Nolf et al. 2014). For experiments dedicated to probe the structure of SiO₂-melt, we inserted Soller slits between the DAC and the CCD detector. These slits filter out a major fraction of the Compton 140 141 diffusion coming from diamonds (Weck et al. 2013), which results in a significant improvement of the signal of X-ray diffuse scattering coming from the melt. We first recorded the X-ray diffuse scattering 142 of SiO₂ glass under compression up to \sim 50 GPa, before we performed measurements at high 143 temperatures in SiO₂ melt up to 80 GPa. In both cases, we also recorded the diffuse background 144 coming from diamond anvils (Compton diffusion). For the melt, small diffraction peaks of not-molten 145 stishovite were eventually visible in the Q region between 1 and 8 Å⁻¹. The peaks were carefully fitted 146 by Gaussian functions and subtracted. Then, after subtraction of the diamond Compton diffusion, the 147 148 scattering signal coming from the sample was normalized to obtain the scattering factor S(Q). Finally,

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III. RESULTS AND DISCUSSION

the wavelength position of the first sharp diffraction peak (FSDP) was extracted from the S(Q) using a

154 **3.1** A steep SiO₂ melting curve up to 45 GPa

Gaussian function coupled with a polynomial background.

We first refine the melting curve of SiO_2 up to ~120 GPa using the LH-DAC coupled with *in situ* 155 X-ray diffraction. With increasing temperature, the glass starting material first crystallizes in coesite 156 below 14 GPa, stishovite or in the CaCl₂-form above ~60 GPa (Andrault et al. 1998). The presence of 157 158 thin and continuous diffraction rings are evidences for fine crystalline SiO₂ powders under a moderate temperature gradient within the X-ray beam (Fig. 2A). At a sample temperature interpreted hereafter 159 as the melting point of SiO₂, several changes occur simultaneously (such as reported in (Andrault et al. 160 2014)): (i) Additional increase of the laser power does not induce an increase of the sample 161 temperature (discussed in (Geballe and Jeanloz 2012)); (ii) texture of diffraction rings becomes 162 163 spottier; (iii) position of diffraction spots on the image plate changes rapidly at constant temperature; 164 (iv) integrated intensity of the SiO₂ diffraction peaks decreases severely relative to the KCl pressure 165 medium (Fig. 2); (v) upon quenching, intense stishovite peaks reappear, which certifies the presence of non-crystalline (i.e. molten) SiO_2 within the X-ray spot at high temperature. The reason why all 166 diffraction peaks of SiO_2 do not disappear abruptly at the melting temperature is the unavoidable axial 167 temperature gradient between the two diamonds. Also, the SiO₂ melt may absorb the laser radiation 168 less than solid SiO₂, resulting in a relatively lower heating efficiency when the sample is partially 169 170 molten. This can explain the temperature plateau despite an increase of the laser power.

171 Our measurements confirm the steep increase of the SiO_2 melting temperature above ~14 GPa (Fig. 3, Table 1), in agreement with previous experimental (Shen and Lazor 1995) and theoretical 172 (Belonoshko and Dubrovinsky 1995; Usui and Tsuchiya 2010) studies. Within experimental 173 uncertainties, the melting curve appears linear between ~15 and ~45 GPa. The Clapeyron slope of ~80 174 K/GPa remains ~4 times smaller than that reported at ~300 K/GPa for the coesite to stishovite 175 transition (Akaogi et al. 2011). Basic thermodynamic relations apply around the triple point where 176 177 coesite, stishovite and the LDM of SiO_2 coexist (Table 2). As a consequence, we estimate the volume of fusion of stishovite $(2x\Delta V_{\text{Melting}}/(V_{\text{LDM}}+V_{\text{Sti}}))$ at ~28% from the combination of the volume change 178 179 of ~28% at the subsolidus transition (Akaogi et al. 2011) and a negligible volume of fusion of coesite derived from the quasi-horizontal melting line between 9 to 14 GPa (Zhang et al. 1993). On the other 180 hand, the entropy of melting of stishovite (Δ Sm) is estimated to ~60 J/molK, a value 4 times larger 181 182 than the entropy change at the coesite-stishovite transition (Akaogi et al. 2011). 183

184 **3.2** A relatively flat SiO₂ melting curve from 45 to 90 GPa

185 The response of our samples to laser heating was different below and above a pressure of ~45 GPa. Between ~15 and ~45 GPa, after the onset of sample melting between ~3000 and ~5400 K (Fig. 3), 186 further increase of the laser power always yield to the same pressure-temperature conditions of ~45 187 188 GPa and ~5400 K. This behavior could result from the very large volume of fusion (~28%, see above). Increasing the amount of melt by LH within a nearly constant sample volume in the DAC would 189 logically result in a significant increase of the sample pressure. Above ~45 GPa, this effect is not 190 observed anymore; full sample melting could be achieved within the X-ray spot at each given 191 192 experimental pressure.

- 193 Between ~45 to ~90 GPa, the SiO₂ melting curve presents a Clapeyron slope of ~13 K/GPa, which is ~6 times lower than that observed at lower pressures (Fig. 3). The difference cannot result from 194 195 subsolidus properties, because the stability field of stishovite extends continuously between ~15 and ~90 GPa below the melting line. Instead, it can be related to the evolution of the structure of the SiO_2 196 melt. A smaller Clapeyron slope (dT/dP) above 45 GPa denotes either a smaller volume of fusion 197 (ΔVm) and/or a higher entropy of fusion (ΔSm). A higher ΔSm above 45 GPa is improbable, because 198 199 we expect large structural similarities between stishovite and HDM, as they both contain most of the 200 Si in a 6-fold coordination (Benmore et al. 2010; Lin et al. 2007; Meade et al. 1992; San et al. 2016; Sato and Funamori 2010). Instead, their structural similarities suggest a lower Δ Sm. Therefore, the 201 flattening of the melting slope above ~45 GPa is likely to result from a decrease of the volume of 202 203 fusion due to the presence of a HDM of SiO₂ above the melting curve.
- Now, if we consider a fictive triple point between LDM, HDM and stishovite at ~45 GPa and 5400 204 K, we can use Clapeyron relations ($\Sigma \Delta V=0$, $\Sigma \Delta S=0$ around triple points, and $dT/dP=\Delta V/\Delta S$) to 205 206 retrieve the thermodynamical parameters of these three phases. Assuming a negligible entropy difference between LDM and HDM ($\Delta S_{LDM-HDM}$), the stishovite volume of fusion (ΔVm) is estimated 207 208 to ~5.3% above ~45 GPa (see details of the calculation below Table 2). Δ Vm becomes 6.6 % or 4.0 % 209 for hypothetical ΔS_{LDM} -HDM values of e.g. -15 or 15 J/mol.K, respectively. Such range of ΔVm values is comparable to the density contrast (reported at ~ 3 %) between stishovite and a glass compressed 210 above 45 GPa at room temperature (Petitgirard et al. 2017). Then, we can estimate a volume difference 211 of ~25% between LDM and HDM melts. This value is marginally affected (+/- 2%) by changing 212 $\Delta S_{LDM-HDM}$ in a range of possible values. Such volume change is comparable to the 28% reported 213 between coesite and stishovite (Akaogi et al. 2011). We finally estimate the SiO₂ LDM and HDM 214 densities at ~3.26 10³ and ~4.17 10³ kg/m³, respectively, at 45 GPa and 5400 K (Table 2). Such LDM 215 density is consistent with a previous theoretical report considering less than 10% Si in 6-fold 216 217 coordination in the melt (Takada et al. 2016).
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219 **3.3 Melting of the CaCl₂-form of SiO₂ above 90 GPa**

Above ~90 GPa, the slope of the SiO₂ melting curve increases slightly to ~35 K/GPa to yield a 220 melting temperature of ~7000 K at ~120 GPa. The progressive increase of the Clapeyron slope could 221 be related to the 2^{nd} order phase transformation from stishovite to the CaCl₂-type polymorph (Fig. 3), 222 even though the location of this phase boundary in the P-T diagram remains controversial (see (Fischer 223 224 et al. 2018) and references therein). In fact, our X-ray measurements confirm that the subsolidus phase transformation occurs between ~80 and ~100 GPa at ~5400 K (Fig. 4). The change of slope of the 225 226 melting curve is unlikely to come from a difference of volume between the two SiO_2 polymorphs, because it was reported to be insignificant up to ~120 GPa, at least at 300 K (Andrault et al. 2003). In 227 228 contrast, the higher entropy of the CaCl₂-form, compared to stishovite, implies a lower entropy of 229 melting (Δ Sm) above the transition pressure and, consequently, a steeper melting curve (because 230 $dT/dP = \Delta Vm/\Delta Sm$). In addition, the strain energy becomes significant with increasing pressure above 231 the transition. The CaCl₂ distortion of the stishovite-based lattice observed at ~ 107 GPa and 5400 K 232 corresponds to a strain energy of ~2 kJ/mole (Andrault et al. 2003). This value is significant compared 233 to the enthalpy of fusion of quartz of 9.4 kJ/mole (Akaogi et al. 2011). Thus, the stabilization energy 234 associated to the CaCl₂-distortion can modify the Clapeyron slope of the melting curve.

235

3.4 Comparison with previous works

237 Previous LH-DAC experiments suggested an even steeper Clapeyron slope than ours above ~14 GPa (Shen and Lazor 1995). However, after correction of the nominal pressures of the early 238 experiments regarding the pressure increase due to laser heating in the DAC, the two data sets are in 239 fact quite similar to each other. On the other hand, our melting curve plot in relatively good agreement 240 241 with theoretical calculations, except in a few aspects (Fig. 3). The experimental melting slope remains steep, with little pressure-induced flattening, up to pressures higher than reported in theoretical studies 242 (Belonoshko and Dubrovinsky 1995; Luo et al. 2002; Usui and Tsuchiya 2010). When it eventually 243 flattens around 45 GPa and up to 90 GPa, the SiO₂ experimental melting curve is found parallel, but at 244 245 200-300 K higher temperatures, compared to the most recent calculation (Usui and Tsuchiya 2010). 246 The significant increase of melting temperature observed experimentally above ~90 GPa is not 247 reported in any theoretical studies yet.

248 We note an important discrepancy, of e.g. ~1000K for the melting temperature at 45 GPa, among 249 the previous theoretical studies (Belonoshko and Dubrovinsky 1995; Luo et al. 2002; Usui and Tsuchiya 2010). It could come from the difficulty to model properly the local structure in SiO_2 melt, in 250 251 particular the strength of the Si-based structural units in LDM and HDM, respectively. Then, 252 compared to shock compressions, our measurements plot at the upper limit of the Hugoniot plots of (solid) quartz and fused quartz (see (Akins and Ahrens 2002) and references therein). Using this 253 technique, one generally identifies the melting temperature when melting is truly achieved at higher 254 255 impact energies, rather than in the solid at the highest temperatures, before the solid melts. However, this remains a matter of interpretation and the range of temperature uncertainties is large. 256

258 **3.5** Analysis of the SiO₂-melt structure factors

Polyamorphism in liquid SiO₂ has been extensively discussed in the past by theoretical approaches 259 (e.g. (Brazhkin et al. 2011; Lin et al. 2007; San et al. 2016; Takada et al. 2016)). To address 260 experimentally the structural evolution of liquid SiO₂ under high pressures, we performed LH-DAC 261 experiments coupled with the Soller slits system installed at the ID27 beamline (Weck et al. 2013). 262 263 This maximizes the signal of the SiO_2 melt over the noise from the sample environment. No KCl pressure medium was used for these experiments. First, we collected diffraction patterns for the cold-264 compression of SiO₂ glass up to ~45 GPa, in order to validate our experimental measurements of the 265 structure factors S(Q). Our results are found in good agreement with previous studies (e.g. (Prescher et 266 al. 2017)). Spectra present a first sharp diffraction peak (FSDP) around 2 Å⁻¹, a second, less intense, 267 around 5 Å⁻¹ and a new contribution growing with increasing the pressure between 3 and 3.5 Å⁻¹ (Fig. 268 5). These experimental features have been correlated to the gradual compaction of the structure of 269 amorphous SiO₂ (Benmore et al. 2010; Prescher et al. 2017; Sato and Funamori 2010). 270

271 We then recorded diffuse patterns of SiO₂ melt between ~ 25 and ~ 75 GPa, which we normalized to obtain S(Q) (Fig. 6). In comparison to the glass, the same FSDP position is found at higher pressure in 272 the melt (Figs. 7 and 8). In addition, the new diffraction peak starts growing between 40 and 54 GPa, 273 274 and 18.5 and 30.5 GPa, for the melt and the glass, respectively (Fig. 5). In the melt, the shift of the FSDP occurs prior to the appearance of the new diffraction peak, a similar behavior as in the glass 275 276 (Benmore et al. 2010; Prescher et al. 2017; Sato and Funamori 2010). This effect has been associated 277 to the first regime of the compression where the structural changes are dominated by the compaction of the oxygen network (Wu et al. 2012), before Si changes its coordination number at higher pressures. 278 At pressures higher than ~54 GPa, the additional contribution between 3 and 3.5 $Å^{-1}$ appears clearly in 279 280 the signal of SiO₂-melt diffuse scattering. At this point, the position of the FSDP is found above 2.2 Å⁻ ¹ and, at 76 GPa, the melt presents the same FSDP position as the glass at ~45 GPa (Fig. 8). All 281 experimental features are similar to those observed in the glass upon change of Si coordination number 282 from 4 to 6 (Benmore et al. 2010; Inamura et al. 2004; Meade et al. 1992; Prescher et al. 2017; Sato 283 284 and Funamori 2010).

It appears clear that the structural changes occur at higher pressures in the liquid than in the glass. Regarding the extremely high temperatures considered (between 5250 K and 6200 K, Fig. 6), this could be an effect of thermal pressure; the sequence of structural compaction may take place for similar SiO_2 volumes in the glass and in the melt. However, higher temperature requires higher pressure to achieve a similar SiO_2 volume. This hypothesis is supported by the large similarities between diffraction patterns recorded in the liquid and in the glass.

291 292

IV. CONCLUSION

The compaction mechanisms of liquid and amorphous SiO2 are complex and involve various 294 295 atomic processes. Our experimental measurements suggest a change of the SiO₂ melt structure in a relatively narrow range of pressures, within 10-20 GPa, around 45 GPa. Such evolution is compatible 296 with a previous work investigating the Si coordination change from 4 to 6, e.g. in a SiO₂-glass using 297 transverse acoustic wave velocities (Murakami and Bass 2010). In our experiments, the change in the 298 SiO₂ melt structure is recognized based on (i) a change in the Clapeyron slope of the SiO₂ melting 299 curve and (ii) the occurrence of the new diffraction peak between 3 and 3.5 Å⁻¹. The relative sharpness 300 301 of the structure change in the melt suggests a dominant role of the covalent character of Si-O bonds in 302 the melt (Lin et al. 2007), which could induce a significant difference in energy between the 303 tetrahedral "SiO₄" and octahedal "SiO₆" electronic configurations as a function of pressure and temperature. The two different melt structures could be relatively incompatible with each other, in 304 agreement with previous reports (Brazhkin et al. 2011; Lin et al. 2007; Takada et al. 2016)). 305

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308	REFERENCES
309	
310	Akaogi M, Oohata M, Kojitani H, Kawaji H (2011) Thermodynamic properties of stishovite
311	by low-temperature heat capacity measurements and the coesite-stishovite transition
312	boundary Am Miner 96:1325-1330
313	Akins JA, Ahrens TJ (2002) Dynamic compression of SiO ₂ : A new interpretation Geophys
314	Res Lett 29:1394
315 316	Andrault D, Angel RJ, Mosenfelder JL, Le Bihan T (2003) Equation of state of stishovite to lower mantle pressures Am Miner 88:301-307
317	Andrault D, Fiquet G, Guyot F, Hanfland M (1998) Pressure-induced landau-type transition in
318	stishovite. Science 23:720-724
319	Andrault D, Pesce G, Bouhifd MA, Bolfan-Casanova N, Henot JM, Mezouar M (2014)
320	Melting of subducted basalt at the core-mantle boundary Science 344:892-895
321	Belonoshko AB, Dubrovinsky LS (1995) Molecular-dynamics of stishovite melting Geochim
322	Cosmochim Acta 59:1883-1889
323	Benmore CJ, Soignard E, Amin SA, Guthrie M, Shastri SD, Lee PL, Yarger JL (2010)
324	Structural and topological changes in silica glass at pressure Phys Rev B 81
325	Brazhkin VV, Lyapin AG, Trachenko K (2011) Atomistic modeling of multiple amorphous-
326	amorphous transitions in SiO2 and GeO2 glasses at megabar pressures Phys Rev B 83
327	De Nolf W, Vanmeert F, Janssens K (2014) XRDUA: crystalline phase distribution maps by
328	two-dimensional scanning and tomographic (micro) X-ray powder diffraction Journal
329	of Applied Crystallography 47:1107-1117 doi:doi:10.1107/S1600576714008218
330	Dewaele A, Belonoshko AB, Garbarino G, Occelli F, Bouvier P, Hanfland M, Mezouar M
331	(2012) High-pressure high-temperature equation of state of KCl and KBr Phys Rev B
332 333	85 doi:10.1103/PhysRevB.85.214105 Fischer RA et al. (2018) Equations of state and phase boundary for stishovite and CaCl2-type
333 334	SiO2 Am Miner 103:792-802 doi:10.2138/am-2018-6267
335	Geballe ZM, Jeanloz R (2012) Origin of temperature plateaus in laser-heated diamond anvil
336	cell experiments Journal of Applied Physics 111:123518
337	Inamura Y, Katayama Y, Utsumi W, Funakoshi K (2004) Transformations in the
338	intermediate-range structure of SiO ₂ glass under high pressure and temperature Phys
339	Rev Lett 93
340	Lin JF et al. (2007) Electronic bonding transition in compressed SiO ₂ glass Phys Rev B 75
341	Luo SN, Cagin T, Strachan A, Goddard WA, Ahrens TJ (2002) Molecular dynamics modeling
342	of stishovite Earth Planet Sci Lett 202:147-157
343	Meade C, Hemley RJ, Mao HK (1992) High-pressure X-ray diffraction of SiO ₂ glass Phys
344	Rev Lett 69:1387-1390
345	Murakami M, Bass JD (2010) Spectroscopic Evidence for Ultrahigh-Pressure Polymorphism
346 247	in SiO ₂ Glass Phys Rev Lett 104 Petitgirard S et al. (2017) SiO ₂ Glass Density to Lower-Mantle Pressures Phys Rev Lett 119
347 348	Prescher C, Prakapenka VB (2015) DIOPTAS: a program for reduction of two-dimensional
348 349	X-ray diffraction data and data exploration High Pressure Research 35:223-230
350	Prescher C, Prakapenka VB, Stefanski J, Jahn S, Skinner LB, Wang YB (2017) Beyond
351	sixfold coordinated Si in SiO ₂ glass at ultrahigh pressures Proc Natl Acad Sci U S A
352	114:10041-10046
353	San LT, Hong NV, Hung PK (2016) Polyamorphism of liquid silica under compression based
354	on five order-parameters and two-state model: a completed and unified description
355	High Pressure Research 36:187-197

- Sanloup C et al. (2013) Structural change in molten basalt at deep mantle conditions Nature
 503:104-+
- Sato T, Funamori N (2010) High-pressure structural transformation of SiO2 glass up to 100
 GPa Phys Rev B 82
- Schultz E et al. (2005) Double-sided laser heating system for in situ high pressure-high
 temperature monochromatic x-ray diffraction at the ESRF High Pressure Research
 25:71-83
- Shen G, Lazor P (1995) Measurement of melting temperatures of some minerals under lower
 mantle conditions J Geophys Res 100:17699-17713
- Stixrude L, Karki BB (2005) Structure and freezing of MgSiO₃ liquid in the Earth's lower
 mantle Science 310:297-299
- Takada A, Bell RG, Catlow CRA (2016) Molecular dynamics study of liquid silica under high
 pressure Journal of Non-Crystalline Solids 451:124-130
- Usui Y, Tsuchiya T (2010) Ab Initio Two-Phase Molecular Dynamics on the Melting Curve
 of SiO₂ J Earth Sci 21:801-810
- Wang FL, Tange Y, Irifune T, Funakoshi K (2012) P-V-T equation of state of stishovite up to
 mid-lower mantle conditions J Geophys Res: Solid Earth 117:B06209
- Weck G, Garbarino G, Ninet S, Spaulding D, Datchi F, Loubeyre P, Mezouar M (2013) Use
 of a multichannel collimator for structural investigation of low-Z dense liquids in a
 diamond anvil cell: Validation on fluid H₂ up to 5 GPa Rev Sci Instr 84
- Wu M, Liang Y, Jiang J-Z, Tse JS (2012) Structure and Properties of Dense Silica Glass
 Scientific Reports 2
- Zhang JZ, Liebermann RC, Gasparik T, Herzberg CT, Fei YW (1993) Melting and subsolidus
 relations of SiO₂ at 9-14 GPa J Geophys Res: Solid Earth 98:19785-19793

FIGURES AND TABLES

Figure 1: Typical temperature measurements. Thermal emission spectra were recorded at 384 the center of the laser spot for sample temperatures of e.g. 5010 K (left) and 6140 K (right), 385 using the online optical system installed at the ID27 beamline. The raw spectra (upper frames) 386 are corrected by the response of the optical system (itself calibrated with a W-lamp at 2600 K) 387 and fitted using the Wien law (middle frames). The so-called "two-color" methods (lower 388 frames) shows a temperature variation of less than 10% of the mean temperature within a 389 range of wavelength from ~670 to ~870 nm, which establishes the good quality of the thermal 390 emission spectra. Temperature measurements are performed with a precision of ± 100 K and a 391 392 reproducibility better than ± 30 K.

393

394 Figure 2A: X-ray evidences for melting of SiO₂ at ~40 GPa. Time elapses from bottom to 395 top. Temperature uncertainty is +/- 100 K. After crystallization of stishovite, continuous rings 396 remain clearly visible up to more than 5000 K. At 5100 K the intensity of SiO₂ diffraction 397 lines is already significantly decreased. At sample temperatures of 5600 and 5400 K, the only 398 remaining diffraction lines come from the KCl pressure medium. Diffraction lines of 399 stishovite reappear when the laser power is decreased to 5050 K, evidencing the SiO₂ 400 recrystallization.

401

402 Figure 2B: X-ray evidence for the melting of SiO₂ at 105 GPa. Time elapses from top to 403 bottom. For temperatures lower than 6200 K, the set of diffraction lines evidences the 404 presence of the CaCl₂-form of SiO₂ (see Fig. 4) in addition to the KCl pressure medium. The 405 intensity of the SiO₂ peaks decreases severely at 6300-6400 K before they almost disappear at 406 6500 K. The stishovite to KCl ratio of diffraction peaks intensity decreases by a factor of 10 407 around between 6200 and 6500 K.

408

409 <u>Figure 3</u>: Melting curve of SiO₂ up to ~120 GPa. We report temperatures of solid (blue) and 410 liquid (red) SiO₂ together with a smooth trend through the data set (grey band). Our 411 temperature and pressure uncertainties are 200 K and 5 GPa, respectively. We also plot 412 previous melting curve determinations based on early experimental works (Shen and Lazor 413 1995; Zhang et al. 1993) (light blue), theoretical calculations (purple (Usui and Tsuchiya 414 2010), blue (Belonoshko and Dubrovinsky 1995) and green (Luo et al. 2002)) and shock wave

- 415 experiments (orange (Akins and Ahrens 2002)). For the latter, we report Hugoniot plots in416 solid and molten states for starting materials composed of quartz and fused-quartz.
- 417

Figure 4: Subsolidus transition from Stishovite to the CaCl₂-form between 75 and 102 GPa: The stishovite model fits well the upper diffraction pattern recorded at ~75 GPa and ~4500 K. In contrast, the diffraction pattern recorded at ~102 GPa and ~5400 K is not well fitted with the stishovite model (middle frame), but instead with the CaCl₂-form model (lower pattern). At these conditions, the CaCl₂-form presents (a, b, c) unit cell parameters of (3.831(6), 3.967(5), 2.543(3)). Black, red and blue profiles are experimental data, background and Rietveld fit, respectively.

425

426 **Figure 5:** Structure factor of the compressed glass: As previously noticed (e.g. (Prescher et 427 al. 2017) and references therein), a new contribution around 3-3.5 Å^{-1} grows above ~20 GPa.

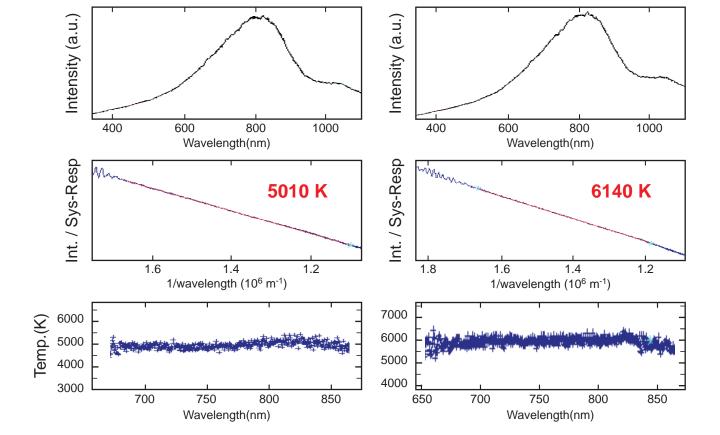
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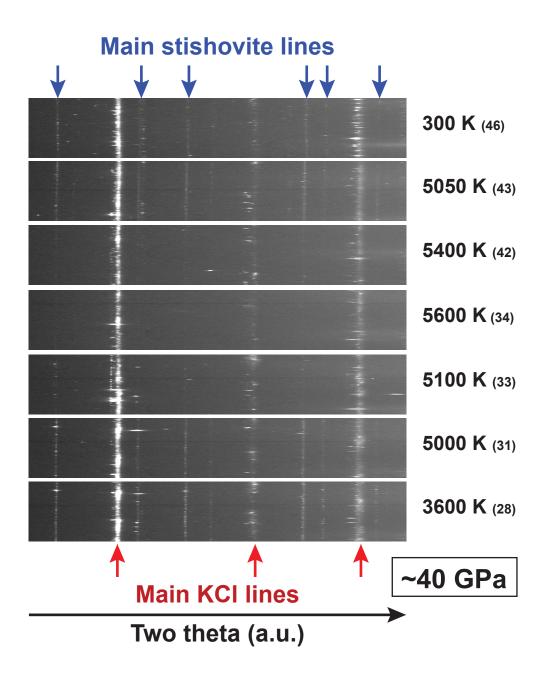
Figure 6: Structure factors S(Q) of SiO₂ recorded in the melt between ~38 and ~76 GPa.
Spectra recorded at high temperature in the melt at pressures of 38 GPa (orange), 40 GPa
(bleu), 54 GPa (grey) and 76 GPa (yellow) are bracketed between spectra recorded at 300 K
in the glass compressed to ~18.5 GPa (green) and ~49.5 GPa (dark blue). The contributions
observed for the melt below and above ~45 GPa are comparable to those of the glass at ~18.5
GPa and ~49.5 GPa, respectively. It indicates an evolution of the SiO₂ melt structure between
40 and 54 GPa comparable to that reported during the compression of the glass (Fig. 5).

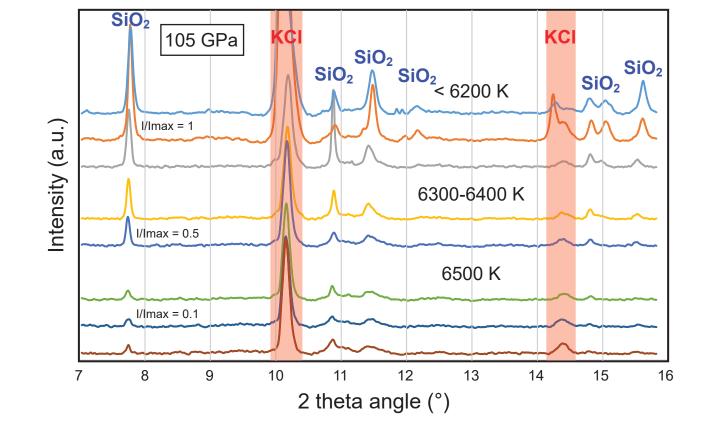
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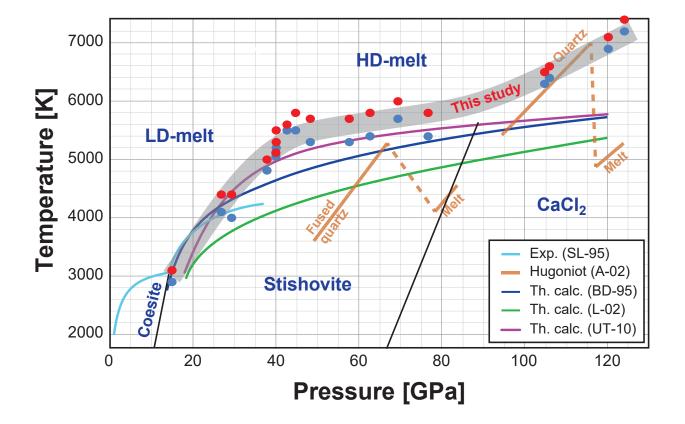
Figure 7: Diffuse scattering signals of melt and quenched glass. The quenched glass (orange, recorded at 8 GPa and 300 K) presents a FSDP position significantly sharper and at slightly higher Q position compared to the corresponding melt (blue, recorded at 26 GPa and ~5150 K). This suggests less densification of the local structure of SiO₂ melt at high temperature, compared to the glass. We did not successfully produced a glass at experimental pressures exceeding 45 GPa, because of full sample crystallization.

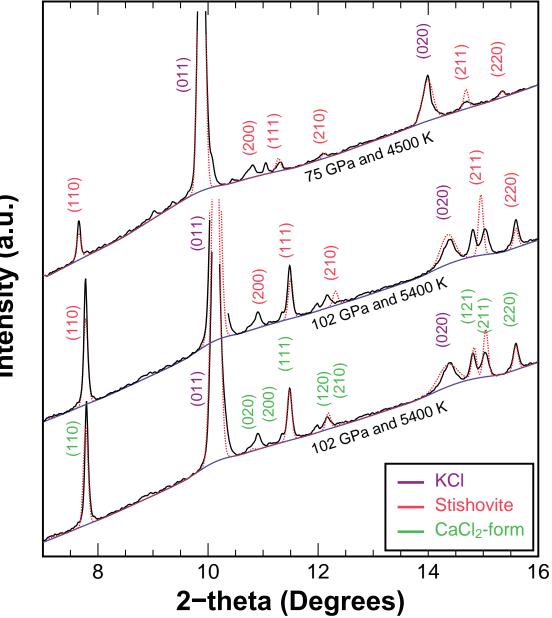
- 445 <u>Figure 8</u>: Pressure dependence of the first sharp diffraction peak. Our measurements 446 performed in melts (Red) and in glasses either quenched from high temperature (Blue) or 447 during cold-compression (Green) are compared with previous reports. The shift of the FSDP 448 to high wavenumbers occurs at lower pressures in the glass, compared to the melt.
- 449
- 450 <u>**Table 1. Melting temperatures:**</u> Sample temperatures are directly provided by radio-451 spectrometry (Fig. 1). Sample pressures are derived from cross correlation between two 452 methods: the P-V-T equation of state of stishovite and an empirical model of thermal pressure 453 (see text). Experimental uncertainties are within 5 GPa and 100 K.
- 454
- 455 <u>Table 2.</u> Clapeyron relations: Volumes and entropies relations calculated around (i) the true
 456 triple point where coesite, stishovite and LDM coexist and (ii) the fictive triple point where
 457 LDM, HDM and stishovite would coexist (see text). Thermodynamical parameters come from
 458 (Akaogi et al. 2011; Wang et al. 2012; Zhang et al. 1993).
- 459



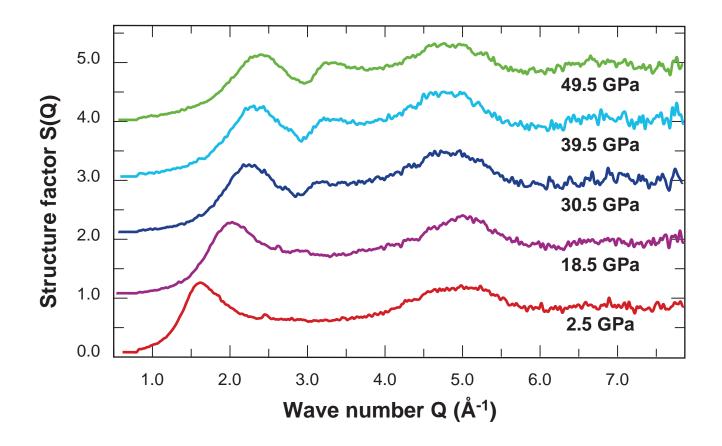


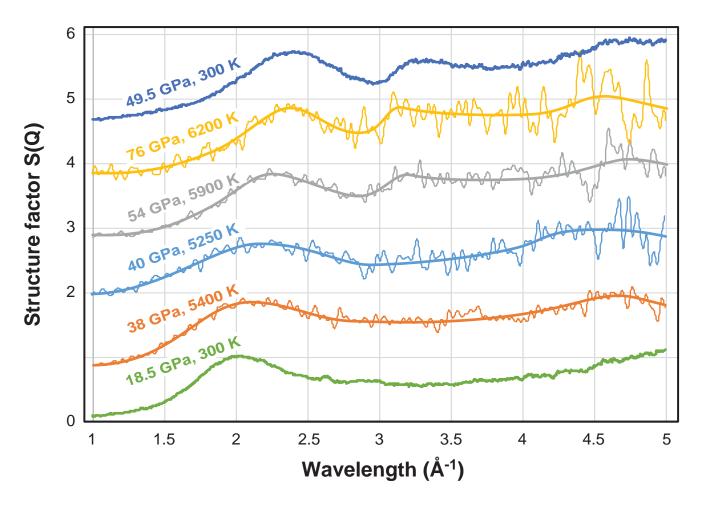


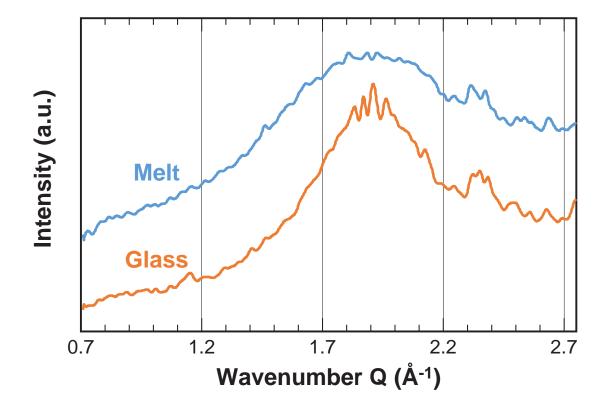


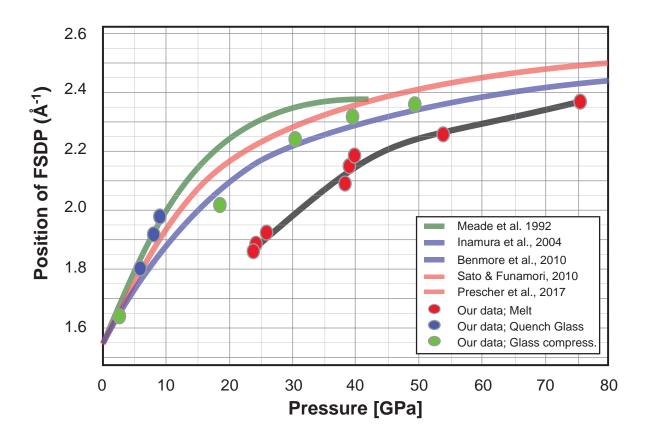


Intensity (a.u.)









Transition	Press.	Temp.	dT/dP	DS	DV	Vstish	Vcoes	VLDM	VHDM	DV/V
considered	GPa	Κ	K/GPa	mol/K	cm3/mol	cm3/mol	cm3/mol	cm3/mol	cm3/mol	
C => S	15	5 3023	312 a	-15.28 a	-4.76736 b	14.6916 g	19.459 h			-0.279 k
$C \Longrightarrow LDM$	15	5 3023		74.9 d	0 e		19.459 h			
S => LDM	15	5 3023	80 *	59.59 b	4.76736 g	14.6916 g		19.459 i		0.2792 k
S => HDM	45	5 5400	12.5 *	74.59 d	0.9324 b	13.6434 g			14.5758 j	0.0661 k
LDM=>HDM	45	5 5400		-15 c	-3.83496 f			18.4108 i	14.5758 j	-0.233 k
S => HDM	45	5 5400	12.5 *	59.59 d	0.7449 b	13.6434 g			14.3883 j	0.0531 k
LDM=>HDM	45	5 5400		0 c	-4.02246 f			18.4108 i	14.3883 j	-0.245 k
S => HDM	45	5 5400	12.5 *	44.59 d	0.5574 b	13.6434 g			14.2008 j	0.04 k
LDM=>HDM	45	5 5400		15 c	-4.20996 f			18.4108 i	14.2008 j	-0.258 k

C, S, LDM and HDM stand for Coesite, Stishovite, low-density and high-density SiO2 melts

*, this study

a, from Akaogi et al. (2011)

b, calculated based on the Clapeyron relation $dT/dP{=}DV/DS$

c, value assumed to be -15, 0 or 15 J/molK (see text)

d, derived from Clapeyron relations around the (C, S, LDM) or (C, LDM, HDM) triple points

e, derived from the flat melting curve of coesite between 9 and 14 GPa (Zhang et al., 1993)

f, calculated using DV=VHDM-VLDM

g, calculated at the considered P-T conditions using the EoS of Wang et al. (2012)

h, calculated using Vcoes=Vsti+DV(C-S)

i, calculated using VLDM=Vsti+DV(LDM-S)

j, calculated using VHDM=Vsti+DV(HDM-S)

k, DV of the transition normalized to the mean value between the two phases considered

P(300K)	Temp.	P(high-T)	Temp	Temp
GPa	К	GPa	solid	Melt
12	3000	15	2900	3100
16	4250	27	4100	4400
19	4200	29	4000	4400
19	5350	40	5200	5500
19.5	5100	40	5050	5100
20	5500	43	5500	5600
21	5600	45	5500	5800
21.5	4900	38	4800	5000
21.5	5100	40	5000	5300
26	5500	48	5300	5700
36	5500	58	5300	5700
39	5700	63	5400	5800
45	5800	69	5700	6000
54	5700	77	5400	5800
76	6500	106	6400	6600
76	6400	105	6300	6500
85	7300	124	7200	7400
85	7000	120	6900	7100